

Ritual Violence and the Perception of Social Difference:
Migration and Human Sacrifice in the Epiclassic Basin of Mexico

by

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ABSTRACT

Archaeologists have long contended that large-scale human migrations played an essential role in the cultural development of pre-Hispanic central Mexico. During the Epiclassic period (600-900 CE), migration is implicated in the appearance of new forms of material culture, sociopolitical disruptions, and the emergence of new regional polities. Sweeping social changes accompanied these developments, including demographic reorganization and increased levels of violence. Research across the social sciences finds that violence directed at individuals perceived as categorically distinct also typically increases during such periods of socio-political upheaval. This dissertation investigates identity-based violence in the Epiclassic Basin of Mexico to consider how diverse social identities contributed to the selection of victims of ritual violence.

This research examines the skeletal remains from a sacrificial deposit at the Epiclassic shrine site of Non-Grid 4 in the Basin of Mexico, where a minimum of 180 human crania were interred as ritual offerings. The project reconstructs patterns of paleomobility and biological relatedness to determine whether individuals with distinct categorical social identities were more likely to become victims of human sacrifice. It answers the questions: (1) Were the sacrificed individuals predominantly locals who lived in the Basin of Mexico throughout their lives?; (2) Were the sacrificed individuals comprised of a single kin-group biologically continuous with pre-extant populations in the Basin of Mexico?; and (3) If victims were migrants biologically discontinuous with antecedent populations, from where in ancient Mesoamerica did they originate?

Results indicate that a majority of sacrificial victims were immigrants originating north and south of the Basin of Mexico. Biogeochemical analyses of sacrificed

individuals find that 80% are non-local migrants into the Basin, suggesting that they were likely targeted for violence based on their divergent residential histories. Multi-scalar biodistance analyses of Non-Grid 4 sacrificial victims demonstrate that they represent two biologically distinct groups. There was evidence, however, for both biological continuity among victims and pre-extant central Mexican populations, as well as for migration from northern and southern Mexico. This project therefore not only improves knowledge of migration during the central Mexican Epiclassic, but also contributes to broader anthropological understandings of the social context of violence.

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I decided I wanted to be an archaeologist when I was seven years old. My family lived in Mexico at the time. We were on a road trip, driving up a narrow winding dirt road through dense jungle before finally arriving. Calakmul. No one was around to direct us, so we started toward a pyramid that was half-swallowed by vines and trees. We didn't realize there were on-going excavations. I remember climbing up the tall steps like a spider monkey, all limbs and exuberance. As I advanced higher and higher, I noticed the rungs of a ladder peeking out from a small rectangular opening in the top of the pyramid. Peering over the edge and into the darkness below, my eyes struggled to make sense of what they were seeing. I was looking inside the pyramid. It was as simple as that. I was sold.

Despite having decided on a career path rather early in life, I hadn't figured out many details beyond that. It wasn't until taking a forensic osteology course with Jill Shapiro, a biological anthropologist at Columbia University, that things began to settle into place for me. I was captivated by the direct window into past lives that human skeletal remains provided. From the first class I took with her, Jill's generosity and mentorship encouraged me to pursue my research interests in bioarchaeology and apply for graduate school. I would not have made it this far without her. Thank you, Jill.

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TABLE OF CONTENTS

| | Page |
|---|------|
| LIST OF TABLES | x |
| LIST OF FIGURES | xi |
| CHAPTER | |
| 1 INTRODUCTION | 1 |
| Migration in Ancient Mesoamerica..... | 3 |
| Migration in Epiclassic Central Mexico..... | 4 |
| Research Orientation..... | 6 |
| Structure of the Dissertation | 8 |
| 2 EXPANDING RADIOGENIC STRONTIUM ISOTOPE BASELINE DATA FOR CENTRAL MEXICAN PALEOMOBILITY STUDIES | 12 |
| Strontium Isotopes in Studies of Paleomobility | 13 |
| Strontium Isotope Systematics | 13 |
| Strontium Isotopes and Paleomobility across Mesoamerica | 15 |
| Central Mexican Geography, Geology, and Geochemistry | 15 |
| The Basin of Mexico in Geological Context..... | 19 |
| Materials and Methods..... | 20 |
| Biogeochemical Methods | 22 |
| Analytical Methods | 24 |
| Results and Discussion | 25 |
| Conclusion..... | 33 |
| References | 34 |

| CHAPTER | Page |
|---------|---|
| 3 | MIGRATION, VIOLENCE, AND THE "OTHER": A BIOGEOCHEMICAL APPROACH TO IDENTITY-BASED VIOLENCE AND HUMAN SACRIFICE IN THE EPICLASSIC 44 |
| | Social Boundaries and Identities45 |
| | Theorizing the Link between Social Identity and Violence46 |
| | Geographic Origins and Social Difference in Ancient Mesoamerica49 |
| | Epiclassic Central Mexico and the Non-Grid 4 Shrine Site53 |
| | Ritual Violence at Non-Grid 4.....54 |
| | Human Sacrifice and Identity-Based Violence58 |
| | Reconstructing Paleomobility through Biogeochemistry60 |
| | Strontium Isotope Baseline Variability in Central Mexico.....61 |
| | Oxygen Isotope Baseline Variability in Central Mexico.....64 |
| | Strontium and Oxygen Isotope Sources in the Basin of Mexico.....65 |
| | Materials67 |
| | Methods68 |
| | Radiogenic Strontium Isotope Analysis69 |
| | Stable Oxygen Isotope Analysis71 |
| | Results73 |
| | Examining Diagenetic Contamination at Non-Grid 474 |
| | Residential Histories of Sacrificial Victims at Non-Grid 475 |
| | Identity-Based Violence at Non-Grid 4.....80 |
| | Conclusion.....82 |

| CHAPTER | Page |
|--|------|
| References | 84 |
| 4 MIGRATION AND BIOLOGICAL CONTINUITY IN CENTRAL MEXICO DURING THE CLASSIC-EPICLASSIC TRANSITION..... | 109 |
| Migration in Epiclassic Central Mexico..... | 110 |
| Archaeological Evidence of Migration | 111 |
| Bioarchaeological Evidence of Migration..... | 113 |
| Examining Epiclassic Migrations through Biodistance Analysis..... | 114 |
| Migration at the Non-Grid 4 Epiclassic Shrine Site..... | 116 |
| Materials and Methods..... | 118 |
| Finite Mixture Analysis at Non-Grid 4 | 124 |
| Relationship (R) Matrix Analysis..... | 125 |
| Results | 126 |
| Discussion | 131 |
| Continuity or Replacement in Epiclassic Central Mexico?..... | 133 |
| Epiclassic Migrant Origins at Non-Grid 4 and Beyond..... | 134 |
| Conclusion..... | 137 |
| References | 139 |
| 5 CONCLUSION | 151 |
| Summary of Results..... | 155 |
| Final Thoughts and Future Directions | 159 |
| REFERENCES | 162 |

| APPENDIX | Page |
|---|------|
| A AUTHOR CONTRIBUTIONS | 204 |
| B CENTRAL MEXICAN BASELINE $^{87}\text{SR}/^{86}\text{SR}$ DATA | 206 |
| C BIOGEOCHEMICAL $^{87}\text{SR}/^{86}\text{SR}$ AND $\delta^{18}\text{O}$ DATA FROM NON-GRID 4 | 212 |

LIST OF TABLES

| Table | Page |
|--|------|
| 1. $^{87}\text{Sr}/^{86}\text{Sr}$ Medians and Interquartile Ranges for Central Mexican Subregions Identified through K-means Cluster Analysis. | 27 |
| 2. Central Mexican Site-Level $^{87}\text{Sr}/^{86}\text{Sr}$ Medians and Interquartile Ranges | 32 |
| 3. Demography of Non-Grid 4 Sacrificial Victims | 58 |
| 4. Sites and Sample Sizes of Individuals Included in the Study..... | 119 |
| 5. Finite Mixture Analysis Top <i>BIC</i> Values and <i>k</i> Group Solutions for <i>m</i> Imputed Datasets..... | 126 |
| 6. Pooled Inter-site Estimated Genetic Distances and Associated Standard Errors... | 128 |
| 7. Z-scores of Inter-site Estimated Genetic Distances and Associated <i>p</i> -values | 129 |

LIST OF FIGURES

| Figure | Page |
|--|------|
| 1. Morphotectonic Provinces of Mexico | 17 |
| 2. Cluster $^{87}\text{Sr}/^{86}\text{Sr}$ Values in Plant Samples by Plant Origin | 26 |
| 3. Sampled Sites within Central Mexico Sorted by Cluster Membership | 27 |
| 4. Medians and Interquartile Ranges of Cluster Subregions | 28 |
| 5. $^{87}\text{Sr}/^{86}\text{Sr}$ Interquartile Ranges of Central Mexican Sites, Shaded by Cluster | 31 |
| 6. The Non-Grid 4 Shrine Site and pre-Hispanic Highland Lake System in Central Mexico..... | 56 |
| 7. REE and U Concentrations (ppm)/MTC for all Analyzed Human Samples | 75 |
| 8. $^{87}\text{Sr}/^{86}\text{Sr}$ and $\delta^{18}\text{O}$ Results across All Sampled Human Elements ($N=194$) | 76 |
| 9. Intra-individual Mobility among Sacrificial Victims Interred at Non-Grid 4 | 79 |
| 10. Sites and Regions Included in the Analysis | 120 |
| 11. Density Plots of Imputed Data by Variable | 123 |
| 12. Multi-dimensional Scaling of Inter-site Estimated Genetic Distances | 131 |

CHAPTER 1

INTRODUCTION

Human migrations can be powerful catalysts for political, economic, and social change. Here, migration is understood as a process involving the long-term movement and relocation of people across significant socio-cultural, political, or environmental boundaries (Tsuda et al., 2015). Such migrations are ubiquitous across human history (Baker & Tsuda, 2015; Cabana & Clark, 2011; Crawford & Campbell, 2012), and recent studies find evidence of evolutionary origins for the practice (B. C. Campbell & Barone, 2012). While contemporary migration occurs on a greater scope and scale than ancient migration, it remains an acceleration of the same process (Burmeister, 2000; B. C. Campbell & Crawford, 2012; Manning, 2006). As such, archaeologists have used causes and patterns evident in modern migrations as models to better recognize and understand the social dynamics of ancient population movements (Anthony, 1990; Burmeister, 2000; Cameron, 2013; Clark et al., 2019).

Studies of contemporary migrations revolve around the motivation behind and consequences of migration (Tsuda, 2011, p. 315). Motivations for past migrations, however, can be difficult to discern from the archaeological record. Archaeologists have reconstructed environmental factors, such as droughts (Dillehay, 2002; Hoggarth et al., 2017; Lachniet et al., 2012, 2017) and natural disasters (González et al., 2000; Plunket & Uruñuela, 2006; Siebe, 2000) that may have pushed individuals to emigrate from their origin communities. Such environmental factors, however, cannot be invoked as monolithic explanations of complex social reactions spurring migration processes

(Beekman, 2015, p. 79). Indeed, often patterns of cultural development and migration do not coincide with unfavorable environmental conditions (e.g., Elliott et al., 2010; Frederick & Cordova, 2018; Knudson & Torres-Rouff, 2015).

Moreover, environmental conditions are not the only factors motivating migration. Long-term studies of multi-scalar population movements suggest that social factors are among the most important in motivating individual and collective decisions to migrate from their origin communities (Hamnett, 1985). Reviews of ethnographic and ethnohistoric literature of more recent migrations indicate that factors such as social hierarchies, kinship obligations, and warfare all play an essential role in informing mobility decisions (e.g., Cameron, 2013; Schachner, 2012). Accessing these factors archaeologically, however, can be challenging given both the imprecise nature of emigration chronologies and difficulties locating migrant origin communities in the past.

In contrast, archaeology has had greater success investigating the consequences of past migrations. Resultant changes in social organization and past social dynamics in receiving communities are more readily reconstructed through the archaeological record (Anthony, 1990; Burmeister, 2000). Many studies have found that migration has both generative impacts, such as the formation of multicultural societies with expanded socioeconomic networks and the revitalization of host societies through cultural exchange, as well as disruptive impacts, such as social exclusion and factionalism, violence, and even political decline (Beekman & Christensen, 2011; J. H. Bentley, 1993; Cowgill, 2015a; Hills, 2015; Manzanilla, 2015; Van De Mieroop, 2011, pp. 127–149; Zakrzewski, 2015). These studies contradict broadly polarized characterizations of the

impact of migration on receiving communities. Rather than clearly supporting oversimplified models of either total assimilation or cultural replacement, archaeological studies of migration demonstrate that migration engenders a complex social landscape where migrant enclaves, zones of cultural hybridization, and areas of local resistance may all represent responses to the same migration event (Clark, 2001).

Migration in Ancient Mesoamerica

A large body of archaeological research on ancient migrations within Mesoamerica has produced evidence of both generative and disruptive impacts (Beekman, 2019; Kelley, 2000; Manzanilla, 2005b; Ragsdale & Edgar, 2018). Migration of elites has been associated with the rise of centralized hierarchical leadership at sites across the Maya region (Price et al., 2010; Stuart, 2000; White et al., 2000; Wright, 2012). Similarly, migration is often implicated in the emergence of new forms of socio-political organization, including the urbanized multi-ethnic states of southern and central Mexico (Beekman, 2015; Feinman & Nicholas, 2013, pp. 63–64; Manzanilla, 2015; Overholtzer & De Lucia, 2016). In particular, at the central Mexican Classic period (250–600 CE) urban center of Teotihuacan, migration has been linked to the creation and maintenance of urban neighborhood districts (Manzanilla, 2015; Rattray, 1989; Smith, 2010), as well as the expansion of socioeconomic exchange networks (Nichols, 2016; Rodríguez Galicia & Valadez Azúa, 2013). Additionally, archaeologists have found evidence that migration also accompanied the spread of novel forms of material culture (Foster, 1999; Hernández & Healan, 2019; Rattray, 1966; Tozzer, 1921), hybrid art styles

(McVicker, 1985; Nagao, 1989, 2014; Turner, 2019), and religious cults (Ringle et al., 1998).

Additionally, there is evidence suggesting that migration within ancient Mesoamerica led to both increased social integration as well as exclusion. Studies of Teotihuacan residential compounds indicate that while some migrants assimilated into the city's multiethnic society (Manzanilla, 2017; White, Storey, et al., 2004), others maintained social boundaries, living in ethnic enclaves (Begun, 2013; Rattray, 1989; Spence, 1992). These patterns of multiculturalism and integration are replicated beyond Teotihuacan as well (Castañeda de la Paz, 2013; Healan & Cobean, 2019; Holt Mehta, 2018; Smith & Lind, 2005; Spence, 2005). However, iconographic and bioarchaeological evidence suggests that, in some cases, periods of increased migration also coincided with higher levels of inter-personal violence (Baird, 1989; Finegold, 2012; Morehart, 2017a; Morehart et al., 2012) and may have ultimately contributed to regional socio-political upheaval (Cowgill, 2015a; Manzanilla, 2015).

Migration in Epiclassic Central Mexico

Despite broader acknowledgement of the importance of migration within the cultural development of Mesoamerica, the role of migration within the socio-political transformation of central Mexico during the Classic (250-600 CE) to Epiclassic (600-900 CE) transition has long been debated (Cowgill, 2013). The Epiclassic period is characterized by the dramatic decline of Teotihuacan's political and economic control of the region (J. H. Anderson et al., 2016; Cowgill, 2015b; Diehl & Berlo, 1989; Healan & Cobean, 2019, p. 82), and the emergence of new competing political centers within

central Mexico (García Cook, 2013; Hirth, 2000; Mendoza, 1992). Additionally, the period is marked by the proliferation of militaristic iconography (Finegold, 2012; Hirth, 1989; Koontz, 1994; McVicker, 2007; Nagao, 1989; Ringle et al., 1998), suggesting increased conflict accompanying regional political decentralization during this period.

The Epiclassic also witnessed a large influx of migrants into the region. A novel ceramic complex known as Coyotlatelco appeared in central Mexico during this time period (Crider, 2011; Rattray, 1966; Sánchez, 2013; Solar Valverde, 2006; Tozzer, 1921). Most scholars agree that Coyotlatelco ceramics represent a distinctive break with local central Mexican ceramic traditions (but see Sanders, 1986, 2006; Sugiura Yamamoto, 2006). Instead, Coyotlatelco's strong resemblance to earlier ceramic industries from elsewhere in Mexico suggested that migrants from northwestern Mexico introduced the ceramic complex to central Mexico (Beekman & Christensen, 2003; Hernández, 2016; Hernández & Healan, 2019; Rattray, 1966, 1989, 1996). Combined with the marked demographic reorganization of central Mexico during the Epiclassic (Parsons et al., 2008; Sanders et al., 1979) and shifts in regional settlement patterns from valley floor to dispersed hilltop settlements (J. H. Anderson et al., 2016; Gorenflo & Sanders, 2007; Healan & Cobean, 2019; Hirth, 2000; Mastache et al., 2002; Morehart, 2016a; Parsons et al., 2008), the prevailing consensus among archaeologists is that substantial migration into central Mexico took place during the Epiclassic period.

Bioarchaeological evidence also supports the arrival of migrants into central Mexico during the Epiclassic period. A biogeochemical analysis of Epiclassic burials from Teotihuacan directly identified three immigrants at the site (Price et al., 2000).

Further applications of biogeochemical analyses to identify Epiclassic migrants in the region, however, have been hindered by the paucity of central Mexican skeletal collections dating to the Epiclassic. Other lines of evidence, including both ancient DNA (Aguirre-Samudio et al., 2017; Álvarez-Sandoval et al., 2015; Morales-Arce et al., 2019) and biodistance analyses of phenotypic variation in skeletal and dental traits (Beekman & Christensen, 2003; Christensen, 1997; Gómez-Valdés et al., 2008; González-José et al., 2007; Meza-Peñaloza et al., 2019; Ragsdale & Edgar, 2018) find evidence of biological discontinuity between central Mexican Classic and Postclassic (900-1521 CE) populations. Although few ancient DNA or biodistance analyses examined skeletal remains dating to the Epiclassic, their results are suggestive of an Epiclassic population replacement within central Mexico.

Research Orientation

Thus, although a great deal of effort has been spent establishing the presence of migrants within Epiclassic central Mexico, relatively little research has been done on the migrants themselves (but see Manzanilla, 2005a). Who were these migrants? From where in ancient Mesoamerica did they originate? What happened to them once they arrived in central Mexico?

This dissertation complements previous research on Mesoamerican migration, using multiple lines of evidence to characterize migration at Non-Grid 4, an Epiclassic shrine site in central Mexico where a minimum of 180 individuals were ritually sacrificed and interred (Morehart et al., 2012). Previous studies of these individuals found they were biologically discontinuous with antecedent populations in central Mexico (Meza-

Peñaloza et al., 2019). Furthermore, the sacrificed individuals were not genetically homogenous, but instead represented multiple biological kinship groups (García Velasco, 2019). Here, biogeochemical analyses are used to directly identify first-generation migrants among the sacrificial assemblage, and biodistance methods are used to posit potential origin communities of identified migrants.

Furthermore, although ethnohistoric sources indicate that human sacrifice was a culturally sanctioned form of violence practiced in ancient Mesoamerica (Graulich, 2006; López Luján & Olivier, 2010; Sahagún, 1951), bioarchaeological analyses of victims of ritual violence throughout Mesoamerica indicate that they do not represent a random cross-section of the population at-large (Crandall & Thompson, 2014; De La Cruz et al., 2008; Duncan, 2011; Kieffer, 2017; Moreiras Reynaga, 2019). As such, this dissertation develops a theoretical framework to examine identity-based violence—violence perpetrated against individuals or groups who are perceived as categorically distinct. This research thus examines how individuals with particular categorical social identities were selected or targeted for violence during the Epiclassic period, examining the violent consequences of migration on the migrants themselves.

The goal of this research is to reconstruct Epiclassic central Mexican migration patterns through the analysis of paleomobility and biological relatedness at the Epiclassic ritual shrine site of Non-Grid 4. Additionally, this project examines the interplay between perceptions of social difference and ritual violence, seeking to determine whether individuals with distinct categorical social identities were more likely to become victims of human sacrifice. In the volatile geopolitical landscape of Epiclassic central Mexico,

aspects of individuals' social identities—including their residential histories or biological kinship affiliations—could have acted as powerful indicators of social difference that culminated in ritual violence.

A bioarchaeological approach is particularly well suited to investigate migration and identity-based violence at Non-Grid 4. Many facets of identity are grounded in the physical body, such as age, sex, or phenotypic expressions of genetic relatedness. Other aspects of identity leave traces on the body through culturally mitigated behaviors, including dietary preferences or various forms of intentional body modifications (Buikstra & Scott, 2009; Knudson & Stojanowski, 2008, 2009). These corporeal embodiments of multivalent social identities allow for the empirical reconstruction of geographic origins and residential histories using biogeochemical analysis, as well as the characterization of biological affinities through biodistance analysis among individuals at Non-Grid 4. Furthermore, they allow for a nuanced consideration of whether particular social identities were predisposed to suffer ritual violence in Epiclassic central Mexico.

Structure of the Dissertation

This dissertation employs a bioarchaeological approach to examine migration and identity-based violence in Epiclassic central Mexico. It is organized into five chapters. The second chapter presents a biogeochemical baseline for bioavailable radiogenic strontium ($^{87}\text{Sr}/^{86}\text{Sr}$) values across central Mexico. Radiogenic strontium isotope values vary according to the age and composition of local underlying geologic bedrock (R. A. Bentley, 2006; Faure & Powell, 1972). Examining $^{87}\text{Sr}/^{86}\text{Sr}$ values in skeletal tissues mineralizing at different times over the life course allows for the reconstruction of past

individuals' movements between geologically distinct zones (R. A. Bentley, 2006; Ericson, 1985; Flockhart et al., 2015). However, expected "local" $^{87}\text{Sr}/^{86}\text{Sr}$ values are essential to the successful implementation of such techniques to characterize past mobility.

While $^{87}\text{Sr}/^{86}\text{Sr}$ has been used widely in the study of paleomobility within ancient central Mexico, the majority of work in the region has focused on the site of Teotihuacan (Price et al., 2000; Solís Pichardo et al., 2017; White et al., 2007). This chapter expands $^{87}\text{Sr}/^{86}\text{Sr}$ baseline data throughout central Mexico. It presents generated ($n=63$) and previously published ($n=16$; Price et al., 2008; Schaaf et al., 2012) $^{87}\text{Sr}/^{86}\text{Sr}$ values from 13 sites across central Mexico to establish a k-means cluster model determining which regions within central Mexico can and cannot be distinguished using radiogenic strontium isotopes.

The third chapter establishes a theoretical framework within which to examine identity-based violence in the past using bioarchaeological methods. Drawing on social science research on the development of in-group preferences and out-group biases (Riek et al., 2006; Tajfel et al., 1971), as well as realistic group conflict theory from social psychology (Jackson, 1993; Sherif, 1966), and empirical data from cross-cultural anthropological, evolutionary psychology, and behavioral case studies (Böhm et al., 2016; Schmidt & Schröder, 2001; Simunovic et al., 2013), the chapter establishes a link between social group formation and identity-based violence in conditions of heightened scarcity and intergroup competition. Bioarchaeological methods are particularly adept at reconstructing multivalent past social identities, but few studies have applied such an

identity-based approach to understand social targeting in acts of violence (but see Kurin, 2014, 2016).

After establishing the cultural salience of distinct geographic origins as an important signifier of social difference within pre-Hispanic central Mexico, the chapter then turns to examine the residential histories of a sample of sacrificed individuals at Non-Grid 4 ($n=73$). Biogeochemical analyses of radiogenic strontium ($^{87}\text{Sr}/^{86}\text{Sr}$) and stable oxygen ($\delta^{18}\text{O}$) isotopes in enamel and bone tissues are used to reconstruct sacrificial victims' residential histories and identify first-generation immigrants. Paleomobility patterns among sampled individuals are then analyzed to discern if and how identity-based violence affected individuals interred at the site.

The fourth chapter examines patterns of biological affinity at Non-Grid 4 and greater Mesoamerica during the Classic-Epiclassic transition. It reviews previous archaeological and bioarchaeological models of Epiclassic migrations into central Mexico to establish testable models and expectations for the subsequent analyses. The chapter investigates patterns of biological affinity (1) among Non-Grid 4 individuals, as well as (2) between Non-Grid 4 individuals and antecedent central Mexican populations, and more broadly (3) between Non-Grid 4 individuals and other Epiclassic populations throughout Mesoamerica.

Multi-scalar biodistance analyses are used to characterize biological diversity at Non-Grid 4 and Epiclassic central Mexico and locate potential origin communities for Epiclassic migrants. Finite mixture analysis (McLachlan & Peel, 2000) of cervicometric tooth dimensions is used to characterize patterns of biological variability among Non-

Grid 4 individuals ($n=115$). Relationship (R) matrix analysis (Relethford et al., 1997; Relethford & Blangero, 1990) is then used to compare Non-Grid 4 individuals with a broader skeletal sample of individuals from across Classic and Epiclassic Mesoamerica ($n=218$) to evaluate archaeological models of biological continuity and population replacement within Epiclassic central Mexico.

The final chapter presents a summary of the dissertation and its conclusions. It evaluates whether the dissertation successfully achieved its stated aims of reconstructing migration patterns among the individuals interred at Non-Grid 4 and examining the impact of that migration on the migrants themselves using an identity-based violence framework. The dissertation closes with a consideration of how future research can contribute to increasingly nuanced research into both the causes and consequences of migration as well as identity-based violence in the past.

CHAPTER 2

EXPANDING RADIOGENIC STRONTIUM ISOTOPE BASELINE DATA FOR CENTRAL MEXICAN PALEOMOBILITY STUDIES

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Researchers have long debated the importance of migration in the cultural development of central Mexico. A number of archaeological (Beekman & Christensen, 2003; Begun, 2013; Millon, 1973; Spence, 1992), morphological (Christensen, 1997; González-José et al., 2007; Meza-Peñaloza et al., 2019; Ragsdale, 2017; Ragsdale & Edgar, 2018), and genetic (Aguirre-Samudio et al., 2017; Kemp et al., 2005), analyses indicate that the Basin of Mexico attracted multiple waves of migrants from across greater Mesoamerica throughout pre-Hispanic times. Biogeochemical studies of radiogenic strontium ($^{87}\text{Sr}/^{86}\text{Sr}$) isotopes have proven effective in directly testing the presence of migrants within the Basin, particularly at the Classic period city of Teotihuacan (Manzanilla et al., 2012; Nado, 2017; Price et al., 2000; Schaaf et al., 2012; Solís Pichardo et al., 2017; White et al., 2007). While determining “local” ranges of variation in $^{87}\text{Sr}/^{86}\text{Sr}$ values is essential for the further application of this method, central Mexican radiogenic strontium data outside of Teotihuacan remain limited. Price and colleagues (Price et al., 2008) established a regional expected $^{87}\text{Sr}/^{86}\text{Sr}$ range for the Basin of Mexico as a whole, but no studies examine $^{87}\text{Sr}/^{86}\text{Sr}$ variability within the Basin or central Mexico.

This study investigates radiogenic strontium variability within the Basin of Mexico and greater central Mexico, further facilitating paleomobility studies within the

region. I first discuss the use of strontium isotopes in paleomobility within Mesoamerica and beyond and then consider geologic expectations for $^{87}\text{Sr}/^{86}\text{Sr}$ values within central Mexico and the Basin of Mexico. Finally, I present biogeochemical data on modern plant and water samples ($n=63$), analyzing them alongside published data ($n=16$; Price et al., 2000; Schaaf et al., 2012) to characterize biogeochemically distinguishable zones within the Basin of Mexico and central Mexico.

Strontium isotopes in studies of paleomobility

Radiogenic strontium isotopes are one of several isotopic systems that have been used to characterize paleomobility (Ehleringer et al., 2008; Ericson, 1985; Gulson et al., 1997; Luz et al., 1984; Schwarcz et al., 1991). $^{87}\text{Sr}/^{86}\text{Sr}$ values reflect regional geologic variability (Faure & Powell, 1972). Biologically available strontium present in soil and groundwater is incorporated into local plants and subsequently into hydroxyapatite, the hard tissues (including bone and enamel) of animals ingesting that vegetation (Blum et al., 2000; Flockhart et al., 2015; Sealy et al., 1991; Turekian & Kulp, 1956). By comparing the strontium isotopic values in human and animal hard tissues mineralizing at different times over the life course, bioarchaeologists can reconstruct prehistoric patterns of mobility between distinct geologic zones over the life course (Ericson, 1985; Knudson et al., 2016; Price et al., 1994, 2000; Sealy et al., 1995).

Strontium isotope systematics

Strontium is an alkaline earth metal typically found in rock, water, soil, plants, and animals at the parts-per-million (ppm) level (Faure, 1998; Faure & Powell, 1972). Of the four naturally occurring strontium isotopes, ^{87}Sr is radiogenic and is produced by the

slow radioactive decay of rubidium (^{87}Rb). Thus, the abundance of ^{87}Sr in a given region varies by the age and composition of local bedrock minerals (Faure, 1998; Faure & Powell, 1972). Geologically older igneous and granitic formations rich in parent ^{87}Rb are enriched in ^{87}Sr ($^{87}\text{Sr}/^{86}\text{Sr} > 0.750$) compared to geologically younger volcanic basalts, rhyolites, or andesites ($^{87}\text{Sr}/^{86}\text{Sr} \approx 0.702\text{--}0.704$), while marine carbonates and metamorphic formations often have intermediate values (R. A. Bentley, 2006; Faure & Powell, 1972; Sillen & Sealy, 1989). There is a large range of variation in comparison to the instrumental error of mass spectrometer measurements, which can generate accurate measurements up to the fourth decimal place or better (± 0.00001) (R. A. Bentley, 2006; Knudson et al., 2016). As such, geologic maps of bedrock types and ages can be used to predict expected $^{87}\text{Sr}/^{86}\text{Sr}$ variation.

Predictions based solely on geologic maps of bedrock types, however, are not always accurate. A number of factors, including the modification of source rock by erosion and preferential weathering of mineral with more radiogenic signatures, the addition of material from wind-derived material, and sea spray, can be mixed to produce different bioavailable strontium ratios that ultimately end up incorporated in hydroxyapatite (R. A. Bentley, 2006; Burton & Price, 2002; Reynolds et al., 2012). Thus, researchers have undertaken strontium isotope studies of local water sources, soils, plants, and animal bones to more accurately characterize bioavailable strontium variability in a given environment (R. A. Bentley et al., 2004; Evans & Tatham, 2004; Flockhart et al., 2015; Hodell et al., 2004; Knudson, Webb, et al., 2014; Pestle et al., 2013; Price et al., 2002; Wright, 2005).

Strontium isotopes and paleomobility across Mesoamerica

Studies using $^{87}\text{Sr}/^{86}\text{Sr}$ isotopes to reconstruct paleomobility throughout Mesoamerica have increased dramatically in recent years as archaeologists seek to directly test models of ancient migration, diaspora, and mobility within the region (Spence & White, 2010). Researchers have used radiogenic strontium isotopes to reconstruct ancient migration patterns (K. A. Miller, 2015; Miller Wolf & Freiwald, 2018; Price et al., 2000, 2014, 2015; Solís Pichardo et al., 2017; Wright, 2005, 2012; Wrobel et al., 2017), the geographic origins of sacrificial victims (Price et al., 2007; White et al., 2007), and past animal trade and management networks (Sharpe et al., 2018; Thornton, 2011), as well as long distance material culture trade networks (Thibodeau et al., 2018) and historic diasporas (Price et al., 2006, 2012).

Other studies have focused on characterizing $^{87}\text{Sr}/^{86}\text{Sr}$ variability across Mesoamerica. Hodell and colleagues (2004) carried out an extensive study of radiogenic strontium variability across the Maya region of southern Mexico, Belize, and Guatemala to identify isotopically distinct sub-regions. Similarly, Price and colleagues (2008) analyzed $^{87}\text{Sr}/^{86}\text{Sr}$ ranges more broadly across Mesoamerica. While they report a local range of $^{87}\text{Sr}/^{86}\text{Sr} = 0.7046\text{-}0.7051$ for the Basin of Mexico, little published data exist examining variability within central Mexico and the Basin itself.

Central Mexican geography, geology, and geochemistry

Understanding regional geology is essential to the study of variability in radiogenic strontium isotope values within central Mexico, which is defined here as including the modern Mexican states of Mexico State, Hidalgo, Puebla, Tlaxcala, and

Morelos, as well as Mexico City. Geologists have divided Mexico into several geologically and physiographically distinct morphotectonic provinces (Fig 1). However, only three morphotectonic provinces—the Sierra Madre Oriental, Mexican Volcanic Belt, and Sierra Madre Sur—make up central Mexico.

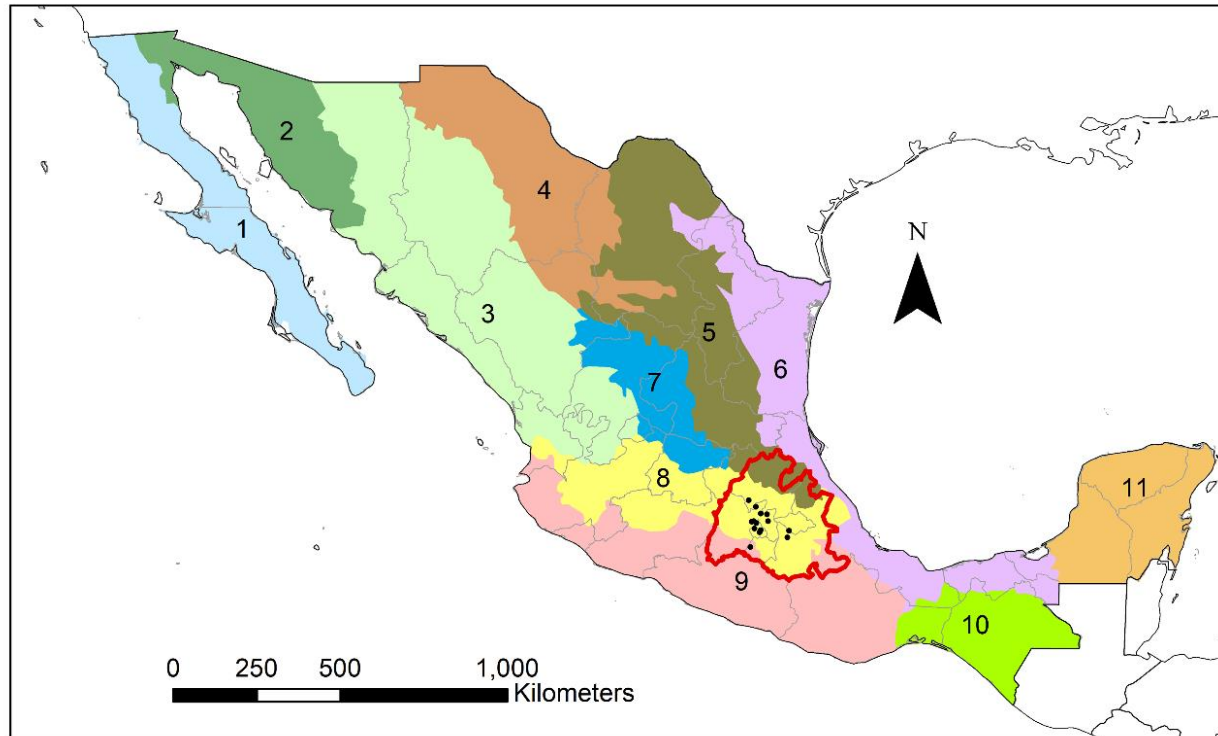


Figure 1. Morphotectonic provinces of Mexico. Central Mexico is outlined in red and is made up by parts of the Sierra Madre Oriental (5), the Mexican Volcanic Belt (8), and the Sierra Madre del Sur (9) morphotectonic provinces. Other morphotectonic provinces include Baja California Peninsula (1), the Northwestern Plains and Sierras (2), the Sierra Madre Occidental (3), the Chihuahua-Coahuila Plateaus and Ranges (4), the Gulf Coast Plain (6), the Central Plateau (7), the Sierra Madre de Chiapas (10), and the Yucatán Platform (11). Sites included in the study are indicated by black dots. Map created with free vector and raster map data (*Natural Earth*, n.d.). Morphotectonic data adapted from the Mexican Geological Service (Servicio Geológico Mexicano, 2007).

The geology of central Mexico is a complex mixture of recent volcanic highlands and older marine sedimentary deposits, along with a variety of metamorphic rocks (de Cserna, 1989; Ferrusquía-Villafranca, 1993; Morán-Zenteno, 1994). The northern portion of central Mexico is comprised of the Sierra Madre Oriental mountain range. The Sierra Madre Oriental is primarily made up of orogenic Mesozoic Jurassic and Cretaceous sedimentary carbonates, sandstones, and shales of marine origin with some metamorphic Precambrian and Paleozoic gneiss and schist outcrops (Morán-Zenteno, 1994; Ohmoto et al., 1966). Immediately to the south and forming the heart of central Mexico is the Mexican Volcanic Belt, which extends from the Pacific to Gulf coasts. The Mexican Volcanic Belt is a Cenozoic volcanic plateau with central basaltic andesites forming during the late Miocene and early Pliocene and younger southern andesites, dacites, and rhyolites forming more recently during the Quaternary (de Cserna, 1989; Demant, 1981; Moorbath et al., 1978; Mooser, 1975; Morán-Zenteno, 1994).

Finally, the southern edge of central Mexico is defined by the Sierra Madre del Sur mountain range. The Sierra Madre del Sur is the most geologically complex morphotectonic province in Mexico, composed of a northern segment of Mesozoic Jurassic and Cretaceous sediments and volcanic rock outcrops partially covered by Cenozoic volcanic and sedimentary rocks, a southern segment of Paleozoic and Mesozoic metamorphic rock outcrops and intrusive Mesozoic and Cenozoic batholiths, and a coastal Pacific area of andesitic Mesozoic Jurassic and Cretaceous volcanic-sedimentary rocks (Ferrusquía-Villafranca, 1993; Morán-Zenteno, 1994).

The Basin of Mexico in geological context

The Basin of Mexico, the primary region of interest in this study, is situated in the central-eastern part of the Mexican Volcanic Belt. It is a late Tertiary and Quaternary graben basin characterized by basaltic and andesitic volcanism with single rhyolite cones, featuring some of the most complex volcanic geology of Mexico (Mooser, 1975; Morán-Zenteno, 1994; Torres-Alvarado et al., 2000; Vázquez-Sánchez & Jaimes-Palomera, 1989). The Basin is enclosed by several mountain ranges, including the Sierra de Tepotzotlán and the Sierra de Pachuca to the north, the Sierra de Río Frío and the Sierra Nevada to the east, the Sierra de Chichinautzin to the south, and the Sierra del Ajusco and Sierra de las Cruces to the west.

While the underlying bedrock geology is likely the dominant contribution to the radiogenic strontium isotope composition of the piedmont and mountains of the Basin of Mexico, the alluvial plain represents a large catchment area for weathered minerals deposited by rivers and streams flowing into the Basin lakes. At high elevations, which tend to have high weathering rates, bioavailable $^{87}\text{Sr}/^{86}\text{Sr}$ and bedrock $^{87}\text{Sr}/^{86}\text{Sr}$ values are more often closely correlated (Aubert et al., 2002; Hoogewerff et al., 2001). At lower elevations, however, correlations between underlying bedrock and river content are less clear, as rivers carry suspended loads of upstream rocks and solids as well as precipitation, all of which could contribute geologically distinct strontium values to alluvial deposits (Hodell et al., 2004; Price et al., 2002; Tricca et al., 1999). This suggests that soils in the Basin of Mexico's alluvial plain may vary considerably in strontium isotope values and will likely average source materials. Thus, though the geology of the

Basin of Mexico provides starting expectations for ranges of radiogenic strontium variability, it is necessary to generate expected “local” ranges of bioavailable strontium values within the region to gain a more comprehensive understanding of variability within and beyond the Basin of Mexico.

Materials and Methods

Modern plant and water samples provide an excellent means of characterizing the bioavailable strontium within ecosystems. While soil $^{87}\text{Sr}/^{86}\text{Sr}$ values in a given geologic zone may vary greatly due to the distinct strontium concentrations and weathering profiles of minerals in the underlying bedrock (R. A. Bentley, 2006; Dasch, 1969), only a proportion of soil strontium is available to plants. As such, plant $^{87}\text{Sr}/^{86}\text{Sr}$ values provide a consistent average of local bioavailable strontium within a given ecosystem (Budd et al., 2000). Similarly, the majority of strontium in water sources is carried as dissolved or suspended sediment and primarily represents bioavailable strontium from rocks undergoing erosion within an ecosystem (Aubert et al., 2002; R. A. Bentley, 2006; Négrel et al., 2003; Palmer & Edmond, 1992; Price et al., 2002; Tricca et al., 1999).

Plant and water samples were collected between December 2015 and June 2017 from a total of 13 archaeological and agricultural sites from distinct ecological zones throughout the Basin of Mexico and greater central Mexico ($n = 63$). Universal Transverse Mercator (UTM) coordinate and elevation data for each sample were collected using a hand-held GPS unit. Plant samples were only collected if it was clear that they had not been treated with fertilizers or irrigation water, as these could skew signatures of local bioavailable strontium with non-local sources of strontium.

Furthermore, plants of varied rooting depths were sampled opportunistically. Plants with shallow rooting depths in topsoil (<1 m deep), such as grasses and many herbaceous plants, tend to exhibit $^{87}\text{Sr}/^{86}\text{Sr}$ values closer to atmospheric dust. In contrast, plants with deeper rooting depths, including many species of tree, exhibit $^{87}\text{Sr}/^{86}\text{Sr}$ values derived from local bedrock in addition to atmospheric sources (Reynolds et al., 2012). Including both of these sources allows for the more accurate characterization of bioavailable strontium in local ecosystems (Grimstead et al., 2017). Similarly, water samples were only collected from uncontaminated springs that would likely have been used by ancient inhabitants of the region (Buzon et al., 2011; Knudson, 2009). The Mexican Instituto Nacional de Antropología e Historia (INAH) does not require specific permissions to collect water or modern plant samples from the study sites. Furthermore, no endangered or protected plant species were involved in the study. Samples were imported to the Arizona State University Archaeological Chemistry Laboratory under permits granted to Pacheco-Forés from the United States Department of Agriculture Animal and Plant Health Inspection Service (PCIP-17-00469).

Additionally, published central Mexican $^{87}\text{Sr}/^{86}\text{Sr}$ values generated by Price and colleagues (2000) and Schaaf and colleagues (2012) were included in the study dataset ($n = 16$). Non-human baseline samples such as soils, plants, or faunal materials (Price et al., 2002) were incorporated. Data from published whole rock samples were not included, as these $^{87}\text{Sr}/^{86}\text{Sr}$ values were likely not bioavailable within the ecosystem. Finally, published data were included only if their provenience could be confirmed via GPS to provide reasonably accurate UTM coordinate and elevation data.

Biogeochemical methods

All samples were prepared at the Arizona State University Archaeological Chemistry Laboratory. Water samples were filtered (2.5 μ m diameter) and acidified to 5% HCl to prevent precipitates from forming, adsorption to bottle walls, and discourage bacterial and algal growth. When possible, pre-Hispanic diets were simulated through the manual isolation and analysis of edible components (e.g., seeds, berries, leaves) of dried plants (Runia, 1987). Plant samples were rinsed with 18.2 M Ω Millipore water to remove adhering dirt and were ashed in a furnace for approximately 10 hours at 800° C. Approximately 25.0 mg of ashed sample was digested in 2 mL of concentrated nitric and hydrochloric acid (HNO₃ + 3HCl) at approximately 50° C for 24 hours. This aggressive leach does not break down the silica tetrahedra structure of most silicate minerals, leaving much of the soil in a solid form while prioritizing the release of bioavailable strontium within plants. Leach solution was evaporated, and sample precipitates were redissolved in concentrated nitric acid and diluted to a 2 M stock solution.

Dissolved samples were analyzed at the Metals, Environmental, and Terrestrial Analytical Laboratory at Arizona State University. An aliquot was taken for elemental concentration by a Thermo Fisher Scientific iCAP quadrupole inductively coupled plasma mass spectrometer (Q-ICP-MS). Strontium was then separated with a PrepFAST, an automated low-pressure ion exchange chromatography system (Romaniello et al., 2015). Strontium was isolated from the sample matrix using Elemental Scientific, Inc. supplied Sr-Ca ion exchange resin (Part CF-MC-SrCa-1000) and ultrapure 5 M nitric acid (HNO₃). Each strontium cut from the PrepFAST was dried down in a Teflon beaker and

digested with concentration nitric acid and 30% hydrogen peroxide to remove organics from the resin. Once digested, samples were again dried down and reconstituted with 0.32 M nitric acid. Using concentration information from the Q-ICP-MS, the samples were diluted with 0.32 M nitric acid to a calculated constant concentration of 50 ppb Sr. Radiogenic strontium isotope ratios were measured on a Thermo-Finnigan Neptune multi-collector inductively coupled plasma mass spectrometer (MC-ICP-MS). The MC-ICP-MS has nine Faraday cups capable of simultaneous ion beam measurement, and this instrument was configured with an Elemental Scientific, Inc., Apex Q high sensitivity sample introduction system with an Elemental Scientific, Inc. 50 or 100 $\mu\text{L}/\text{minute}$ PFA-ST microflow nebulizer. This instrument has seven 1011 amplifiers and three 1012 amplifiers which can be designated for any of the Faraday cups.

Data was collected by measuring 60 simultaneous ratios integrating 4.194 seconds each. Samples were corrected for on-peak blanks, and in-line correction of the contributions of ^{84}Kr on ^{84}Sr and ^{86}Kr on ^{86}Sr using $^{83}\text{Kr}/^{84}\text{Kr}$ ratio of 0.201750 and $^{83}\text{Kr}/^{86}\text{Kr}$ ratio of 0.664533, after instrumental mass bias correction using a normalizing $^{88}\text{Sr}/^{86}\text{Sr}$ ratio of 8.375209. Samples were analyzed in three different analytical sessions. Typical sensitivity was >10 V on ^{88}Sr with a 50 ppb Sr solution, with ^{83}Kr values <0.0001 V. ^{85}Rb voltages for samples were typically <0.004 V due to the low Rb/Sr initial ratios of the samples and effective chemical purification, but all data was interference-corrected using a $^{85}\text{Rb}/^{87}\text{Rb}$ ratio of 2.588960, normalized to $^{88}\text{Sr}/^{86}\text{Sr}$ as above. Ratio outliers two standard deviations outside the mean were removed using a Matlab 2D-mathematical correction routine written by Dr. Stephen Romaniello, now at

University of Tennessee. Typical internal $^{87}\text{Sr}/^{86}\text{Sr}$ two standard error (SE) precision was $\sim 1\text{e-}6$.

Sequences included bracketing concentration-matched SRM 987 standards. SRM 987 was run as a bracketing standard with a measured value of $^{87}\text{Sr}/^{86}\text{Sr} = 0.710252 \pm 0.000026$ (2σ , $n = 89$). Each analytical session included a sequence incorporating SRM 987 standard in a range of variable concentrations to verify the accuracy of $^{87}\text{Sr}/^{86}\text{Sr}$ values for samples; reported values are all above the threshold for accurate $^{87}\text{Sr}/^{86}\text{Sr}$ values within the range of error of the bracketing standards. In addition, SRM 987 doped with calcium up to a ratio of Ca/Sr of 500 was run to simulate the accuracy and precision of isotope ratios in poorly purified samples with low yields. SRM 987 run at 50% concentration doped to a Ca/Sr of 500 was run as a check standard with a measured value of $^{87}\text{Sr}/^{86}\text{Sr} = 0.710253 \pm 0.000025$ (2σ , $n = 15$). IAPSO seawater (Ocean Scientific International Ltd., Havant, UK) as a secondary check standard had a measured value of 0.709182 ± 0.000010 (2σ , $n = 11$), within error of the published value of 0.709182 ± 0.000004 (Ma et al., 2013). NIST 1400 purified in parallel with samples had a measured value of 0.713124 ± 0.000023 (2σ , $n = 12$), similar to the published value of 0.713150 ± 0.0000160 (Galler et al., 2007).

Analytical methods

K-means cluster analysis was used to sort observed and published $^{87}\text{Sr}/^{86}\text{Sr}$, UTM, and elevation data into groups in R using the cluster and ggplot2 packages (Maechler et al., 2017; R Core Team, 2016; Wickham, 2009). K-means cluster analysis is a divisive iterative non-hierarchical pure locational clustering method (Kintigh, 1990; Kintigh &

Ammerman, 1982) that has been applied to the analysis of analysis of bioavailable $^{87}\text{Sr}/^{86}\text{Sr}$ isotopes (Hodell et al., 2004). Clusters were defined based on Euclidean distances to minimize the sum of squares error (SSE), thus minimizing variability within clusters while maximizing variability between clusters. A randomization procedure assessing changes in the global SSE for different cluster levels was conducted. A cluster solution was selected by comparing the difference in SSE in the original data to the mean SSE of 1,000 randomized iterations of the data.

Results and Discussion

Appendix B reports observed and published $^{87}\text{Sr}/^{86}\text{Sr}$ values of water, plant, faunal, and soil samples included the study $^{87}\text{Sr}/^{86}\text{Sr}$ values varied from 0.70432 to 0.70641. Among plant samples, opportunistically sampled non-native and non-edible plants did not provide significantly different values from native edible plants simulating pre-Hispanic diets (Fig 2).

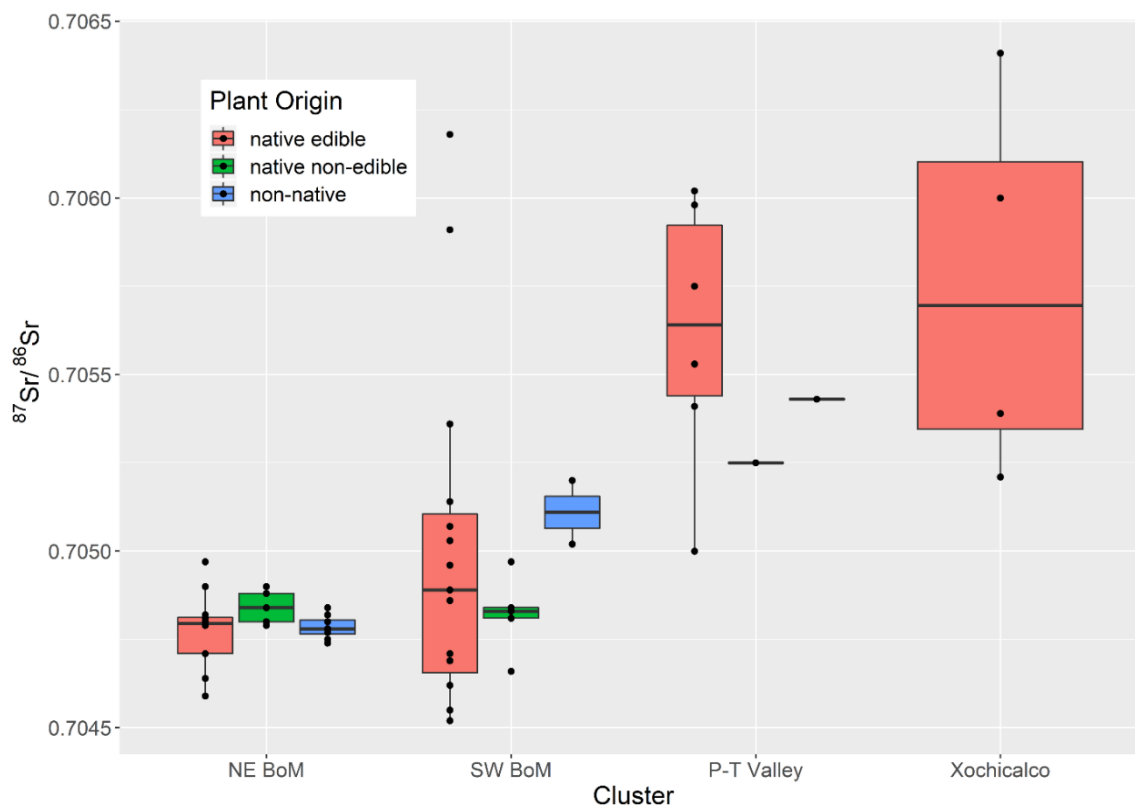


Figure 2. Cluster $^{87}\text{Sr}/^{86}\text{Sr}$ values in plant samples by plant origin. There were no significant differences between edible native plants and non-edible native plants or non-native plants. While non-edible native and non-native plants would not have contributed to past human and animal bioavailable $^{87}\text{Sr}/^{86}\text{Sr}$ values, they are included in this study to further characterize bioavailable strontium values in local ecosystems.

The randomization procedure indicates a five-cluster solution represents the greatest departure in the global SSE from randomness. The data are not normally distributed. Medians and interquartile ranges are therefore used to characterize $^{87}\text{Sr}/^{86}\text{Sr}$ variability within each cluster, following Price and colleagues (2008) (Table 1, Figs 3-4). In cases where sites or clusters have fewer than three samples, simple ranges are provided in lieu of interquartile ranges.

Table 1. $^{87}\text{Sr}/^{86}\text{Sr}$ medians and interquartile ranges for central Mexican subregions identified through k-means cluster analysis.

| Cluster | Geographic Subregion | Median $^{87}\text{Sr}/^{86}\text{Sr}$ | Interquartile $^{87}\text{Sr}/^{86}\text{Sr}$ Range | <i>n</i> |
|---------|------------------------------|--|---|----------|
| 1 | North of the Basin of Mexico | 0.70469 | 0.70466 - 0.70488 | 7 |
| 2 | Basin of Mexico Northeast | 0.70476 | 0.70464 - 0.70481 | 38 |
| 3 | Basin of Mexico Southwest | 0.70488 | 0.70470 - 0.70506 | 22 |
| 4 | Puebla-Tlaxcala Valley | 0.70548 | 0.70537 - 0.70581 | 8 |
| 5 | Xochicalco | 0.70570 | 0.70535 - 0.70610 | 4 |

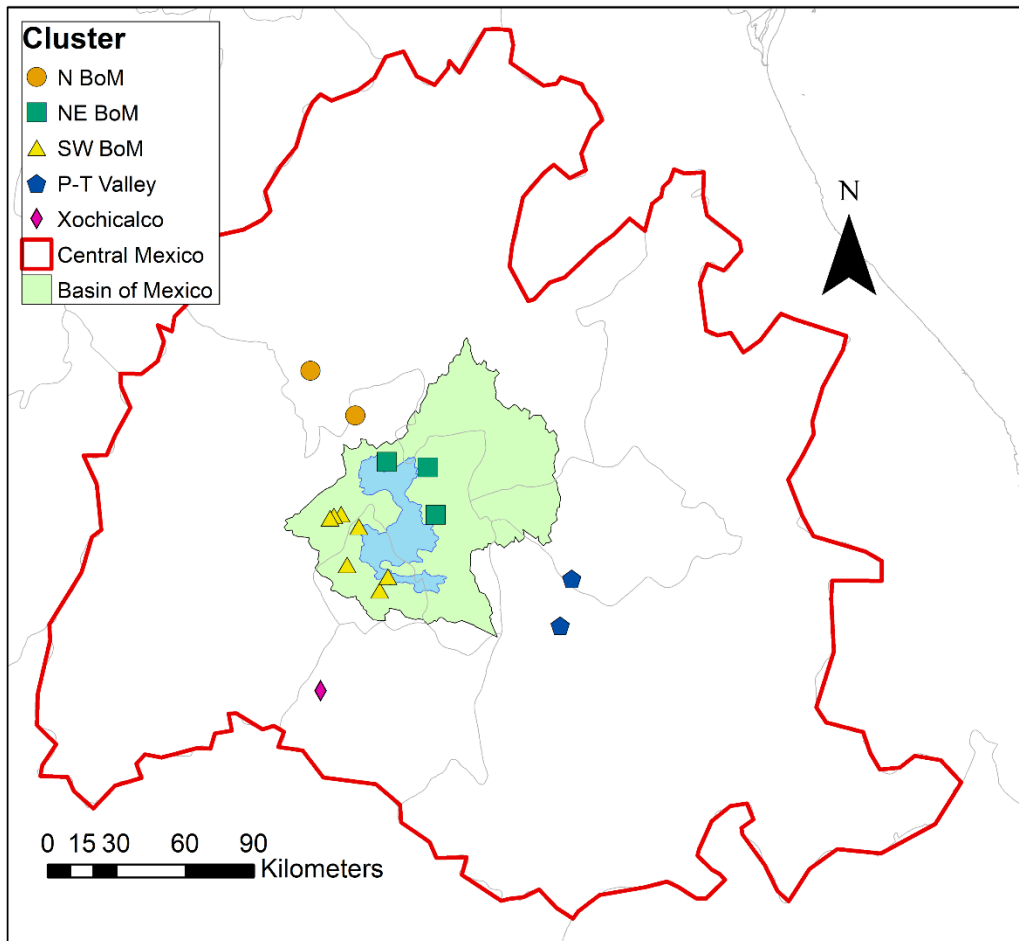


Figure 3. Sampled sites within central Mexico sorted by cluster membership. The Basin of Mexico is highlighted in green, and the extinct highland lake system is shown in blue. Map created with free vector and raster map data (*Natural Earth*, n.d.).

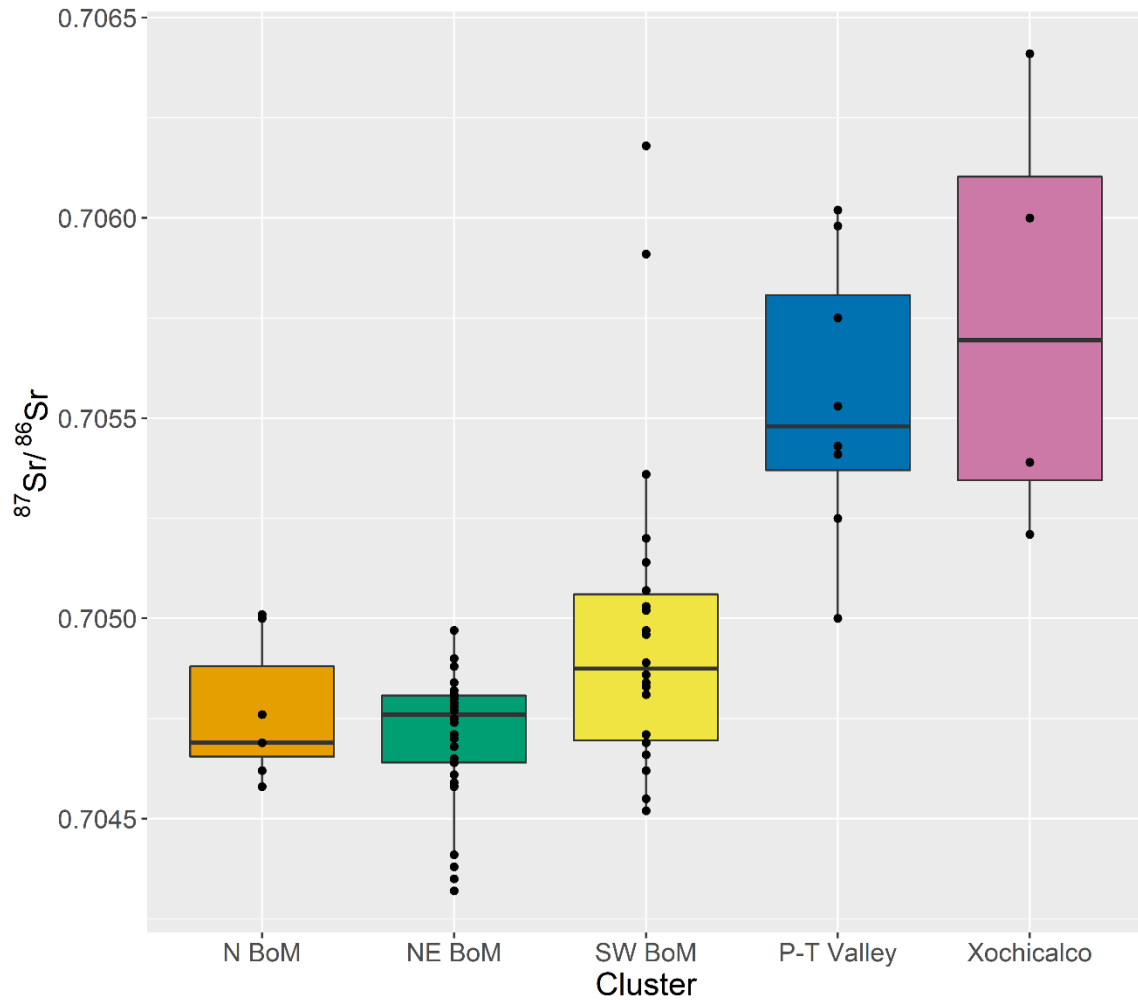


Figure 4. Medians and interquartile ranges of each clustered subregion. Data points are superimposed individual. BoM = Basin of Mexico, P-T = Puebla-Tlaxcala.

Each of the five clusters form culturally meaningful geographically distinct subregions within central Mexico (Fig 4). Cluster 1 is made up of two sites north of the Basin of Mexico. The Basin of Mexico itself is divided into two clusters, Cluster 2 which comprises the northeast of the Basin (three sites), and Cluster 3 which makes up the southwest of the Basin (seven sites). Cluster 4 is comprised of two sites in the Puebla-Tlaxcala Valley, and Cluster 5 is made up of the site of Xochicalco, south of the Basin of Mexico. Overall, cluster $^{87}\text{Sr}/^{86}\text{Sr}$ ranges conform to geologic expectations. The Basin of

Mexico clusters have the lowest $^{87}\text{Sr}/^{86}\text{Sr}$ values, reflecting the Basin's origins in Cenozoic volcanism (Servicio Geológico Mexicano, 2009; Vázquez-Sánchez & Jaimes-Palomera, 1989). In contrast, the Xochicalco cluster has the highest $^{87}\text{Sr}/^{86}\text{Sr}$ values, indicating the region's Mesozoic origins (Fries, Jr., 1958; Servicio Geológico Mexicano, 2009), although the intra-region variability is poorly constrained given the number of data points ($n = 4$). Finally, the Puebla-Tlaxcala Valley cluster has intermediate values consistent with the region's Mesozoic platforms overlain by Cenozoic volcanic rocks (Servicio Geológico Mexicano, 2008).

While the five-cluster model divides the Basin into two distinct groups, it is notable that there is significant overlap in $^{87}\text{Sr}/^{86}\text{Sr}$ values between Basin clusters, as well as with $^{87}\text{Sr}/^{86}\text{Sr}$ values in the cluster north of the Basin (Fig 4). Interestingly, $^{87}\text{Sr}/^{86}\text{Sr}$ values of the southwest Basin of Mexico cluster are most variable within the Basin of Mexico. This may reflect the greater diversity in age of the geologic substrate, as the southwestern Basin is made up by some of the oldest and youngest geologic formations in the Basin, including the Xochitepec Formation (Oligocene, 33.9-23.0 Ma) and the Chichinautzin mountain range (Quaternary, 2.6 Ma-present). Despite overlapping ranges among Basin of Mexico clusters, $^{87}\text{Sr}/^{86}\text{Sr}$ interquartile ranges indicate that sites in the Basin of Mexico are readily distinguishable from those in the Puebla-Tlaxcala Valley to the east, as well as Xochicalco to the south. Radiogenic strontium isotopes can thus be used to address questions of paleomobility at the regional level within central Mexico.

The generated Basin of Mexico interquartile range is consistent with previously published ranges. The two Basin of Mexico clusters (2-3) have a combined interquartile

range of $^{87}\text{Sr}/^{86}\text{Sr} = 0.70465\text{-}0.70487$ ($n = 60$). While this range is consistent with the $^{87}\text{Sr}/^{86}\text{Sr} = 0.7046\text{-}0.7051$ ($n = 86$) published by Price and colleagues (2008), examination of site-specific $^{87}\text{Sr}/^{86}\text{Sr}$ interquartile ranges indicates that this local range belies a great deal of variability within the Basin. Many sites in Basin of Mexico clusters can still be distinguished using radiogenic strontium analysis (Fig 5, Table 2). Furthermore, with a few notable exceptions, including Teotihuacan in the northeast Basin cluster and Cuicuilco and Tezozomoc in the southwest Basin cluster, all site-specific $^{87}\text{Sr}/^{86}\text{Sr}$ “local” ranges are narrower than the $^{87}\text{Sr}/^{86}\text{Sr}$ ranges of their assigned clusters. This suggests that while the k-means cluster analysis is useful on a larger scale for isotopically distinguishing the Basin of Mexico from surrounding regions within central Mexico, it does not perform well dividing the Basin itself into isotopically distinct subregions.

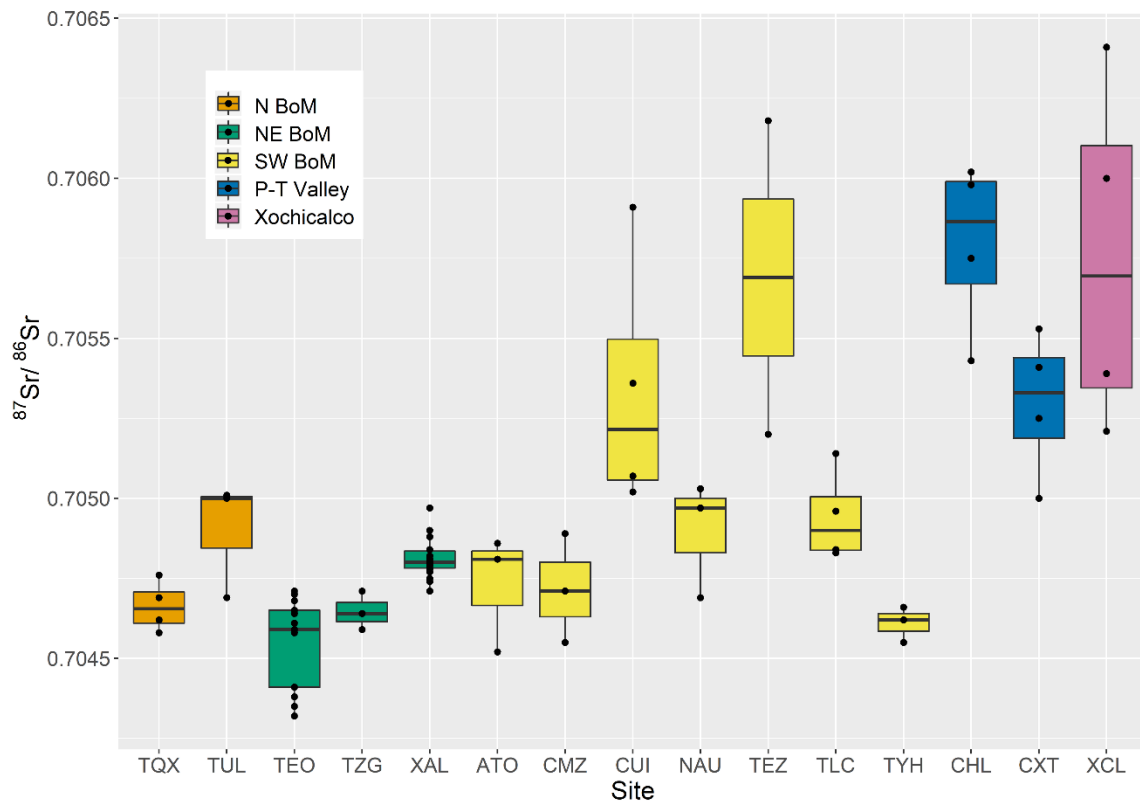


Figure 5. $^{87}\text{Sr}/^{86}\text{Sr}$ interquartile ranges of central Mexican sites, shaded by cluster. Individual data points are overlain. TQX = Tequixquiac, TUL = Tula, TEO = Teotihuacan, TZG = Texcotzingo, XAL = Xaltocan, ATO = San Pedro Atocpan, CMZ = Cerro Moctezuma, CUI = Cuicuilco, NAU = Naucalli, TEZ = Tezozomoc, TLC = Tlatelolco, THY = Santiago Tulyehualco, CHL = Cholula, CXT = Cacaxtla, XCL = Xochicalco.

Table 2. Central Mexican site-level $^{87}\text{Sr}/^{86}\text{Sr}$ medians and interquartile ranges.

| Cluster | Site | Median $^{87}\text{Sr}/^{86}\text{Sr}$ | Interquartile $^{87}\text{Sr}/^{86}\text{Sr}$ Range | <i>n</i> |
|---------|--------------------|--|---|----------|
| 1 | TQX | 0.70466 | 0.70461 - 0.70471 | 4 |
| 1 | TUL ^a | 0.70500 | 0.70485 - 0.70501 | 3 |
| 2 | TEO ^{a,b} | 0.70459 | 0.70441 - 0.70465 | 13 |
| 2 | TZG | 0.70464 | 0.70462 - 0.70468 | 3 |
| 2 | XAL | 0.70480 | 0.70478 - 0.70484 | 22 |
| 3 | ATO | 0.70481 | 0.70467 - 0.70484 | 3 |
| 3 | CMZ | 0.70471 | 0.70463 - 0.70480 | 3 |
| 3 | CUI | 0.70522 | 0.70506 - 0.70550 | 4 |
| 3 | NAU | 0.70497 | 0.70483 - 0.70500 | 3 |
| 3 | TEZ ^c | 0.70569 | 0.70520 - 0.70618 | 2 |
| 3 | TLC | 0.70490 | 0.70484 - 0.70501 | 4 |
| 3 | TYH | 0.70462 | 0.70459 - 0.70464 | 3 |
| 4 | CHL | 0.70587 | 0.70567 - 0.70599 | 4 |
| 4 | CXT | 0.70533 | 0.70519 - 0.70544 | 4 |
| 5 | XCL | 0.70570 | 0.70535 - 0.70610 | 4 |

^a From data on bulk soil samples published in Schaaf et al. (2012)

^b From data on faunal remains published in Price et al. (2000)

^c Only two samples were analyzed from Tezozomoc, so a simple range is presented rather than an interquartile range.

In the context of paleomobility studies, the use of cluster (Table 1) or site-specific (Table 2) $^{87}\text{Sr}/^{86}\text{Sr}$ interquartile ranges as a “local” bioavailable baseline should be determined by the scale of the research question. For example, if a study seeks to identify individuals who migrated into the Basin of Mexico from greater central Mexico and beyond, using cluster “local” $^{87}\text{Sr}/^{86}\text{Sr}$ ranges provides a robust mechanism for establishing individuals as non-locals within the Basin of Mexico. If, however, a study seeks to identify an individual’s residential mobility within the Basin of Mexico, using site-specific “local” $^{87}\text{Sr}/^{86}\text{Sr}$ ranges will provide a higher resolution analysis. With all such analyses, it is important to keep in mind that $^{87}\text{Sr}/^{86}\text{Sr}$ values are not unique and may mask the presence of non-locals if these individuals were from a region with similar

$^{87}\text{Sr}/^{86}\text{Sr}$ values. For this reason, the use of multiple lines of evidence and isotopic systems is essential (Knudson & Price, 2007; Price et al., 2007; White et al., 2007).

Conclusion

Analysis of presented and published bioavailable radiogenic strontium isotope ratios from central Mexico indicates that the Basin of Mexico can be distinguished isotopically from neighboring central Mexican regions. Furthermore, many sites within the Basin of Mexico itself can be distinguished from each other using radiogenic strontium isotopes, despite some overlap in $^{87}\text{Sr}/^{86}\text{Sr}$ cluster expected local ranges. This indicates that radiogenic strontium isotopes remain a powerful tool for examining paleomobility within central Mexico, particularly if used in concert with other isotopic systems, such as oxygen ($\delta^{18}\text{O}$) (Moreiras Reynaga, 2019).

Expanding knowledge of radiogenic strontium isotope variability within central Mexico is essential for future paleomobility work in the region, particularly given the hypothesized importance of migration in the cultural development of the region (Beekman & Christensen, 2003; Cowgill, 2015a). Future work will focus on augmenting the baseline data presented here with samples from additional sites throughout greater central Mexico. These data will be stored in an open-access comprehensive database of strontium isotopes throughout central Mexico with the ultimate goal of developing an $^{87}\text{Sr}/^{86}\text{Sr}$ isoscape for the region.

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CHAPTER 3

MIGRATION, VIOLENCE, AND THE "OTHER": A BIOGEOCHEMICAL APPROACH TO IDENTITY-BASED VIOLENCE AND HUMAN SACRIFICE IN THE EPICLASSIC

Target Journal – Journal of Anthropological Anthropology

A broad base of research in the social sciences demonstrates that identity-based violence—violence directed at individuals or groups perceived as categorically distinct or “other”—typically increases during periods of socio-political upheaval (Arkush & Tung, 2013; Bowman, 2001; Howard, 2014; Kurin, 2016; Messner et al., 2019; Riek et al., 2006; Simunovic et al., 2013; Tajfel & Turner, 2004). While archaeologists have documented innumerable cases of dramatic and punctuated social change in the past (e.g., Diehl & Berlo, 1989; Manahan, 2004; McAnany & Yoffee, 2012), only limited archaeological evidence of identity-based violence in the past exists (see Kurin, 2016). Here, I use a bioarchaeological approach to directly examine identity-based violence at the Epiclassic (600-900 CE) central Mexican shrine site of Non-Grid 4, where the remains of over 180 sacrificial victims were interred (Morehart et al., 2012). I use biogeochemical techniques to consider how diverse geographic origins and residential histories contributed to the selection of victims of ritual violence during a period of political, social, and demographic reorganization in the Basin of Mexico.

This manuscript first delineates a theoretical framework within which to investigate identity-based violence. I integrate social science research on the development of in-group preferences and out-group biases with realistic group conflict theory from

social psychology, as well as with empirical data from cross-cultural anthropological, evolutionary psychology, and behavioral case studies to establish a link between social group formation and identity-based violence in conditions of heightened scarcity and intergroup competition. I then contextualize this work within Mesoamerican migration and identity studies and previous research at the Non-Grid 4 shrine site. I introduce the biogeochemical methods used to identify residential mobility and migration among the Non-Grid 4 sacrificed individuals, presenting biogeochemical data on a sample of sacrificed individuals ($n = 73$). I conclude with a discussion of paleomobility patterns at Non-Grid 4 and their implications for understanding identity-based violence in Epiclassic central Mexico.

Social boundaries and identities

Social group formation is a fundamental part of the human experience. A large body of social and behavioral science research has shown that as humans, I innately sort ourselves into categories of “self” and “other”, creating ingroups and outgroups (Bigler et al., 1997; Tajfel et al., 1971; Tajfel & Turner, 2004). This categorizing behavior begins at an early age (Kelly et al., 2005; Moreira et al., 2017), and even the most minimal and arbitrary categorical differences reliably produce ingroup preferences and outgroup biases in controlled laboratory experiments among children (Bigler et al., 1997; Dunham & Emory, 2014; Dunham Yarrow et al., 2011; Nesdale et al., 2007) and adults (Ashburn-Nardo et al., 2001; Everett et al., 2015; Locksley et al., 1980; Mullen et al., 1992; Otten & Wentura, 1999; Riek et al., 2006; Tajfel et al., 1971).

Contemporary culturally salient social groups based on shared social identities also often produce similarly polarized responses of ingroup preferences and outgroup biases (Tajfel & Turner, 2004). Social identities are constructed via particularistic categorical attributes including but not limited to age, gender, ethnicity, or geographic origin (Brubaker & Cooper, 2000; Knudson & Stojanowski, 2008). These identities exist on both individual and communal scales and are not mutually exclusive but can intersect and interact in myriad ways (Collins, 2015; Walby et al., 2012). While some aspects of social identity can be invoked or deemphasized situationally, many others are highly visible, acting as mechanisms signaling group membership and exclusion (Jenkins, 2004; Tajfel & Turner, 2004). Indeed, particularly salient social identities such as ethnicity (Whitt & Wilson, 2007), race (Bobo, 1983), religiosity (Sosis & Ruffle, 2003), language or accent (Kinzler et al., 2011; Purnell et al., 1999), political affiliation (Fowler & Kam, 2007; Rand et al., 2009), and geographic origin or immigrant status (Esses et al., 2001; Zárate et al., 2004) have all been documented to produce negative stereotyping and prejudicial behavior towards individuals perceived to belong to an outgroup.

Theorizing the link between social identity and violence

Despite this consistent bias against outgroup members, ingroup/outgroup social boundary formation often takes place non-violently through the course of intergroup interactions. In Barth's (1969) seminal essay on ethnic groups and social boundaries, he argues that intergroup interactions are essential in the reification of social boundaries, as there can be no "self" without an "other" against which to identify.

However, social group formation can become violent during periods of scarcity or stress. Within social psychology, the realistic group conflict theory posits that opposing claims to scarce resources, such as power, prestige, or wealth will generate antagonism and hostility between groups (D. T. Campbell, 1965; Jackson, 1993; Sherif, 1966; Tajfel & Turner, 2004). Empirical data from cross-cultural anthropological studies, as well as evolutionary psychology and behavioral studies lend further support to this theory, suggesting that modern humans often carry out preemptive violence to defend their ingroup if they perceive outgroups as posing a potential threat (Böhm et al., 2016; Schmidt & Schröder, 2001; Simunovic et al., 2013).

Socio-political upheaval is thus often accompanied by increased identity-based violence. While states are generally characterized as having a stabilizing effect on violence (Andrushko & Torres, 2011; D'Altroy, 1992; Gómez et al., 2016; Sołtysiak, 2017), comprehensive analyses considering diachronic patterns in violence find that periods of political decentralization and the disintegration of pre-extant authority systems evince the highest rates of bioarchaeological (Andrushko & Torres, 2011; Arkush & Tung, 2013; Duncan, 2005; Serafin et al., 2014; Tiesler & Cucina, 2012) and archaeological evidence of violence (Arkush & Tung, 2013; Inomata, 2003, 2014; Mock, 1994). In such contexts, individuals or groups with distinct social identities may become targets of violence.

Thus, during times of political decentralization or reorganization, the presence of diverse social groups in competition over finite resources may escalate into violence. A social group defined by single or multiple common aspects of identity may perceive an

outgroup with a distinct identity as a potential threat or competitor (Riek et al., 2006; Zárate et al., 2004). This perception of threatening social difference can culminate in the enactment of violence against outgroups with divergent social identities (Bowman, 2001; Bush & Keyman, 1997; Sen, 2009). Historical examples of such identity-based violence during periods of political disruption abound, including in the ethnic violence following the collapse of the former Yugoslavia (Bowman, 1994), the tribalist violence in the chaos of the Albanian post-Soviet transition (Schwandner-Sievers, 2001), and the genocidal ethnic violence during the Rwandan Civil War (Mamdani, 2002). Moreover, the contemporary Fragile State Index includes identity-based violence as an indicator of state breakdown (Haken et al., 2014; Messner et al., 2019).

Despite this growing literature demonstrating that the perception of threatening social difference can motivate violence, there are few archaeological examples of identity-based violence. This may be because the subtleties of distinct social identities can be difficult to identify in the past. A bioarchaeological approach, however, is well suited to empirically reconstruct past individuals' multivalent social identities, as many facets of identity such as age, sex, or phenotypic expressions of biological relatedness are grounded in the physical body, while others leave traces on the body through behavior, such as dietary preferences or cultural modification (Buikstra & Scott, 2009; Knudson & Stojanowski, 2008, 2009).

A rare archaeological example from the pre-Hispanic central Andes provides convincing evidence of identity-based violence in the past. The ethnohistorically-documented Chanka ethnic group, marked by a visibly distinct form of cranial

modification, experienced increased rates of both non-lethal and lethal trauma during the period of political decentralization following the decline of the Wari Empire, indicating that these individuals were targeted for violence at least in part based on their highly visible ethnic identity (Kurin, 2014, 2016). Thus, historical and limited archaeological evidence indicates that during periods of socio-political disruption, groups with visibly distinct social identities may perceive each other as potentially threatening, culminating in intergroup violence.

This study further develops bioarchaeological approaches to identity-based violence in the past. Although there are countless dimensions of social identity and social difference that can be detected using bioarchaeological techniques, here I focus on one specific aspect of social identity—geographic origin and residential history as reconstructed through biogeochemical analysis. I use the central Mexican shrine site of Non-Grid 4 as a case study to explore how the perception of social difference contributed to the selection of victims of ritual violence.

Geographic origins and social difference in ancient Mesoamerica

The mere existence of differences between social groups is often not sufficient to produce tangible intergroup rivalries (Moya, 2013; Tajfel, 1959). An experimental study examining the evolutionary mechanisms behind social group formation found that while some arbitrary markers of group membership became accurate predictors for in-group favoritism and out-group discrimination, others failed to take on this social symbolism and did not lead to ingroup biases (Efferson et al., 2008). Similarly, in contemporary anthropological and behavioral economic studies, certain categorical aspects of identity

such as religious affiliation or political party membership are only salient within particular contexts, such as during prescribed periods of ritual celebration or phases of election campaigning (Moya & Boyd, 2015; Rand et al., 2009). Thus, in order to examine identity-based violence in any context—past or present—the cultural import of the specific categorical attributes of identity must be established (Esses et al., 2001; Jackson, 1993).

Ample ethnohistoric evidence indicates that geographic origin was a salient indicator of social difference in pre-Hispanic Mesoamerica. Contact period Nahua codices such as the Codex Boturini, the Codex Chimalpopoca, the Codex Azcatitlan, and others all emphasize migration stories and the importance of sacred places of origin (Akademische Druck, 1979; *Codex Azcatitlan*, 1949; *Codex Boturini: Tira de la peregrinación mexicana*, 1944; *Códice Aubin: Manuscrito azteca de la Biblioteca Real de Berlín*, 1963; Durán, 1867; Feliciano Velázquez, 1975; Kirchhoff et al., 1976). Similarly, the Florentine Codex, which describes daily life within the Aztec Empire, consistently associates non-Nahua peoples with the geographic locations of their city-states of origin. Indeed, many ethnic descriptors in the Florentine Codex come from the name of the lands these peoples inhabited. For example, in a passage describing the Toloque peoples, Sahagún writes: “By [the city of Toloacan] lies their mountain...It is said its name is Tolotzin [or] Tolotepetl. Some say – furthermore the Toloque also say – that because many reeds grow there the city is Toloacan, and the people Toloque” (Sahagún, 1961, p. 182). Moreover, in addition to cataloguing the importance of distinct ethnic groups’ places of origin, the Florentine Codex also describes specific customs tied to their place

of origin, including characteristic forms of dress, hairstyles, cuisines, languages, forms of worship, and skilled labor. These customs would likely have persisted and marked these peoples as different even after leaving their places of origin.

While ethnohistoric data reflect a Nahuatl-centric perspective from the Late Postclassic and Contact periods, archaeological evidence provides additional support and time depth for the persistent cultural salience of place of origin throughout ancient Mesoamerica. Some of the best evidence for the symbolic importance of geographic origins comes from the Classic period (200-600 CE) multi-ethnic urban center of Teotihuacan in central Mexico. Phenotypic, isotopic, and genetic studies have all identified the presence of multi-ethnic immigrant groups at Teotihuacan (Álvarez-Sandoval et al., 2015; Meza Peñaloza, 2015; Nado, 2017; Price et al., 2000; Solís Pichardo et al., 2017; Spence, 1974; White et al., 1998; White, Storey, et al., 2004). While many of these immigrants may have assimilated into the Teotihuacano host culture, others maintained a distinctive cultural identity over centuries of life at this massive central Mexican city (Begun, 2013; Manzanilla, 2017; Spence, 2005).

For example, several neighborhoods within Teotihuacan exhibit a persistent identification with and invocation of distant places of origin. The Tlailotlacan and N1W5:19 neighborhoods near the western edge of the city are respectively associated with the Oaxaca Valley approximately 360 kilometers south, and the Zacapu Basin in Michoacán, some 300 kilometers to the west. Archaeologists identified these ethnic enclaves through the presence of high volumes of non-local material culture, including foreign-style ceramics made out of local clays and produced using non-local technologies

(Abascal et al., 1974). Furthermore, non-local mortuary patterns (Begun, 2013; Gómez Chavez, 2002; Spence & Gamboa Cabezas, 1999) and the incorporation of foreign architectonic elements into Teotihuacan-style facades and apartment compounds (Cabrera Castro, 1998; Gómez Chavez, 1998; Spence, 2005) indicate a deliberate maintenance of a distinct social identity tied to residents' place of origin even after having relocated to central Mexico.

Beyond Teotihuacan, archaeologists have identified ethnic enclaves throughout Mesoamerica. In addition to his discussion of the Tlailotlacan Zapotec neighborhood at Teotihuacan, Spence (2005) uses material culture, mortuary practices, and architectural patterns to identify additional ethnic enclaves at several central Mexican sites, suggesting they may have made up a Zapotec diaspora network. Similarly, Winter (1998) identifies a possible Teotihuacano neighborhood at the Zapotec state capital of Monte Albán, offering further material evidence of the intentional invocation of distant places of origin throughout Mesoamerica.

Broad ethnohistoric and archaeological evidence thus indicates that geographic origin was indeed a salient indicator of social difference in pre-Hispanic Mesoamerica. I therefore investigate if geographic place of origin acted as a potential catalyst for identity-based violence during a period of socio-political upheaval in ancient central Mexico. I examine the skeletal human remains deposited as an apparent sacrificial offering at the site of Non-Grid 4, a shrine site in the northern Basin of Mexico dating to the Epiclassic period (Morehart et al., 2012). The site does not appear to be affiliated with a state power, setting it apart from other central Mexican examples of ritual violence

which are typically understood through the lens of state or elite power consolidation (e.g., Moreiras Reynaga, 2019; Price et al., 2007; White et al., 2002, 2007). The Non-Grid 4 shrine site is thus ideal for the investigation of identity-based violence as it represents an enigmatic example of large-scale violence during a socio-politically fraught time period. Here, I use biogeochemical methods to reconstruct the geographic origin and residential histories of the apparent sacrificial victims from the Epiclassic Non-Grid 4 shrine site to examine identity-based violence in ancient Mesoamerica.

Epiclassic central Mexico and the Non-Grid 4 shrine site

The Epiclassic period (600-900 CE) in the central Mexican highlands was an epoch of major social and political reorganization. Archaeologists typically characterize this time period on the basis of the decline of the Classic period regional center of Teotihuacan, which had previously exerted political control over most of central Mexico (Cowgill, 2015a; Diehl & Berlo, 1989; Rattray, 1996), as well as the dramatic demographic reorganization of the Basin of Mexico (Parsons et al., 2008; Sanders et al., 1979). Teotihuacan's rapid political decline seems to have created a political vacuum in the region. New competing political centers emerged in the areas surrounding the Basin (García Cook, 2013; Healan, 2012; Mendoza, 1992). Furthermore, the rise in dispersed hilltop settlements (Anderson et al., 2016; Gorenflo & Sanders, 2007; Hirth, 2000; Mastache et al., 2002; Morehart, 2016a; Parsons et al., 2008) and the proliferation of militaristic iconography (Finegold, 2012; Hirth, 1989; Koontz, 1994; McVicker, 2007; Nagao, 1989; Ringle et al., 1998) suggest increased conflict accompanying the political decentralization of the region during this period.

The Epiclassic period is also associated with large-scale migrations into the Basin of Mexico. The appearance of the Coyotlatelco ceramic complex throughout central Mexico (Tozzer, 1921) has been interpreted as evidence of immigration into the region (Beekman & Christensen, 2003; Cowgill, 2013; Crider, 2013; Healan & Cobean, 2019; Hernández & Healan, 2019; Rattray, 1966; Sánchez, 2013; Solar Valverde, 2006), although some maintain that the complex represents either a local or autonomous development (Fournier & Bolaños, 2007; Sanders, 1986, 2006; Sugiura Yamamoto, 2006). Bioarchaeological evidence, however, further supports Epiclassic migration hypotheses. Limited biogeochemical analyses of Epiclassic human remains have demonstrated the presence of migrant individuals (three of five individuals examined) buried in the Cueva de las Varillas at Teotihuacan (Price et al., 2000). Furthermore, biodistance analyses suggest a population replacement took place in the Basin of Mexico during the Epiclassic period (Beekman & Christensen, 2003; Christensen, 1997; García Velasco, 2019; González-José et al., 2007; Meza-Peñaloza et al., 2019; Ragsdale & Edgar, 2018). Thus, as a period characterized by dramatic social change including political reorganization, competition for access to resources, increased conflict, immigration, and demographic change, the Epiclassic provides an optimal setting within which to examine identity-based violence.

Ritual violence at Non-Grid 4

Identified through the course of regional survey, the site of Non-Grid 4 is located on an anthropogenic rise in the now-extinct Lake Xaltocan in the northern Basin of Mexico (Fig 6). The site has been interpreted as a shrine as it was the locus of repeated

ritual activities and lacks a residential component (Morehart, 2010, 2017b; Morehart et al., 2012). It is comprised of a slightly elevated platform associated with multiple temporally distinct deposits of human remains and ritual paraphernalia (Meza-Peñaloza et al., 2019; Morehart et al., 2012). Deposits of incense burners and figurines at the base of the constructed platform indicate that it was likely established to petition and repay deities associated with rain, water, and fertility.

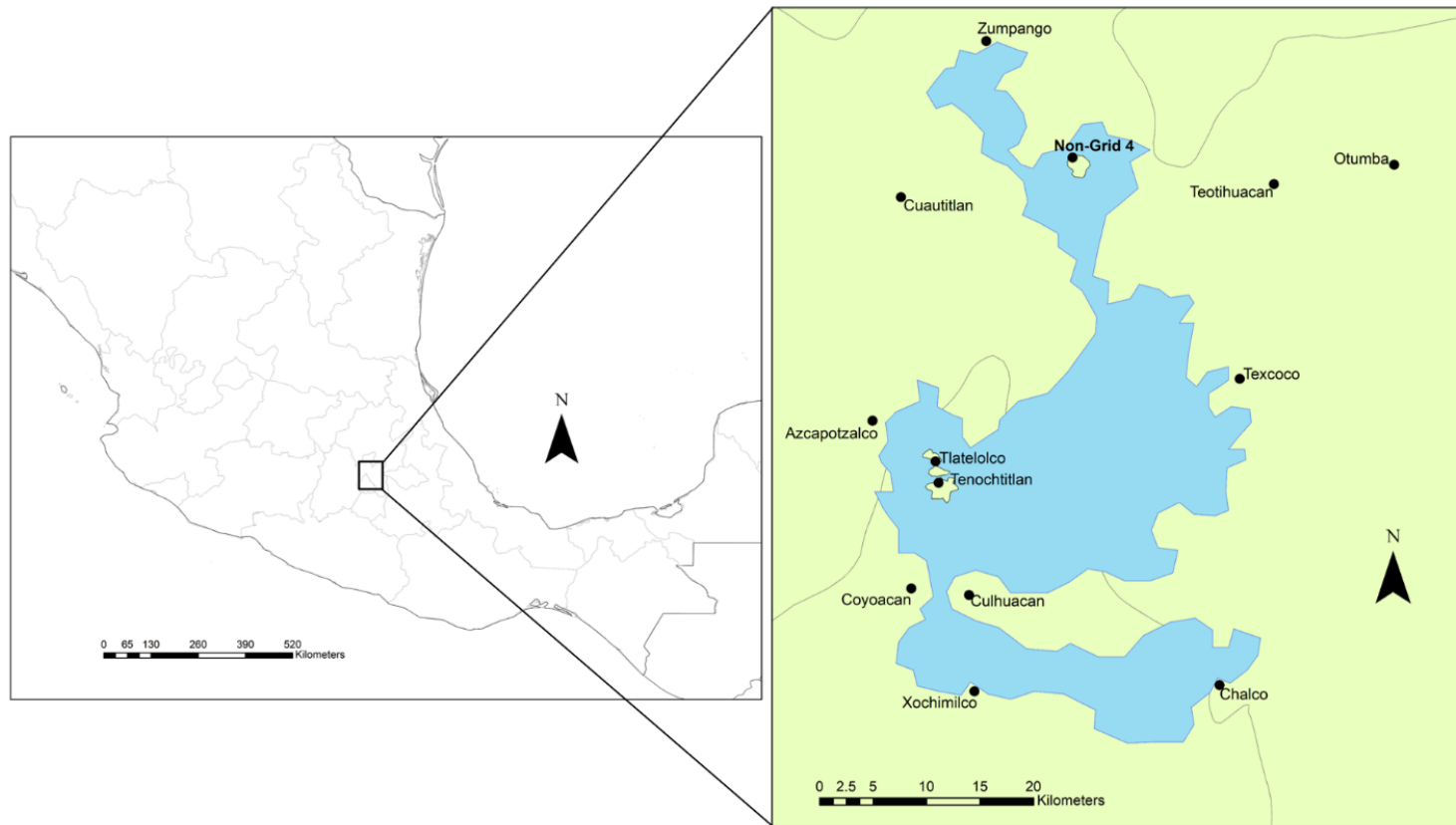


Figure 6. The Non-Grid 4 shrine site and pre-Hispanic highland lake system in central Mexico.

AMS radiocarbon dates of recovered botanical remains ranging from 550 – 890 CE place the site within the Epiclassic period (Morehart et al., 2012), and the creation of the shrine coincides with Teotihuacan's terminal phase after the burning of the city (Cowgill, 2015a; Manzanilla, 2003, p. 94). A range of artifacts recovered from the site, including Coyotlatelco Red on Natural, Garita Black-Brown ollitas, and Mazapan Red on Natural ceramics also suggest the site dates to the Epiclassic period (Morehart, 2010; Morehart et al., 2012), as these ceramics are widely accepted as Epiclassic and transitional Early Postclassic chronological markers (see Crider, 2011 for a regional discussion of the chronology of ceramics from this period).

The human remains interred at Non-Grid 4 postdate Teotihuacan's decline and have been interpreted as sacrificial deposits (Morehart et al., 2012). Although bioarchaeological evidence of violence in central Mexico during the Epiclassic period is rare, the human remains at Non-Grid 4 represent a major exception. Many of the deposited crania exhibit sharp force peri-mortem trauma consistent with throat slitting, exsanguination, and subsequent decapitation (Meza-Peñaloza et al., 2019; Pacheco-Forés, 2017). The presence of articulated mandibles and cervical vertebrae in association with the crania, along with the lack of evidence of sharp force trauma consistent with defleshing or trophy head preparation, and the lack of weathering all suggest that individuals were decapitated in situ and their heads were interred shortly after decapitation (Meza-Peñaloza et al., 2019; Morehart et al., 2012), most likely as part of a ritual offering (Morehart, 2017b). The remains' presence at the shrine site has thus been interpreted as evidence of competition and conflict over political influence in the wake of

Teotihuacan’s collapse and access to resources in a possible period of climatic change (Morehart et al., 2012).

The human remains within the sacrificial deposits were extensive. A minimum of 180 human crania associated with mandibles and cervical vertebrae were unearthed in east-facing rows and pits beneath the shrine platform (Morehart, 2017b; Morehart et al., 2012; Pacheco-Forés, 2016). Limited additional postcranial material was recovered from the site, with the exception of distal manual and pedal phalanges, which were found in the eye orbits of sacrificed individuals (Morehart et al., 2012). Most individuals were adult males, although some females and adolescents of both sexes were also present (Table 3), and many of the deposited crania exhibit cranial and dental modification (García Velasco, 2014; Pacheco-Forés, 2016; Sholts et al., 2014).

Table 3. Demography of Non-Grid 4 sacrificial victims.

| Sex | Adolescents | | Adults | | Unobservable | | Total | |
|---------------|-------------|-----|----------|-----|--------------|-----|----------|------|
| | <i>n</i> | % | <i>n</i> | % | <i>n</i> | % | <i>n</i> | % |
| Male | 12 | 7% | 114 | 63% | 7 | 4% | 133 | 74% |
| Female | 5 | 3% | 6 | 3% | 0 | 0% | 11 | 6% |
| Indeterminate | 5 | 3% | 17 | 9% | 14 | 8% | 29 | 20% |
| Total | 22 | 12% | 137 | 76% | 14 | 12% | 180 | 100% |

Human sacrifice and identity-based violence

Ethnohistoric and archaeological evidence indicate that human sacrifice was a culturally sanctioned form of violence practiced ubiquitously throughout ancient Mesoamerica (Galtung, 2013; Graulich, 2006; López Luján & Olivier, 2010; Sahagún, 1951). Most bioarchaeological analyses of victims of ritual violence are framed in the context of state or elite power consolidation. For example, studies investigating the age-

at-death and sex of sacrificed individuals argue that victims represented deity impersonators who were sacrificed at times of great adversity to curry favor with the gods and gain political prestige for ruling elites (De La Cruz et al., 2008; López Luján et al., 2010; Román Berrelleza, 2010). Similarly, researchers investigating the geographic origins of sacrificial victims at Teotihuacan argue the diverse residential histories of sacrificial victims interred at the Pyramid of the Feathered Serpent and the Pyramid of the Moon reflect the city's far-reaching influence and militaristic presence within Mesoamerica (Spence et al., 2004; White et al., 2002, 2007), while isotopic studies finding locals and non-locals among Aztec sacrificial victims at the Templo Mayor in the Postclassic imperial capital of Tenochtitlan have been used to argue for the effectiveness of the Mexica imperial system (Moreiras Reynaga, 2019). Likewise, in the Maya region, analyses of residential mobility have been used to characterize victims in terms of their relationships to powerful elites as either retainer burials or captive warriors put to death to commemorate a military victory or a ruler's passing (Hoffmeister & Wright, 2016; Lorenz et al., 2016; Price et al., 2007). The absence of an affiliated state or centralized power, or indeed even of a residential component of the site at the Non-Grid 4 shrine sets it apart from these previous studies. The site cannot be understood through a lens of state or elite power consolidation.

I thus use the Non-Grid 4 shrine for investigating identity-based violence, as it represents instances of large-scale violence and human sacrifice during a period of socio-political upheaval. Analyses of victims of ritual violence throughout Mesoamerica indicate that sacrificial victims do not represent a random cross-section of the population

at various archaeological sites. For example, biologically frail or disabled individuals were specifically selected for sacrifice in some contexts (Crandall & Thompson, 2014; De La Cruz et al., 2008; Kieffer, 2017), while particular kinship groups appear to have been disproportionately selected for sacrifice in others (Duncan, 2011, 2012). All members of society were therefore not equally likely to be selected as sacrificial victims. Sacrificial deposits thus provide an excellent context in which to examine identity-based violence, allowing bioarchaeologists to discern whether specific aspects of victims' social identities predisposed them to suffer violence. In the volatile socio-political landscape of the central Mexican Epiclassic, aspects of victims' social identities, including their geographic origins and residential histories, could have acted as powerful indicators of social difference that culminated in violence.

Reconstructing paleomobility through biogeochemistry

Biogeochemical methods provide a unique opportunity to directly examine paleomobility on an individual scale. Radiogenic strontium ($^{87}\text{Sr}/^{86}\text{Sr}$) and stable oxygen ($\delta^{18}\text{O}$) isotope analyses are often used to reconstruct the residential histories of past peoples (Ericson, 1985; Price et al., 2000; Schwarcz et al., 1991; White et al., 1998). Strontium isotopes are incorporated into calcified tissues from the environment through consumed food and imbibed water (Flockhart et al., 2015; Sealy et al., 1991; Turekian & Kulp, 1956). They reflect the age and composition of geologic bedrock in the region where an individual lived during tissue development (R. A. Bentley, 2006; Dasch, 1969; Faure & Powell, 1972). In contrast, oxygen isotopes are incorporated into human hard tissues primarily through imbibed liquids (Longinelli, 1984; Luz et al., 1984). Oxygen

isotopes reflect regional hydrology and are correlated with environmental factors such as elevation, temperature, humidity, and latitude in the region where an individual lived during tissue development (Bowen et al., 2005; Craig, 1957; Dansgaard, 1964; Gat, 1996; M. J. Kohn, 1996; Luz & Kolodny, 1989). Additionally, oxygen isotope data in enamel samples may also be influenced by the consumption of ^{18}O -enriched breast milk in infancy (Herring et al., 1998; Wright & Schwarcz, 1998, 1999), as well as through culturally mediated brewing, cooking, and water storage practices (Brettell et al., 2012; Gagnon et al., 2015).

Strontium isotope baseline variability in central Mexico

The underlying geology of central Mexico indicates that radiogenic strontium isotopes can be used to reconstruct paleomobility throughout the region (Morán-Zenteno, 1994; Price et al., 2008). Younger geologic formations tend to have lower $^{87}\text{Sr}/^{86}\text{Sr}$ values, while older formations have higher $^{87}\text{Sr}/^{86}\text{Sr}$ values (R. A. Bentley, 2006; Faure & Powell, 1972). The Non-Grid 4 shrine is located in the Basin of Mexico, which lies within the Mexican Volcanic Belt morphotectonic province. This geologic region is characterized by Late Cenozoic andesitic, dacite, and rhyolite pyroclastic flows (Moorbath et al., 1978; Morán-Zenteno, 1994; Salinas Prieto et al., 2007). Late Cenozoic rocks from the Mexican Volcanic Belt exhibit an average of $^{87}\text{Sr}/^{86}\text{Sr}=0.7040 \pm 0.00099$ (2σ , $n=329$; Torres-Alvarado et al., 2000). In contrast, mean radiogenic strontium values from rocks to the north in the Central Mesa province are lower, exhibiting an average of $^{87}\text{Sr}/^{86}\text{Sr}=0.7032 \pm 0.00058$ (2σ , $n=28$; Torres-Alvarado et al., 2000). Rocks from other neighboring morphotectonic provinces all exhibit radiogenic strontium values higher than

those reported in the Mexican Volcanic Belt. For example, rocks from Paleozoic marine sedimentary deposits in the Sierra Madre Oriental mountain range to the northeast have an average of $^{87}\text{Sr}/^{86}\text{Sr}=0.7070 \pm 0.005$ (2σ , $n=5$; Ohmoto et al., 1966), Cenozoic rocks from rhyolitic and andesitic volcanic deposits in the Sierra Madre Occidental mountains to the northwest exhibit mean values of $^{87}\text{Sr}/^{86}\text{Sr}=0.7058 \pm 0.0034$ (2σ , $n=105$; Torres-Alvarado et al., 2000), and rocks from the complex mixed Cenozoic, Mesozoic, and Paleozoic geology of the Sierra Madre Sur mountains to the south exhibit an average of $^{87}\text{Sr}/^{86}\text{Sr}=0.7041 \pm 0.00256$ (2σ , $n=22$; Torres-Alvarado et al., 2000). It is important to note that these whole rock strontium values do not necessarily reflect bioavailable $^{87}\text{Sr}/^{86}\text{Sr}$ values that would be taken up into human and animal hard tissues. Instead, variability in bioavailable $^{87}\text{Sr}/^{86}\text{Sr}$ values is typically obtained from plant, faunal, alluvial soil, and water samples and often—but not always—mirrors $^{87}\text{Sr}/^{86}\text{Sr}$ bedrock geology variability (R. A. Bentley, 2006; Evans & Tatham, 2004; Price et al., 2002).

Studies of bioavailable radiogenic strontium sources, including plants, water, and alluvial soils within the Basin of Mexico itself confirm that the Basin can be distinguished from neighboring parts of central Mexico using radiogenic strontium isotopes (Pacheco-Forés et al., 2020; Price et al., 2008). For example, the Basin of Mexico has a reported local range of $^{87}\text{Sr}/^{86}\text{Sr}=0.7046\text{-}0.7048$ ($n=60$; Pacheco-Forés et al., 2020). This range stands in contrast to other central Mexican subregions with higher reported local ranges of $^{87}\text{Sr}/^{86}\text{Sr}=0.7053\text{-}0.7058$ ($n=8$; Pacheco-Forés et al., 2020) in the Puebla-Tlaxcala Valley to the east, and $^{87}\text{Sr}/^{86}\text{Sr}=0.7053\text{-}0.7061$ ($n=4$; Pacheco-Forés et al., 2020) in the Xochicalco Formation in Morelos immediately to the south of the Basin.

Furthermore, significant variability in strontium isotopes exists within the Basin of Mexico itself, allowing many sites within the Basin to be distinguished isotopically using strontium isotopes (Pacheco-Forés et al., 2020).

The Basin of Mexico is also isotopically distinguishable from regions beyond central Mexico. Plant samples from western Mexico exhibit a lower reported local range of $^{87}\text{Sr}/^{86}\text{Sr}=0.7039\text{-}0.7040$ ($n=15$; Price et al., 2008), while plant samples from further to the south in the Oaxaca Valley have a higher reported local range of $^{87}\text{Sr}/^{86}\text{Sr}=0.7059\text{-}0.7077$ ($n=95$; Pacheco-Forés et al., in preparation) faunal samples from northern Mexico exhibit a higher reported local range of $^{87}\text{Sr}/^{86}\text{Sr}=0.7066\text{-}0.7072$ ($n=28$; Offenbecker, 2018). Similarly, the Basin is isotopically distinct from several subregions in the Maya region, including the higher expected local $^{87}\text{Sr}/^{86}\text{Sr}$ values in the northern lowlands ($^{87}\text{Sr}/^{86}\text{Sr}=0.7077\text{-}0.7092$, $n=40$), the southern lowlands ($^{87}\text{Sr}/^{86}\text{Sr}=0.7069\text{-}0.7077$, $n=86$), and the Maya Mountains ($^{87}\text{Sr}/^{86}\text{Sr}=0.7119\text{-}0.7151$, $n=3$); lower expected local $^{87}\text{Sr}/^{86}\text{Sr}$ values in the volcanic highlands ($^{87}\text{Sr}/^{86}\text{Sr}=0.7038\text{-}0.7041$, $n=34$); and only partially overlaps with the lower portion of the expected local $^{87}\text{Sr}/^{86}\text{Sr}$ range of the Montagua Valley ($^{87}\text{Sr}/^{86}\text{Sr}=0.7041\text{-}0.7072$, $n=26$; Hodell et al., 2004).

The expected local range for the Non-Grid 4 shrine site is $^{87}\text{Sr}/^{86}\text{Sr}=0.7047\text{-}0.7048$ ($n=22$; Pacheco-Forés et al., 2020). However, because the Non-Grid 4 shrine site did not contain a residential component, using the Non-Grid 4 site-specific “local” range may artificially inflate the number of non-local individuals identified, as there is no archaeological evidence that individuals interred at the site lived in and consumed strontium and oxygen sources from the immediate shrine environment. Therefore, for the

purposes of this study, I use the combined Basin of Mexico local range of $^{87}\text{Sr}/^{86}\text{Sr}=0.7046\text{-}0.7048$ ($n=60$; Pacheco-Forés et al., 2020) to represent “local” Non-Grid 4 values. This broader “local” range provides a more conservative and robust means of identifying individuals who were non-local not only to the immediate vicinity of the shrine site, but to the whole Basin. Unless otherwise indicated, all expected local $^{87}\text{Sr}/^{86}\text{Sr}$ ranges represent interquartile ranges of sampled $^{87}\text{Sr}/^{86}\text{Sr}$ values, as these are often non-normally distributed (Price et al., 2008).

Oxygen isotope baseline variability in central Mexico

In addition to geologic variability, there is also great diversity in rainfall, temperature, humidity, and elevation across central Mexico and Greater Mesoamerica, indicating that stable oxygen isotopes are suitable for examining paleomobility in the region (Wallén, 1955; Wassenaar et al., 2009; White et al., 1998). General models of oxygen isotope behavior indicate that low elevation, hot, humid, or rainy regions typically exhibit higher $\delta^{18}\text{O}$ values, while those at higher elevations, or are cooler and dried have lower $\delta^{18}\text{O}$ values (Dansgaard, 1964; Gat, 1996). Thus, the Basin of Mexico, which is at a high elevation and receives relatively little rain annually (~ 750 mm; Wallén, 1955) has a reported meteoric water range of $\delta^{18}\text{O}_{\text{mw(VSMOW)}} = -10.8\text{‰}$ to -8.8‰ (Issar et al., 1984; Ortega-Guerrero et al., 1997). This local range includes data on shallow groundwater as well as surface precipitation, as several studies have found negligible differences in $\delta^{18}\text{O}_{\text{VSMOW}}$ values between these two water sources throughout Mexico (IAEA, 1992; Issar et al., 1984; Wassenaar et al., 2009).

In contrast to the available $\delta^{18}\text{O}_{\text{mw (VSMOW)}}$ local range for the Basin of Mexico, oxygen isotope isoscapes of $\delta^{18}\text{O}_{\text{mw(VSMOW)}}$ values throughout Mexico show that the eastern Gulf Coast region of Mexico exhibits a relatively high range of $\delta^{18}\text{O}_{\text{mw(VSMOW)}} = -6.0$ to -3.0‰ , the southern Mexican highlands exhibit an intermediate range of $\delta^{18}\text{O}_{\text{mw(VSMOW)}} = -9.0$ to -6.0‰ , and the western highlands exhibit a wide range of $\delta^{18}\text{O}_{\text{mw(VSMOW)}} = -12.0$ to -4.0‰ (Moreiras Reynaga, 2019; Wassenaar et al., 2009). These local ranges conform with expected altitudinal and environmental effects.

According to oxygen isotope isoscapes, the Non-Grid 4 shrine site's local range is $\delta^{18}\text{O}_{\text{mw(VSMOW)}} = -10.0$ to -9.0‰ (Moreiras Reynaga, 2019; Wassenaar et al., 2009). Recent studies, however, have shown that differences in preparation methods and analytical techniques can lead to substantial inter-laboratory differences in reported $\delta^{18}\text{O}$ values (Pestle et al., 2014). I therefore use the Minimum Meaningful Difference (MMD) value of $\pm 3.1\text{‰}$ to establish a $\delta^{18}\text{O}_{\text{mw (VSMOW)}}$ local range for the Non-Grid 4 shrine site that considers inter-laboratory variability as well as observed $\delta^{18}\text{O}_{\text{mw}}$ variability. Thus, for the purposes of this study, the Non-Grid 4 shrine site's local range is $\delta^{18}\text{O}_{\text{mw (VSMOW)}} = -12.6\text{‰}$ to -6.4‰ .

Strontium and oxygen isotope sources in the Basin of Mexico

Dietary sources of strontium and oxygen isotopes within the Basin of Mexico were largely from local rather than non-local imported sources (Nado, 2017). The greatest source of high-calcium high-strontium foods in the pre-Hispanic Basin of Mexico was likely terrestrial plant sources, such as beans (*Phaseolus vulgaris*) and

amaranth (*Amaranthus spp.*; Santley & Rose, 1979). There is extensive evidence of intensive agricultural field systems throughout the Basin over the course of the pre-Hispanic period (Sanders et al., 1979), including in the Xaltocan lake zone (Morehart, 2012), indicating that peoples living in the Basin were likely consuming primarily locally-produced plants as part of their diet.

While maize (*Zea mays*), the largest component of pre-Hispanic Basin diets (Santley & Rose, 1979), contains negligible amounts of strontium (Harmon et al., 1969), the nixtamalization process represents a substantial source of dietary strontium. This cooking technique involves treating maize with alkaline solution often made from strontium-rich slaked lime to increase nutritional value of maize (Barba Pingarrón, 2013; Johnson & Marston, 2020). While several studies suggest the Tula region just north of the Basin may have served as a source for lime in archaeological plaster (Barba et al., 2009; Miriello et al., 2011; Parsons et al., 2008), there are also rich lime deposits in the Zumpango region in the northwestern Basin of Mexico (Parsons & Gorenflo, 2008; Sanders et al., 1979, p. 145) that could have been exploited for nixtamalization purposes.

Marine resources are also a significant source of strontium and reflect seawater values of $^{87}\text{Sr}/^{86}\text{Sr}=0.7092$ (Veizer, 1989). Diet in the Basin of Mexico, however, largely excluded marine resources (Nado et al., 2016; Parsons, 2008; Widmer & Storey, 2017). While there is evidence of importation of marine fish to one elite residential compound at Teotihuacan (Manzanilla, 2007; Rodríguez Galicia & Valadez Azúa, 2013), wider paleodietary reconstructions indicate that consumption of marine resources was atypical (Nado, 2017). Similarly, the consumption of marine salt, which was widely traded

throughout the Maya region, could also be a non-local source of strontium that would reflect seawater rather than local values (Andrews, 1983; Fenner & Wright, 2014; Wright, 2005). There is extensive evidence, however, of local salt production within the Basin of Mexico along the shores of the three saline highland lakes (Parsons, 2001; Sanders et al., 1979, pp. 57–58). Therefore, individuals living in the Basin would most likely have consumed these local sources of salt.

Similarly, drinking water was likely the single largest source of oxygen isotopes in pre-Hispanic central Mexican diets, and it would also have contributed to strontium values. Past drinking water sources were likely local to the Basin of Mexico. Although three of the five highland lakes on the Basin floor contained non-potable brackish water, the two southernmost lakes contained potable freshwater (Durazo & Farvolden, 1989). Moreover, numerous ethnohistorically-documented freshwater springs throughout the Basin were exploited for drinking water during pre-Hispanic times and into the Colonial era (Berdan & Anawalt, 1997; Morehart, 2016b; Sahagún, 1963). The available evidence thus indicates that strontium and oxygen isotope sources in past peoples' diets were predominantly local to the Basin, making radiogenic strontium and stable oxygen isotope analysis an effective means of reconstructing paleomobility and residential histories among the individuals interred at the Non-Grid 4 shrine.

Materials

Approximately 40% of the total number of individuals ($n=73$) were selected for biogeochemical analysis. Sampled individuals replicate the demographic structure (see Table 3) of the sacrificial deposit and were included in the sample based on a secure

association of the cranium with dentition. I compared intra-individual $^{87}\text{Sr}/^{86}\text{Sr}$ and $\delta^{18}\text{O}$ values across several dental and skeletal elements ($N=194$) representing discrete developmental periods to reconstruct sacrificial victims' patterns of mobility between distinct geologic or environmental zones over the life course (Ericson, 1985; Knudson et al., 2016; Luz & Kolodny, 1989; Price et al., 1994; Sealy et al., 1995). Tooth enamel forms during the first years of life and does not remodel, preserving isotopic signatures from the time of its initial mineralization, while bone remodels continuously throughout life (Hillson, 1996; Katzenberg, 2008). I sampled first molars ($n=73$), which mineralize in utero until three years of age, and third molars ($n=73$), which mineralize from 7 to 12 years of age, from all sampled individuals (Hillson, 1996). Additionally, I sampled cranial bone fragments ($n=48$) from a subset of sampled individuals with fragmented crania to limit destruction of osteological material. Cranial bones remodel at a rate of 1.8% per year in adults and reflect an average of the last 40-50 years of life (Pate, 1994; Torres-Lagares et al., 2010).

Methods

Radiogenic strontium and stable oxygen isotope systems are distributed differently across central Mexico (White et al., 2007). All archaeological enamel and bone samples were therefore analyzed for both $^{87}\text{Sr}/^{86}\text{Sr}$ and $\delta^{18}\text{O}$, allowing for further discrimination of individuals' movement between geological or climatically similar areas that would otherwise be indistinguishable using a single isotopic system (Knudson & Price, 2007). All samples were prepared for analysis at the Arizona State University (ASU) Archaeological Chemistry Laboratory.

Radiogenic strontium isotope analysis

Enamel and bone samples were mechanically cleaned to remove contaminants, as well as any dentine or trabecular bone. Additionally, mechanically cleaned cortical bone samples were chemically cleaned in a 0.8M acetic acid (CH_3COOH) solution and ashed at 800°C for 10 hours. Three milligrams of powdered tooth enamel powder or chemically cleaned and ashed bone were then dissolved in concentrated nitric acid (HNO_3) and diluted to a 2M stock solution.

Dissolved samples were analyzed at the Metals, Environmental, and Terrestrial Analytical Laboratory at ASU. An aliquot of stock solution was taken for major, minor, and trace elemental concentration analysis by a Thermo Fisher Scientific iCAP quadrupole inductively coupled plasma mass spectrometer (Q-ICP-MS). Strontium was then separated from the sample matrix with a PrepFAST automated low-pressure ion exchange chromatography system using an established method (Romaniello et al., 2015) to elute the sample matrix (2M HNO_3 + 1 wt% hydrogen peroxide [H_2O_2]) and strontium (6M HNO_3). Lastly, all other elements were removed using 1M hydrofluoric acid (HF). Aliquots from approximately 12% of samples were measured on the Q-ICP-MS before and after chemical purification to determine chemical recovery. Average chemistry yield was $84\% \pm 0.12$ ($n=24$).

The remaining portion of each strontium cut from the PrepFAST was dried in a Teflon beaker and digested with concentrated HNO_3 and 30% H_2O_2 to remove organics from the resin. Once digested, samples were dried down and reconstituted with 0.32M HNO_3 to a calculated constant concentration of 50 ppb Sr. Radiogenic strontium isotope

ratios were then measured on a Thermo-Finnigan Neptune multi-collector inductively coupled plasma mass spectrometer (MC-ICP-MS).

The MC-ICP-MS has nine Faraday cups capable of simultaneous ion beam measurement, and this instrument was configured with an Elemental Scientific, Inc. Apex Q high sensitivity sample introduction system with an Elemental Scientific, Inc. 50 or 100 $\mu\text{L}/\text{minute}$ PFA-ST microflow nebulizer. This instrument has seven 1011 amplifiers and three 1012 amplifiers which can be designated for any of the Faraday cups.

Data was collected by measuring 60 simultaneous ratios integrating 4.194 seconds each. Samples were corrected for on-peak blanks and in-line correction of the contributions of ^{84}Kr on ^{84}Sr and ^{86}Kr on ^{86}Sr using $^{83}\text{Kr}/^{84}\text{Kr}$ ratio of 0.201750 and $^{83}\text{Kr}/^{86}\text{Kr}$ ratio of 0.664533, after instrumental mass bias correction using a normalizing $^{88}\text{Sr}/^{86}\text{Sr}$ ratio of 8.375209. Sensitivity was 18 V on ^{88}Sr with a 50 ppb Sr solution, with ^{83}Kr values <0.0001 V. ^{85}Rb voltages were typically <0.01 V due to the low Rb/Sr initial ratios of the samples and effective chemical purification, but all data was interference corrected using a $^{85}\text{Rb}/^{87}\text{Rb}$ ratio of 2.588960, normalized to $^{88}\text{Sr}/^{86}\text{Sr}$ as above. Ratio outliers two standard deviations outside the mean were removed using a Matlab 2D-mathematical correction routine written by Dr. Stephen Romaniello now at University of Tennessee. Typical internal $^{87}\text{Sr}/^{86}\text{Sr}$ two standard error (SE) precision was $\sim 1\text{e-}6$.

Sequences included bracketing concentration matched SRM 987 standards. SRM 987 was run as a bracketing standard with a measured value of $^{87}\text{Sr}/^{86}\text{Sr}$ of 0.710257 ± 0.000020 (2σ , $n=60$), which is in agreement with reported values of $^{87}\text{Sr}/^{86}\text{Sr}$ of 0.710251 ± 0.000013 , (2σ , $n=5$; Balcaen et al., 2005). The analytical session included a sequence

incorporating SRM 987 standard in a range of variable concentration to verify $^{87}\text{Sr}/^{86}\text{Sr}$ values for samples. Reported values are all above the threshold for accurate $^{87}\text{Sr}/^{86}\text{Sr}$ values within the range of error of the bracketing standards. In addition, SRM 987 doped with varying amounts of calcium up to a ratio of Ca/Sr of 500 was run to simulate the accuracy and precision of isotope ratios in poorly purified samples with low yields. SRM 987 run at 50% concentration doped to a Ca/Sr of 500 was run as a check standard with a measured value of $^{87}\text{Sr}/^{86}\text{Sr}=0.710245 \pm 0.000028$ (2σ , $n=13$). NIST 1400 purified in parallel with samples had a measured value of 0.713100 ± 0.000051 (2σ , $n=7$), similar to the published value of 0.713150 ± 0.000016 (Galler et al., 2007). Approximately 5% of samples ($n=11$) were run in triplicate, with an average precision on individual measurements of ± 0.00004 (2σ).

Stable oxygen isotope analysis

Sample preparation for stable oxygen isotope analysis of archaeological hydroxyapatite carbonate ($\delta^{18}\text{O}_{\text{carbonate}}$) followed established methods (Koch et al., 1997). Samples of 15-20 mg of mechanically cleaned enamel or bone powder were treated with 0.04 mL of 2% bleach (NaOCl) per each milligram of sample and agitated on a mini-vortexer for 60 seconds and left to sit at room temperature for 24 hours. Samples were rinsed three times with 18.2 M Ω deionized water and treated with 0.04 mL of 0.1 M CH_3COOH per each milligram of tooth enamel or bone powder. Samples were mixed on a Mini-Vortexer for 60 seconds and left to sit at room temperature for 24 hours. Finally, samples were decanted and rinsed three times with 18.2 M Ω deionized water and then dried at 50°C for 24 hours.

After preparation in the ASU Archaeological Chemistry Laboratory, $\delta^{18}\text{O}_{\text{carbonate}}$ samples were analyzed at the Colorado Plateau Stable Isotope Laboratory at Northern Arizona University (NAU). Samples were analyzed using a Delta V Advantage isotope ratio mass spectrometer (IRMS) equipped with a Gas Bench II. Sequences included international standards of LSVEC, NBS-18, and NBS-19, and internal NAU laboratory standards of Joplin calcite (CC) and a calcium carbonate (CaCO_3) standard. Measured LSVEC values of $\delta^{18}\text{O}_{\text{carbonate(VPDB)}} = -26.16 \pm 0.20\text{‰}$ (1σ , $n=8$) correspond with published values of $\delta^{18}\text{O}_{\text{carbonate(VPDB)}} = -26.46 \pm 0.25\text{‰}$ (1σ , $n=10$, Stichler, 1995), while measured NSB-18 values of $\delta^{18}\text{O}_{\text{carbonate(VPDB)}} = -23.03 \pm 0.14\text{‰}$ (1σ , $n=14$) agree with reported values of $\delta^{18}\text{O}_{\text{carbonate(VPDB)}} = -23.04 \pm 0.17\text{‰}$ (1σ , $n=17$; Stichler, 1995), and measured NSB-19 values of $\delta^{18}\text{O}_{\text{carbonate(VPDB)}} = -2.20 \pm 0.19\text{‰}$ (1σ , $n=14$) agree with published values of $\delta^{18}\text{O}_{\text{carbonate(VPDB)}} = -2.20\text{‰}$ (Friedman et al., 1982; Hut, 1987). Similarly, measured Joplin CC values of $\delta^{18}\text{O}_{\text{carbonate(VPDB)}} = -23.49 \pm 0.19\text{‰}$ (1σ , $n=23$) correspond with reported values of $\delta^{18}\text{O}_{\text{carbonate(VPDB)}} = -23.41 \pm 0.19\text{‰}$ (1σ , $n=19$), and measured CaCO_3 standard values of $\delta^{18}\text{O}_{\text{carbonate(VPDB)}} = -12.99 \pm 0.18\text{‰}$ (1σ , $n=19$) correspond with reported values of $\delta^{18}\text{O}_{\text{carbonate(VPDB)}} = -13.01 \pm 0.14\text{‰}$ (1σ , $n=19$; Knudson et al., 2014). Approximately 6% of samples ($n=12$) were run in triplicate, with an average precision on individual measurements of $\pm 0.20\text{‰}$ (1σ).

Generated data are expressed in per mil (‰) using the standard formula:

$$\delta^{18}\text{O} = \left[\left(\frac{^{18}\text{O}/^{16}\text{O}_{\text{sample}}}{^{18}\text{O}/^{16}\text{O}_{\text{standard}}} \right) - 1 \right] \times 1,000 \text{ (Coplen et al., 1983; Craig, 1961).}$$

First molar $\delta^{18}\text{O}_{\text{carbonate}}$ values typically exhibit an average of $\sim 0.7\text{‰}$ $\delta^{18}\text{O}$ -enrichment as

a result of infant breastfeeding trophic level effects and were adjusted downwards (White et al., 2000; Wright & Schwarcz, 1998, 1999). All $\delta^{18}\text{O}_{\text{carbonate(VPDB)}}$ values were converted to $\delta^{18}\text{O}_{\text{drinking water(VSMOW)}}$ values for comparison with expected local $\delta^{18}\text{O}_{\text{mw(VSMOW)}}$ baseline ranges using the following equations: (1)

$$\delta^{18}\text{O}_{\text{carbonate(VSMOW)}} = (1.03091 \times \delta^{18}\text{O}_{\text{carbonate(VPDB)}}) + 30.91 \text{ (Coplen et al., 1983); (2)}$$

$$\delta^{18}\text{O}_{\text{phosphate(VSMOW)}} = (0.98 \times \delta^{18}\text{O}_{\text{carbonate(VSMOW)}}) - 8.5 \text{ (Iacumin et al., 1996); (3)}$$

$$\delta^{18}\text{O}_{\text{drinking water(VSMOW)}} = (1.54 \times \delta^{18}\text{O}_{\text{phosphate(VSMOW)}}) - 33.72 \text{ (Daux et al., 2008).}$$

It is important to note that while these formulae introduce some error into the oxygen values (Pellegrini et al., 2011; Pollard et al., 2011), they allow us to directly compare oxygen isotope values in hydroxyapatite to the oxygen isotope meteoric water values from established baselines.

Results

In 194 enamel and bone samples collected from 73 individuals, Ca/P values varied from 1.94 – 2.33, with a mean of Ca/P=2.03 \pm 0.12 (1 σ , n =194), U/Ca values varied from 3.21E-08 – 6.97E-05, with a mean of U/Ca=6.82E-06 \pm 1.30E-05 (1 σ , n =194), and Nd/Ca values varied from 2.24E-07 – 2.17E-06, with a mean of Nd/Ca=2.98E-07 \pm 3.15E-07 (1 σ , n =193; Appendix C). Observed $^{87}\text{Sr}/^{86}\text{Sr}$ values ranged from 0.70387 – 0.70738, with a mean of $^{87}\text{Sr}/^{86}\text{Sr}$ =0.70500 \pm 0.00061 (1 σ , n =194), and $\delta^{18}\text{O}_{\text{carbonate(VPDB)}}$ ranged from –11.26 – –0.38‰, with a mean of $\delta^{18}\text{O}_{\text{carbonate(VPDB)}}$ = –6.33 \pm 1.11‰ (1 σ , n =194; Appendix C).

Examining diagenetic contamination at Non-Grid 4

Results from major, minor, and trace elemental concentration analyses suggest little diagenetic contamination affected archaeological enamel and bone tissues recovered from the Non-Grid 4 shrine site. Most samples fall beneath the expected biogenic apatite ratio of $\text{Ca/P}=2.15$ (Price et al., 1992). Bone samples exhibited slightly higher Ca/P values, which may indicate at least some diagenetic contamination in these samples (see Appendix C). However, most enamel and bone observed U/Ca and Nd/Ca ratios fall below reported biogenic U/Ca values of $4.9\text{E-}05 \pm 1.8\text{E-}04$ (1σ , $n=52$; Knudson et al., 2014) and Nd/Ca values of $2.31\text{E-}06 \pm 3.52\text{E-}06$ (2σ , $n=14$; Knudson et al., 2012).

Similarly, comparison of analyzed enamel and bone samples' rare earth element (REE) and U concentrations (ppm) relative to Kamenov and colleagues' (2018) experimentally established Maximum Threshold Concentrations (MTC) for each of these elements indicates that the majority of samples exhibit no or only minor diagenetic contamination (Fig 7). REE and U concentrations in most samples fall beneath the MTC threshold of 1, indicating an expected biogenic chemical signal, or beneath the MTC threshold of 10, indicating expected low-grade diagenetic alteration that has not significantly affected the samples' biogenic Sr values (see discussion in Kamenov et al., 2018). Samples exceeding the MTC threshold of 10 exhibit advanced diagenesis that may indicate compromised Sr isotopic values. MTC sample values consistent with major isotopic alteration were primarily made up of U concentrations in bone samples. While this suggests these samples may exhibit significant diagenetic environmental contamination in Sr values, the same samples exhibited unaltered or only mildly altered

MTC values in all other analyzed REE. Furthermore, many of these bone samples with high U concentrations exhibit $^{87}\text{Sr}/^{86}\text{Sr}$ values distinct from the Non-Grid 4 shrine site expected local range of $^{87}\text{Sr}/^{86}\text{Sr}=0.70478 - 0.70484$, indicating that they preserve a pristine $^{87}\text{Sr}/^{86}\text{Sr}$ signal despite some contamination from the deposition environment.

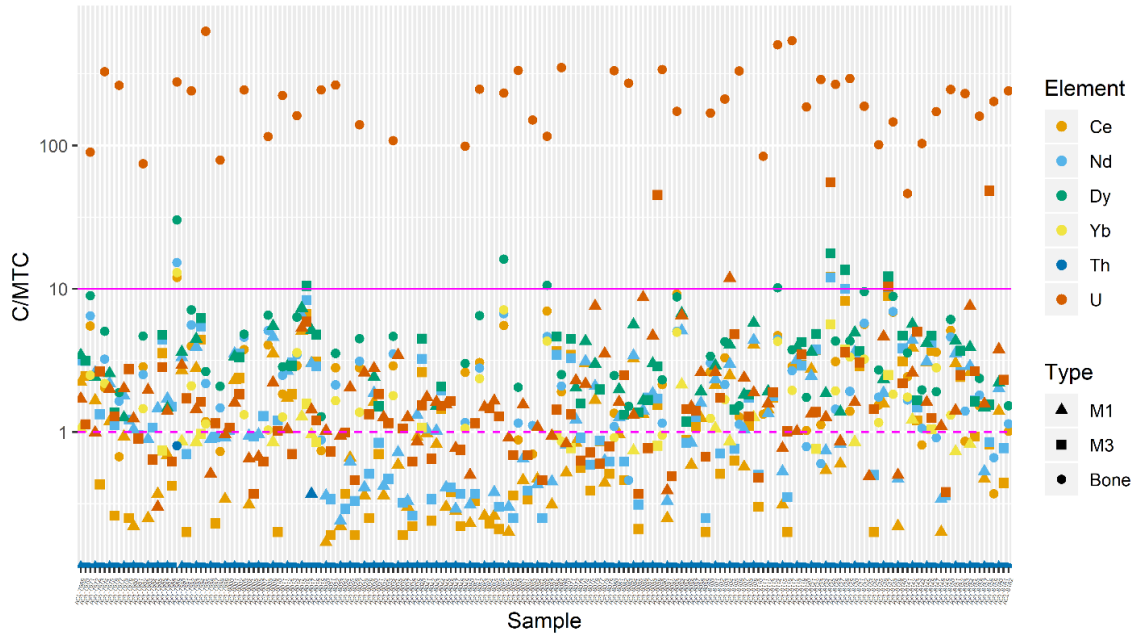


Figure 7. REE and U concentrations (ppm)/MTC for all analyzed human samples. Sample C/MTC values were plotted for each element (color). The dashed line demarcates the C/MTC threshold for biogenic values of 1, while the solid line demarcates the C/MTC threshold for minor diagenetic alteration of 10, following Kamenov et al. (2018). Sample type is indicated by shape.

Residential histories of sacrificial victims at Non-Grid 4

As previously discussed, individuals interred at Non-Grid 4 consuming primarily local sources of strontium were expected to exhibit $^{87}\text{Sr}/^{86}\text{Sr}$ values falling within the combined Basin of Mexico local range of $^{87}\text{Sr}/^{86}\text{Sr}=0.70465-0.70487$ ($n=60$; Pacheco-Forés et al., 2020). Similarly, individuals consuming local drinking water sources—the largest contributor to $\delta^{18}\text{O}$ values in human hard tissues—were expected to fall within a

local range of $\delta^{18}\text{O}_{\text{mw(VSMOW)}} = -12.6$ to -6.4‰ . Individuals with enamel or bone $^{87}\text{Sr}/^{86}\text{Sr}$ or $\delta^{18}\text{O}_{\text{mw(VSMOW)}}$ values falling outside of this conservative local range thus likely represent migrants.

A summary plot of $^{87}\text{Sr}/^{86}\text{Sr}$ and $\delta^{18}\text{O}$ results in all sampled tissues reveals that many of the individuals interred at the Epiclassic Non-Grid 4 shrine were migrants (Fig 8). A majority of first and third molar enamel samples fall outside the expected local Basin of Mexico $^{87}\text{Sr}/^{86}\text{Sr}$ range. Non-local $^{87}\text{Sr}/^{86}\text{Sr}$ values in these tissues, which form during infancy and early childhood (first molars), and early adolescence (third molars), indicate that these individuals originated from regions that were geologically distinct from the Basin of Mexico. This provides further supporting evidence of the long-debated presence of Epiclassic migrants within the Basin of Mexico.

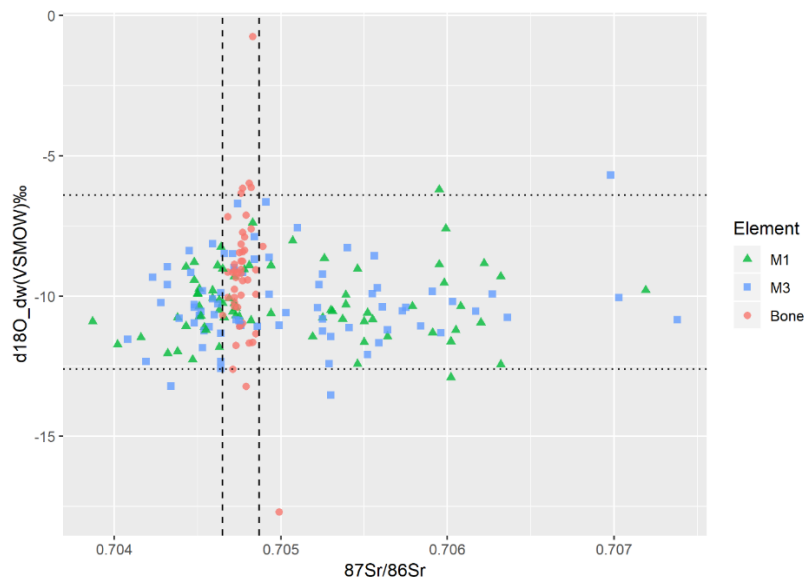


Figure 8. $^{87}\text{Sr}/^{86}\text{Sr}$ and $\delta^{18}\text{O}$ results across all sampled human elements ($N=194$). This includes first molars ($n=73$), third molars ($n=73$), and bone ($n=48$). The expected local $^{87}\text{Sr}/^{86}\text{Sr}$ range is indicated by the dashed lines, while the expected local $\delta^{18}\text{O}$ range is indicated by the dotted lines.

Nearly all of these enamel samples, however, fall within the $\delta^{18}\text{O}$ range for the Basin of Mexico, suggesting that these individuals hailed from locales environmentally indistinguishable from the Basin of Mexico. Conversely, the relative paucity of non-local $\delta^{18}\text{O}$ values could also indicate that the established Non-Grid 4 $\delta^{18}\text{O}$ local range, necessarily expanded by the MMD of 3.1‰ to account for interlaboratory variability in sample treatment, is too broad to capture meaningful variation in distinct environments within ancient Mesoamerica, which tend to range clinally in increments of 1‰ (Moreiras Reynaga, 2019; Wassenaar et al., 2009). In stark contrast to enamel tissues, all bone samples with the exception of two (ACL-7995 XAL-C4-4 and ACL-8126 XAL-C28-2), fall within the combined local Basin of Mexico $^{87}\text{Sr}/^{86}\text{Sr}$ and $\delta^{18}\text{O}$ ranges, suggesting that migrant sacrificed individuals lived locally within the Basin for a significant period of time prior to their deaths.

Interestingly, there is great diversity in non-local $^{87}\text{Sr}/^{86}\text{Sr}$ values among migrant and foreign individuals, who together make up 80% of sampled sacrificial victims. Observed non-local $^{87}\text{Sr}/^{86}\text{Sr}$ values range from $^{87}\text{Sr}/^{86}\text{Sr}=0.70387\text{-}0.70738$, falling significantly outside the expected Basin local range of $^{87}\text{Sr}/^{86}\text{Sr}=0.70465\text{-}0.70487$. Because $^{87}\text{Sr}/^{86}\text{Sr}$ values are not unique, but rather reflect the age and composition of local bedrock, it is not possible to geolocate from where individuals with non-local $^{87}\text{Sr}/^{86}\text{Sr}$ signatures originated. However, comparison of observed $^{87}\text{Sr}/^{86}\text{Sr}$ values in sacrificial victims with reported local $^{87}\text{Sr}/^{86}\text{Sr}$ ranges throughout Mesoamerica discussed above offer some suggestions as to potential source locations for Non-Grid 4 migrants.

Sacrificial victims with non-local $^{87}\text{Sr}/^{86}\text{Sr}$ values falling below the expected local Basin of Mexico range overlap with local ranges from the Maya volcanic highlands, western Mexico, and the Montagua Valley within the Maya metamorphic province (Hodell et al., 2004; Price et al., 2008). Similarly, sacrificed individuals with $^{87}\text{Sr}/^{86}\text{Sr}$ values falling above the expected Basin of Mexico range coincide with local $^{87}\text{Sr}/^{86}\text{Sr}$ ranges from the Puebla-Tlaxcala Valley and the Xochicalco Formation within central Mexico, as well as with further regions, including northern Mexico, the southern Maya lowlands, and the Oaxaca Valley (Hodell et al., 2004; Offenbecker, 2018; Pacheco-Forés et al., 2020). Countless other regions throughout Mesoamerica have yet to be characterized biogeochemically. Therefore, further investigations using biogeochemistry as well as other bioarchaeological methods are needed to better discern the source of the Non-Grid 4 Epiclassic migrants in the Basin of Mexico.

Examining intra-individual patterns in $^{87}\text{Sr}/^{86}\text{Sr}$ values among sacrificial victims at the Non-Grid 4 shrine allows us to reconstruct sacrificial victims' residential histories across the life course to gain a more nuanced understanding of mobility patterns at the site (Fig 9). A majority of sampled individuals (53%, $n=39$) exhibited non-local first molar $^{87}\text{Sr}/^{86}\text{Sr}$ values and local third molar and/or bone $^{87}\text{Sr}/^{86}\text{Sr}$ values, suggesting they were immigrants who were born in a geologically distinct region but moved to the Basin of Mexico later in life. Similarly, 27% ($n=20$) of sampled individuals exhibited non-local $^{87}\text{Sr}/^{86}\text{Sr}$ values among all sampled tissues, suggesting they were foreigners who lived outside the Basin of Mexico for most of their lives, relocating to the Basin only shortly before their deaths, before their bone tissues incorporated a significant Basin $^{87}\text{Sr}/^{86}\text{Sr}$

signature. However, due to sampling limitations, only two of these foreigners have data from bone samples. The remaining 18 “foreign” individuals therefore may or may not have exhibited local bone $^{87}\text{Sr}/^{86}\text{Sr}$ values.

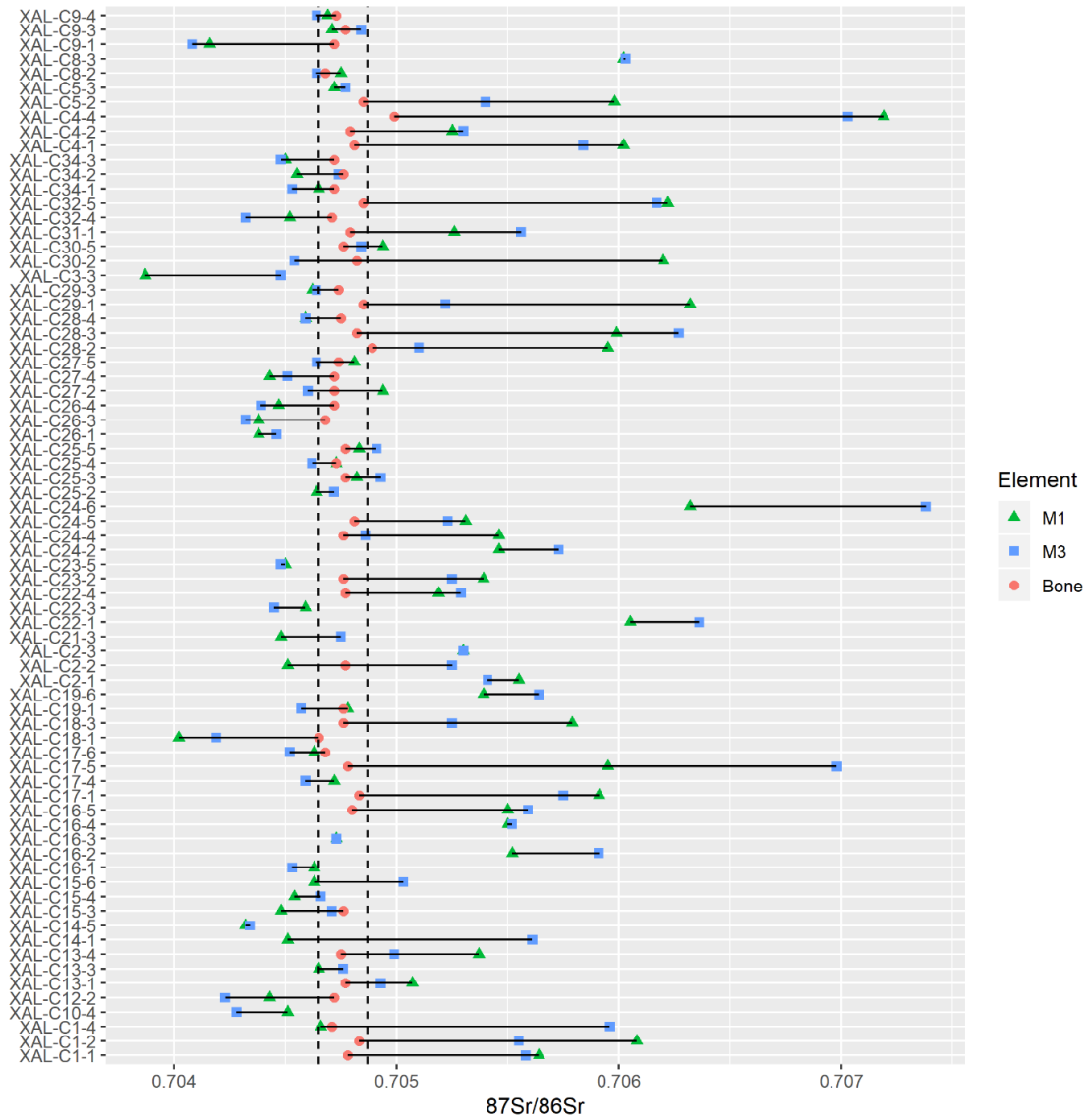


Figure 9. Intra-individual mobility among sacrificial victims interred at Non-Grid 4. Solid horizontal lines connect sampled skeletal elements taken from a single individual. Vertical dashed lines indicate the expected local $^{87}\text{Sr}/^{86}\text{Sr}$ range for the Basin of Mexico.

Furthermore, 12% ($n=9$) of sampled individuals were circular migrants who were born in the Basin but would leave and return periodically. Although the $^{87}\text{Sr}/^{86}\text{Sr}$ values in one individual (XAL-C17-4) suggest he was an emigrant (1% of the total sample) who was born in the Basin of Mexico and left later in life, this interpretation is based solely on enamel samples reflecting his early life and adolescence. He thus may or may not have also been a circular migrant. Finally, only 6% ($n=4$) of sampled individuals exhibited local Basin $^{87}\text{Sr}/^{86}\text{Sr}$ values among all sampled tissues, suggesting they were locals who lived in the region their entire lives. Thus, while some locals were present among the Non-Grid 4 shrine sacrificial victims, the vast majority of sampled sacrificed individuals were either immigrants or foreigners to the Basin of Mexico.

Identity-based violence at Non-Grid 4

Sacrificial victims' residential histories clarify the relationship between geographic origin, migration, and violence in central Mexico during the Epiclassic period. The high proportion of immigrant and foreign individuals (80%) and the low proportion of locals (6%) among the sampled sacrificial victims at Non-Grid 4 suggests that individuals with divergent geographic origins were more likely to suffer ritual violence during this period of socio-political reorganization.

Given that the Epiclassic is associated with migration into the Basin of Mexico, however, it is possible that local communities simply had higher numbers of migrants who, upon their deaths, may have been interred in normative as well as ritual contexts. Ideally, an examination of identity-based violence at Non-Grid 4 would compare the prevalence of migrants among ritual and normative burial contexts to see if migrants are

disproportionately represented in the sacrificial context. The lack of a normative burial context at the Non-Grid 4 site thus complicates interpretations of the prevalence of non-local individuals within the sacrificial deposits. Furthermore, the scarcity of skeletal material from Epiclassic Basin sites limits potential comparisons of Non-Grid 4 migrant prevalence with other normative burial assemblages from this time period.

In the absence of normative Epiclassic central Mexican burial contexts, I compare Non-Grid 4 paleomobility patterns with those from Classic period Teotihuacan. Teotihuacan was known to have maintained its population through immigration (Nichols, 2016; White, Storey, et al., 2004). I thus use the prevalence of migrants at Teotihuacan as a rough approximation for expectations for a normative Epiclassic community dynamic. Previous biogeochemical studies examining migration at Classic Teotihuacan residential compounds found that migrants made up approximately one-third of sampled individuals, ranging from 27% to 37% depending on the compound (Nado, 2017; Schaaf et al., 2012; Solís Pichardo et al., 2017; White, Storey, et al., 2004). Furthermore, biogeochemical studies of other ancient Mesoamerican cities in the Maya region found that migrants made up between 14% to 16% of sampled individuals at Calakmul and Tikal (Price et al., 2018; Wright, 2012). In contrast, at Non-Grid 4, 80% of individuals are migrants. Migrant individuals are thus present in sacrificial contexts in far greater numbers than would be expected in a community experiencing sustained immigration.

As discussed previously, extensive anthropological, behavioral science, evolutionary psychology, and peace studies research indicate that periods of socio-political upheaval are often accompanied by a rise in identity-based violence (Böhm et

al., 2016; Bowman, 1994; Haken et al., 2014; Howard, 2014; Kurin, 2016; Mamdani, 2002; Messner et al., 2019; Schmidt & Schröder, 2001; Schwandner-Sievers, 2001; Simunovic et al., 2013). In conjunction with ethnohistoric and archaeological evidence indicating that migrants often intentionally maintained and invoked visible material ties to their places of origin that may have set them apart in their new communities (e.g., Sahagún, 1961; Spence, 2005; Begun, 2013; Manzanilla, 2017), observed paleomobility patterns among sacrificial victims at Non-Grid 4 therefore strongly suggest that these individuals were victims of identity-based violence who were targeted for violence based on their divergent geographic origins in the balkanized socio-political landscape of the Epiclassic Basin of Mexico.

Conclusion

The biogeochemical data presented here provide compelling evidence that the individuals ritually sacrificed and interred at the Epiclassic Non-Grid 4 shrine site in the northern Basin of Mexico were victims of identity-based violence. A large majority of sampled individuals (80%) were born and lived their early lives outside of the Basin of Mexico, though most of these individuals relocated to the Basin and lived there for a significant period of time prior to their deaths (see Fig 9). I therefore provide direct evidence of first-generation migrants in the Epiclassic Basin of Mexico.

Furthermore, the relative paucity of sacrificial victims who were born and lived their entire lives in the Basin (6%) suggests that migrant individuals were disproportionately more likely to suffer ritual violence during the Epiclassic period. As geographic origin was an ethnohistorically and archaeologically demonstrated salient

indicator of social difference in pre-Hispanic Mesoamerica, these migrant individuals were most likely targeted for violence based on their divergent residential histories and geographic origins. This work further develops a theoretical framework within which to examine identity-based violence archaeologically, connecting broader research on identity and violence within the social sciences to bioarchaeologically observable aspects of identity in the past in contexts of large-scale violence. It thus contributes to anthropological knowledge of the social context of violence, examining the interaction between specific social identities; complex social processes like political fragmentation, migration, and demographic change; and instances of mass violence.

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CHAPTER 4

MIGRATION AND BIOLOGICAL CONTINUITY IN CENTRAL MEXICO DURING THE CLASSIC-EPICLASSIC TRANSITION

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The role of migration within the cultural development of central Mexico has long been debated. Archaeologically, migration has been implicated in the appearance of new forms of material culture within central Mexico (Z. Nelson, 2009; Rattray, 1966, 1987), as well as in sociopolitical disruptions including the dissolution of the Teotihuacan state (Cowgill, 2013, 2015a), and the emergence of new polities within central Mexico (Beekman, 2015; Beekman & Christensen, 2011). Moreover, indigenous accounts of migration play a central role in the ethnohistory of Postclassic central Mexican populations (see review in Beekman & Christensen, 2003), and Nahuatl, the language dominating central Mexico at the time of the Spanish Conquest, is not native to central Mexico (Hopkins, 1984; Kaufman, 2001; W. R. Miller, 1983), suggesting that migrants introduced it into the region. Despite multiple lines of evidence indicating that in-migration likely did take place within central Mexico, precise chronologies and trajectories of migration into the region remain unclear.

In this manuscript, I take a bioarchaeological approach to examine migration within central Mexico, evaluating possible sources of migrants into the region during the Epiclassic period (600-900 CE). I analyze a skeletal assemblage from Non-Grid 4, an Epiclassic ritual shrine site in the northern Basin of Mexico where approximately 180 individuals were sacrificed and interred (Meza-Peñaloza et al., 2019; Morehart et al.,

2012). Recent biogeochemical analyses found that among sampled individuals ($n=73$), 80% exhibited isotopic values identifying them as non-local migrants into the Basin of Mexico (Pacheco-Forés, Morehart, et al., in preparation). With the exception of three individuals interred in Cueva de las Varillas in Teotihuacan and identified biogeochemically as Epiclassic period migrants (Manzanilla, 2005a; Manzanilla et al., 1996; Price et al., 2000), the Non-Grid 4 individuals represent the first directly identified migrants within the Epiclassic Basin of Mexico. I use multivariate biodistance analyses to compare Non-Grid 4 individuals with Classic (250-600 CE) and Epiclassic skeletal populations throughout Mesoamerica to investigate the relationship between Non-Grid 4 individuals and pre-extant central Mexican groups, as well as from where migrant individuals at Non-Grid 4 may have originated.

Migration in Epiclassic Central Mexico

The Epiclassic period was an epoch of major social and political change in the central Mexican highlands (Diehl & Berlo, 1989; Rattray, 1996). In addition Teotihuacan's decline as a regional power (Anderson et al., 2016; Cowgill, 2015a, 2015b; Parsons & Sugiura, 2012) and the demographic reorganization of the Basin of Mexico (Parsons et al., 2008; Sanders et al., 1979), this period is associated with large-scale migrations into the Basin of Mexico. Here, migration is understood as the movement and long-term relocation of peoples across significant socio-cultural, political, or environmental boundaries (Tsuda et al., 2015).

Archaeological evidence of migration

The appearance of the Coyotlatelco red-on-buff ceramic complex in Epiclassic central Mexico has long been a central focus of migration debates within the archaeology of the region. This ceramic tradition occurs throughout central Mexico and is widely accepted as an Epiclassic chronological marker (Crider, 2011; Solar Valverde, 2006; Tozzer, 1921). Rattray (1966, 1996, 2006) argues that Coyotlatelco ceramics represent a complete break from previous ceramic forms in central Mexico, indicative of population replacement. Based on stylistic and technological similarities, many have suggested the Bajío region of north-central Mexico, which includes parts of the modern Mexican states of Michoacán, Guanajuato, Querétaro, Jalisco, and Aguascalientes as a possible source for the central Mexican Coyotlatelco ceramic complex (Beekman, 2015; Beekman & Christensen, 2003; Braniff, 2005; Cobean & Mastache, 1989; Healan, 2012; Healan & Cobean, 2019; Hernández, 2016; Hernández & Healan, 2019; Mastache et al., 2002). Other proposed origins for Coyotlatelco ceramics include the Malpaso Valley-La Quemada settlement system region of Zacatecas (Mastache et al., 2002; B. A. Nelson & Crider, 2005).

In contrast, others argue that Coyotlatelco ceramics developed directly out of Classic period Teotihuacan ceramic traditions (Fournier & Bolaños, 2007; Sanders, 1986, 1989; Sugiura Yamamoto, 2006). This argument of cultural continuity within central Mexico is used to bolster arguments against in-migration and support claims of biological continuity within the region during the Epiclassic period (Sanders, 2006). It thus

dismisses possible influences of northern and western Mexican groups on the cultural development of central Mexico (Sugiura Yamamoto, 2006).

More recently, scholars have identified important spatial and temporal variants in Coyotlatelco ceramics (Crider, 2011, 2013; Cyphers, 2000; Gaxiola González, 2006). These studies allow for the introduction of new styles through immigration into Epiclassic central Mexico, while also maintaining a degree of demographic and cultural continuity with Classic period Teotihuacan. Nevertheless, Cowgill (2013, 2015a) notes that migration is only one possible interpretation of the apparent stylistic break between Coyotlatelco and earlier Teotihuacan ceramic traditions. Other potential explanations of the stylistic discontinuity could include ideological, political, or moral distancing from the Teotihuacan state following its decline and the burning of its ritual center, or local emulations of foreign styles.

Interestingly, however, the distinctive change in ceramics during the Epiclassic also corresponds with other broad regional cultural changes. The appearance of Coyotlatelco coincides with changes in lithic production forms throughout central Mexico (Carballo, 2011; Z. Nelson, 2009; Rattray, 1987). Furthermore, lithic assemblages at many central Mexican Coyotlatelco sites are dominated by obsidian from the Ucareo-Zinapécuaro source area located in northeastern Michoacán within the Bajío region (Healan, 1997). Some authors therefore suggest that Coyotlatelco ceramics and Ucareo-Zinapécuaro sourced obsidian form two parts of a single intrusive material culture complex suggestive of migration into Epiclassic central Mexico (Cowgill, 2013; Hernández & Healan, 2019).

Additionally, archaeological surveys note major shifts in Epiclassic settlement patterns within central Mexico. Epiclassic settlements are primarily located on dispersed hilltops, in contrast with Classic period valley floor settlements (Mastache et al., 2002; Parsons et al., 2008; Sanders et al., 1979). Furthermore, many Classic period central Mexican settlements believed to be Teotihuacan colonies are abandoned during the Epiclassic (Healan & Cobean, 2019). While demographic shifts could be a result of the collapse of the Teotihuacan state (Cowgill, 2013), the abandonment of many structures at Teotihuacan itself with later Epiclassic reoccupations (Cabrera Castro & Gómez Chavez, 2008; Rattray, 2006) suggests that these cultural changes could be a result of an influx of migrants into central Mexico during the Epiclassic period.

Bioarchaeological evidence of migration

Bioarchaeological investigations of migration within Epiclassic central Mexico have been hampered by the limited availability of relevant Epiclassic skeletal assemblages. Biogeochemical studies identifying immigrants within central Mexico during the Classic period abound (e.g., Buckley, Storey, Hirth, Kennett, & Culleton, 2016; Manzanilla, 2005; Nado, 2017; Price et al., 2000; Schaaf et al., 2012; Solís Pichardo et al., 2017; White, Spence, Longstaffe, Stuart-Williams, & Law, 2002; White, Spence, Stuart-Williams, & Schwarcz, 1998; White, Storey, Longstaffe, & Spence, 2004). To date, however, only three central Mexican Epiclassic individuals interred in the Cueva de las Varillas at Teotihuacan have been directly identified as migrants using biogeochemical methods (Price et al., 2000).

Genetic analyses of central Mexican populations also provide support for an Epiclassic migration into central Mexico. Nuclear DNA studies of the Epiclassic Cueva de las Varillas burials, as well as of Epiclassic Cueva del Pirul burials at Teotihuacan found genetic discontinuities between these individuals and antecedent Classic period burials from residential and ritual contexts at Teotihuacan, suggesting they were migrants from outside of central Mexico (Manzanilla, 2005a). Interestingly, these Epiclassic Teotihuacan burials were also genetically distinct from contemporaneous Epiclassic individuals from Chapatongo, Hidalgo in the Tula region (Fournier & Vargas Sanders, 2002), indicating that they may have migrated from farther away. Similarly, in a mitochondrial DNA study of Postclassic central Mexican burials, Morales-Arce and colleagues (2019) find evidence of genetic divergence from Classic central Mexican populations (Aguirre-Samudio et al., 2017; Álvarez-Sandoval et al., 2015), suggesting a population replacement took place during the Epiclassic period. Conclusions from these studies, however, are limited by the scarcity of available comparative ancient DNA data from ancient Mesoamerica. As such, they do not indicate possible source regions for hypothesized migration events.

Examining Epiclassic migrations through biodistance analysis

Biodistance analyses of phenotypic variability among pre-Hispanic Mesoamerican populations provide essential further clarification of potential Epiclassic migrations. Biodistance is the measure of biological relatedness or divergence between individuals or populations based on genetic relationships (Buikstra, 1980; Larsen, 1999; Stojanowski & Schillaci, 2006). Biodistance analyses use phenotypic variation in

polygenic skeletal and dental traits to estimate the degree of genetic relatedness and reconstruct microevolutionary processes such as genetic drift and gene flow within and between populations (Buikstra et al., 1990; Cheverud, 1988; Harpending & Jenkins, 1973; Relethford & Blangero, 1990; Stojanowski & Schillaci, 2006).

Analyses of between-group phenotypic variation have been used to elucidate population histories and evaluate migration hypotheses within Mesoamerica. Studies using cranial non-metric skeletal variants (Beekman & Christensen, 2003; Christensen, 1997), facial morphology (González-José et al., 2007), dental morphology (Aubry, 2009; Gómez-Valdés, Bautista Martínez, Pompa y Padilla, Talavera González, & Castrejón-Caballero, 2008; Ragsdale & Edgar, 2018), and cervicometrics (Aubry, 2009) all find evidence of biological discontinuity between central Mexican Classic and Postclassic skeletal populations. These findings provide further support for migration and population replacement during the Epiclassic period.

Importantly, these studies also suggest possible origins of hypothesized migrants. The majority of these studies found that Postclassic central Mexican groups shared the closest biological affinities with Postclassic populations from northwestern (Beekman & Christensen, 2003; Gómez-Valdés et al., 2008; González-José et al., 2007; Ragsdale & Edgar, 2018) and northern (Aubry, 2009; Christensen, 1997) Mexico. These findings are consistent with archaeological models that suggest the Bajío or Malpasos Valley-La Quemada settlement system region as possible sources of Coyotlatelco ceramics and Epiclassic migrants. It is important to note, however, that González-José and colleagues (2007) found Postclassic central Mexican groups did not share close affinities with other

sampled populations, including northern populations, but hypothesized potential migrant origins in west Mexico based on lacunae in the study sample. Furthermore, while these studies provide compelling evidence for population replacements coinciding with the Epiclassic period, none of the studies examined skeletal remains dating to the Epiclassic.

Migration at the Non-Grid 4 Epiclassic shrine site

The skeletal remains recovered from the Non-Grid 4 shrine site provide an invaluable opportunity to further elucidate Epiclassic migration into central Mexico. The shrine site consists of an anthropogenic raised platform with associated sacrificial deposits of at least 180 human crania located in the now-extinct Lake Xaltocan in the northern Basin of Mexico (Meza-Peñaloza et al., 2019; Morehart et al., 2012). AMS radiocarbon dates of recovered botanical remains ranging from 550-890 CE, as well as the presence of Coyotlatelco ceramics confirm the site dates to the Epiclassic period (Morehart et al., 2012).

Biogeochemical analyses have directly identified large numbers of first-generation immigrants among the individuals interred at Non-Grid 4. An analysis of radiogenic strontium ($^{87}\text{Sr}/^{86}\text{Sr}$) and stable oxygen ($\delta^{18}\text{O}$) isotopes found that in a sample ($n=73$) of sacrificial victims, approximately 80% of sampled individuals exhibited isotope signatures identifying them as non-local migrants into the Basin of Mexico (Pacheco-Forés, Morehart, et al., in preparation). While the Non-Grid 4 migrants are among the first Epiclassic migrants to be biogeochemically identified within central Mexico, such analyses are limited in the information they provide about migrants' geographic origins. Radiogenic strontium and stable oxygen isotope values are not

unique, but instead reflect local geologic and hydrologic variation (Faure & Powell, 1972; Gat, 1996). It is therefore not possible to conclusively associate non-local isotopic signatures with specific locations within Mesoamerica.

Recent biodistance analyses shed more light on the potential origins of the Non-Grid 4 migrants. In a study of cranial non-metric traits at Non-Grid 4, Meza-Peñaloza and colleagues (2019) found that the Non-Grid 4 individuals ($n=118$) were biologically discontinuous with Formative and Classic, as well as other Epiclassic populations from central Mexico. Similarly, in a craniometric study, García Velasco (2019) found Non-Grid 4 individuals ($n=65$) to be biologically distinct from other central Mexican as well as northern Mexican populations. The latter analysis, however, was limited by the presence of cranial modification among 63% of Non-Grid 4 crania included in the study. Because cranial modification represents a substantial alteration of heritable phenotypic variation in the craniometric dimensions considered in the study (Anton, 1989; Cheverud et al., 1992; Cheverud & Midkiff, 1992; L. A. P. Kohn et al., 1993; Püschel et al., 2020), it precludes the examination of biological affinities among modified individuals and comparative populations (García Velasco, 2019, p. 124). Therefore, in this study, I use phenotypic variation in cervicometric tooth dimensions (Hillson et al., 2005; Stojanowski, 2007), which are robust to cultural body modifications, to examine biological distances between Non-Grid 4 individuals and antecedent and contemporaneous Mesoamerican populations to discern potential source populations for the Non-Grid 4 migrants.

Materials and Methods

A total of 333 individuals from Non-Grid 4 and comparative Classic and Epiclassic sites throughout Mexico were analyzed in this study (Table 4). Most potential regions of origin are represented in the comparative Mesoamerican sample, including central, southern, and northern Mexico, as well as the Bajío and Malpaso Valley-La Quemada settlement system regions (Fig 10). Although many sites included in the analysis have occupations preceding or postdating the Classic and Epiclassic periods, only burials dating to these two periods were included in the comparative sample. While precise chronologies are not available for the northern Mexican sites of Cusihuirachi, Sierra de la Fragua, and Charcas, they are believed to be from the late pre-Hispanic period and are included in the analysis to ensure an important potential source of Epiclassic migrants was represented in the analysis.

Table 4. Sites and sample sizes of individuals included in the study.

| Site (Abbreviation) | N | Time period | Region | Location^a |
|--|----------|------------------------|-----------------|-----------------------------|
| Teotihuacan, Mexico State | | | | |
| La Ventilla B (Teo-LVB) | 51 | Classic | Central Mexico | INAH-DAF |
| Tlajinga 33 (Teo-TLJ) | 28 | Classic | Central Mexico | ASU TRL |
| Cuevas del Pirul y de las Varillas (Teo-CUE) | 3 | Epiclassic | Central Mexico | UNAM |
| Non-Grid 4, Mexico State (NG4) | 115 | Epiclassic | Central Mexico | UNAM |
| Tula, Hidalgo (TUL) | 7 | Epiclassic | Central Mexico | INAH-DAF |
| Cholula, Puebla (CHL) | 19 | Epiclassic | Central Mexico | INAH-DAF |
| Los Pilarillos, Zacatecas (PIL) | 14 | Epiclassic | Malpaso Valley | LQMVAP |
| El Ocote, Aguascalientes (OCO) | 10 | Epiclassic | Bajío | INAH-Aguascalientes |
| Monte Albán, Oaxaca (MAB) | 55 | Epiclassic | Southern Mexico | INAH-DAF |
| Cusihuirachi, Chihuahua (CHI) | 17 | Epiclassic-Postclassic | Northern Mexico | NMNH |
| Sierra de la Fragua, Coahuila (COH) | 8 | Epiclassic-Postclassic | Northern Mexico | NMNH |
| Charcas, San Luis Potosí (SLP) | 6 | Epiclassic-Postclassic | Northern Mexico | NMNH |

^a INAH-DAF=Instituto Nacional de Antropología e Historia-Dirección de Antropología Física, Mexico City; ASU TRL = Arizona State University Teotihuacan Research Laboratory, San Juan Teotihuacan; UNAM=Universidad Nacional Autónoma de México, Mexico City; LQMVAP=La Quemada Malpaso Valley Archaeological Project, Zacatecas; NMNH=National Museum of Natural History, Washington D.C

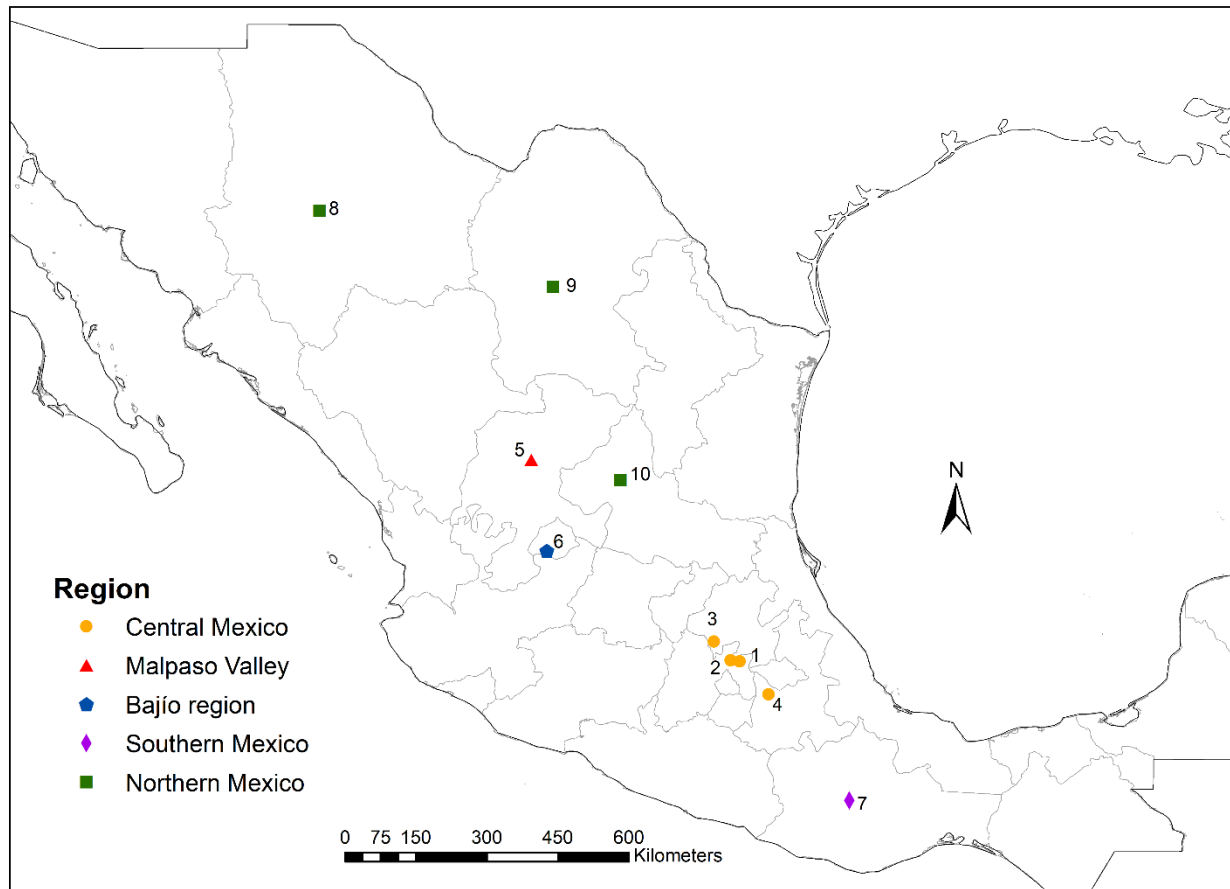


Figure 10. Sites and regions included in the analysis. 1=TEO, 2=NG4, 3=TUL; 4=CHL, 5=PIL, 6=OCO, 7=MAB, 8=CHI, 9=COH, 10=SLP.

Moreover, burials from the central Mexican site of Teotihuacan were treated as three separate contexts for the purposes of the analysis. Epiclassic period Pirul and Varillas cave burials were segregated from Classic period burials, as they represent a temporally distinct occupation (Manzanilla et al., 1996). Similarly, Classic period burials were divided into two groups—those from the elite residential compound of La Ventilla B (Cabrera Castro & Gómez Chavez, 2008) and those from the craft specialist residential compound Tlajinga 33 (Widmer, 1987). Various studies show Teotihuacan residential compounds were predominantly occupied by kin groups of consanguinal and affinal relatives (Millon, 1976; Spence, 1974, 1994; Storey, 1992), indicating that these two populations were unlikely to have engaged in mate exchange and were thus effectively two distinct biological populations.

Cervicometric mesio-distal and bucco-lingual dimensions were collected from all individuals, following Hillson and colleagues (2005), with updates in premolar and molar caliper placement following Aubry (2014). All measurements were taken on the left side, using antimere substitution for unobservable or missing teeth. Cervicometric dimensions are highly correlated with crown dimensions, exhibiting similar degrees of phenotypic variability and producing similar patterns of biological affinity as crown measurements (Hillson et al., 2005; Stojanowski, 2007). In addition to being more resistant to dental attrition, cervicometric dimensions are typically unaffected by cultural practices of body modification including cranial and dental modification, both of which occur among Non-Grid 4 individuals.

Collected cervicometric variables exhibiting high intra-observer error were eliminated (Utermohle & Zegura, 1982). Missing data were minimized through antimere substitution, and cases and variables with high levels of missing data were removed to produce a data matrix with less than 40% missing data. Examination of margin plots of missing data distributions suggests that most values were missing at random. I used multiple imputation to create and analyze 10 multiply imputed datasets ($m=10$; Fig 11). Multiple imputation improves accuracy and statistical power relative to other missing data imputation techniques by accounting for the missing data uncertainty (Schafer, 1999; Schafer & Olsen, 1998; van Buuren, 2018). Missing values were imputed under fully conditional specification, using the random forest settings of the mice 3.0 package in R (R Core Team, 2016; van Buuren & Groothuis-Oudshoorn, 2011). The random forest imputation method is a machine learning technique that can accommodate nonlinearities and interactions and does not require a particular regression model to be specified. This method has been shown to perform well in the imputation of complex data sets where some individuals have missing data and the data do not conform to a normal distribution (Shah et al., 2014). Parameters of interest were estimated separately for each imputed dataset and pooled following Rubin (1987).

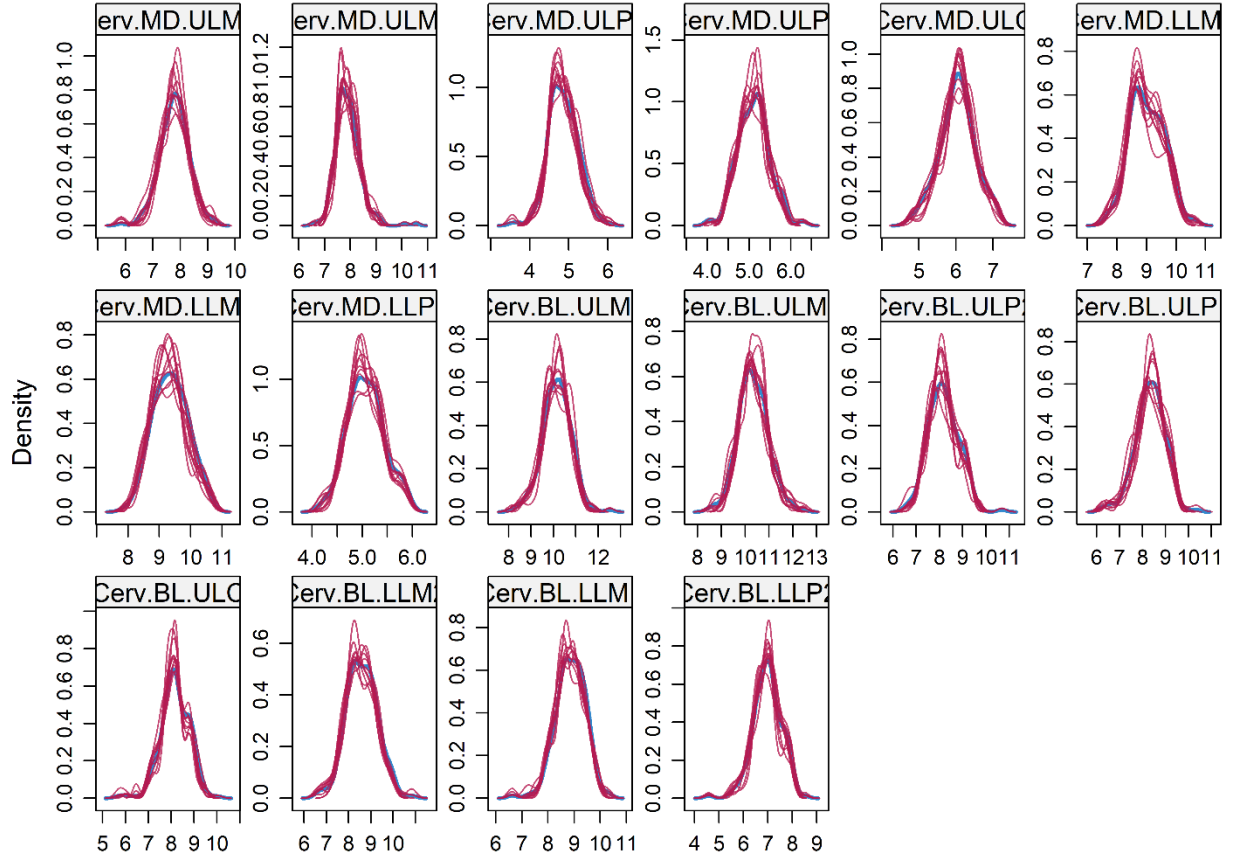


Figure 11. Density plots of imputed data by variable. Imputed data (red) conform to the same distribution of the observed data (blue). Matching distributions indicate the imputed values are plausible value.

Imputed data sets were subjected to pre-analysis data treatments. Size differences due to sexual dimorphism and allometric effects were minimized through a Q-mode correction (Corruccini, 1973). I then used a principal components analysis to generate uncorrelated variables from Q-mode corrected cervicometric data to eliminate redundancy in the phenotypic data. Resulting uncorrelated eigenvectors were used in all further analyses.

Finite Mixture Analysis at Non-Grid 4

Previous geometric morphometric analysis by García Velasco (2019) found three morphologically distinct groups among Non-Grid 4 individuals, one of which isolated individuals exhibiting cranial modification. The presence of two groups among unmodified individuals indicates that there is greater phenotypic—and therefore likely greater genetic—diversity at Non-Grid 4 than would normally be expected from a single population of consanguinal and affinal relatives. García Velasco’s identification of at least two biological groups, in conjunction with biogeochemical results indicating the presence of a handful of local individuals as well as non-local individuals from multiple geologically distinct regions (Pacheco-Forés, Morehart, et al., in preparation), suggests that Non-Grid 4 individuals likely originated from multiple distinct sources. I therefore used finite mixture analysis to reveal latent biological population structure among Non-Grid 4 individuals to further refine my ability to pinpoint potential migrant source populations.

Finite mixture models provide powerful but flexible tools to detect the presence of multiple biological groups within a larger sample (McLachlan & Peel, 2000). This analysis assumes no prior knowledge of the number of groups presence, or individuals’ sex or group affiliation (Algee-Hewitt, 2016). Instead, the model iteratively selects the most likely number of groups present in the sample via maximum likelihood estimation using the expectation-maximization algorithm (McLachlan & Krishnan, 1997). Additionally, the analysis generates estimated posterior probabilities of group membership for each individual for k groups. I performed a finite mixture analysis on all

observable Non-Grid 4 individuals ($n=115$) in R using the MCLUST 4.3 mixture clustering package (Fraley & Raftery, 2002, 2012). I used estimated posterior probabilities to assign individuals to k groups for subsequent analysis.

Relationship (R) matrix analysis

I used each imputed dataset ($m=10$; $n=333$), including Non-Grid 4 k groups, to conduct a relationship (R) matrix analysis in RMET 5.0. R-matrix analysis evaluates the extent of regional differential external gene flow, as well as the average genetic differentiation among populations (Harpending & Ward, 1982; Relethford, 2003; Relethford et al., 1997; Relethford & Blangero, 1990). The analysis produces two parameters of interest. The first, the F_{ST} statistic, ranges between 0 and 1 and represents overall levels of genetic variation within a regional mating network. Values closer to 0 indicate high levels of mate exchange between populations, while values closer to 1 indicate very limited mate exchange and increased genetic isolation. The second parameter of interest is estimated genetic distance matrix between individual populations. Using equal population size weighting schemes and an average heritability of $h^2=0.62$ for odontometric traits following Stojanowski (2005a, 2005b), I generated inter-site genetic distance matrices for each imputed dataset. Distance matrices, unbiased F_{ST} statistics, and their respective standard errors for m imputations were then averaged into a pooled set of results. This method combines the results from m imputations into a single set of results that reflects the statistical uncertainty due to missing data (Rubin, 1987). The pooled distance matrix was visualized using multidimensional scaling and the ggplot2 package in R (Wickham, 2009).

Results

Top estimated *BIC* values for finite mixture candidate models for all Non-Grid 4 imputed datasets was $k=2$ (Table 5). Non-Grid 4 individuals were thus classified into two groups for each iterated dataset based on generated cluster membership posterior probabilities for best fitting models. Final sample sizes for each Non-Grid 4 group (designated as Non-Grid 4 A and Non-Grid 4 B) are reported in Table 5.

Table 5. Finite mixture analysis top *BIC* values and k group solutions for m imputed datasets.

| <i>m</i> | Top model <i>BIC</i> | <i>k</i> | Non-Grid 4 A (<i>n</i>) | Non-Grid 4 B (<i>n</i>) |
|----------|----------------------|----------|---------------------------|---------------------------|
| 1 | -4345.3 | 2 | 45 | 70 |
| 2 | -4342.2 | 2 | 69 | 46 |
| 3 | -4459.7 | 2 | 53 | 62 |
| 4 | -4409.2 | 2 | 98 | 17 |
| 5 | -4470.3 | 2 | 37 | 78 |
| 6 | -4390.1 | 2 | 53 | 62 |
| 7 | -4539.8 | 2 | 95 | 20 |
| 8 | -4273.0 | 2 | 83 | 32 |
| 9 | -4409.0 | 2 | 51 | 64 |
| 10 | -4413.8 | 2 | 44 | 71 |

The unbiased pooled average genetic differentiation among all populations was $F_{ST}=0.05$. This statistic differed significantly from 0 ($se=0.01$; $t=4.69$, $p<.001$). Pooled estimated genetic distance values between populations and their associated standard errors are presented in Table 6. Z-distribution tests were used to determine if inter-site estimated genetic distances were statistically significant (Table 7). Many calculated p -values of intersite estimated genetic distances do not fall under the $\alpha=0.05$ threshold, indicating they are not statistically significantly different from 0. This is likely due to

small sample sizes among many of the reference populations. I discuss them below regardless, as they are suggestive of trends, even if not statistically significant.

Table 6. Pooled inter-site estimated genetic distances^a (upper triangle) and associated standard errors (lower triangle).

| Site ^b | Teo-LVB | Teo-TLJ | Teo-CUE | NG4A | NG4B | TUL | CHL | PIL | OCO | MAB | CHI | COH | SLP |
|-------------------|---------|-------------|-------------|-------------|-------------|-------------|-------------|-------------|-------------|-------------|-------------|-------------|-------------|
| Teo-LVB | - | 0.06 | 0.26 | 0.10 | 0.12 | 0.07 | 0.08 | 0.04 | 0.05 | 0.02 | 0.04 | 0.08 | 0.12 |
| Teo-TLJ | 0.02 | - | 0.05 | 0.07 | 0.09 | 0.12 | 0.05 | 0.02 | 0.07 | 0.04 | 0.03 | 0.11 | 0.05 |
| Teo-CUE | 0.13 | 0.10 | - | 0.18 | 0.20 | 0.31 | 0.10 | 0.09 | 0.31 | 0.22 | 0.15 | 0.22 | 0.17 |
| NG4A | 0.02 | 0.02 | 0.12 | - | 0.16 | 0.12 | 0.09 | 0.10 | 0.13 | 0.07 | 0.08 | 0.14 | 0.13 |
| NG4B | 0.03 | 0.03 | 0.12 | 0.03 | - | 0.17 | 0.13 | 0.08 | 0.15 | 0.09 | 0.11 | 0.16 | 0.17 |
| TUL | 0.05 | 0.06 | 0.17 | 0.06 | 0.07 | - | 0.12 | 0.13 | 0.08 | 0.09 | 0.11 | 0.18 | 0.16 |
| CHL | 0.03 | 0.03 | 0.11 | 0.03 | 0.04 | 0.07 | - | 0.07 | 0.11 | 0.06 | 0.08 | 0.12 | 0.13 |
| PIL | 0.03 | 0.03 | 0.12 | 0.04 | 0.04 | 0.08 | 0.04 | - | 0.06 | 0.05 | 0.02 | 0.09 | 0.07 |
| OCO | 0.04 | 0.04 | 0.16 | 0.05 | 0.05 | 0.07 | 0.06 | 0.05 | - | 0.04 | 0.04 | 0.11 | 0.08 |
| MAB | 0.01 | 0.02 | 0.13 | 0.02 | 0.02 | 0.05 | 0.03 | 0.03 | 0.04 | - | 0.04 | 0.08 | 0.10 |
| CHI | 0.03 | 0.03 | 0.12 | 0.03 | 0.04 | 0.07 | 0.04 | 0.03 | 0.04 | 0.03 | - | 0.13 | 0.06 |
| COH | 0.05 | 0.06 | 0.15 | 0.06 | 0.06 | 0.10 | 0.06 | 0.06 | 0.07 | 0.05 | 0.07 | - | 0.21 |
| SLP | 0.07 | 0.06 | 0.15 | 0.07 | 0.07 | 0.10 | 0.08 | 0.07 | 0.08 | 0.06 | 0.06 | 0.11 | - |

^a D² values with significant associated *p*-values bolded.

^b Abbreviations based on those in Table 4.

Table 7. Z-scores of inter-site estimated genetic distances (upper triangle) and associated *p*-values^a (lower triangle).

| Site ^b | Teo_LVB | Teo_TLJ | Teo_CUE | NG4A | NG4B | TUL | CHL | PIL | OCO | MAB | CHI | COH | SLP |
|-------------------|---------|---------|---------|------|------|------|------|------|------|------|------|------|------|
| Teo_LVB | - | 2.65 | 1.98 | 4.39 | 4.49 | 1.37 | 2.57 | 1.43 | 1.32 | 1.58 | 1.63 | 1.61 | 1.78 |
| Teo_TLJ | 0.01 | - | 0.54 | 3.07 | 3.35 | 1.82 | 1.71 | 0.83 | 1.48 | 1.97 | 1.00 | 1.92 | 0.86 |
| Teo_CUE | 0.05 | 0.59 | - | 1.49 | 1.63 | 1.81 | 0.87 | 0.81 | 1.93 | 1.74 | 1.21 | 1.45 | 1.11 |
| NG4A | 0.00 | 0.00 | 0.13 | - | 5.48 | 2.01 | 2.95 | 2.65 | 2.68 | 3.61 | 2.62 | 2.42 | 1.94 |
| NG4B | 0.00 | 0.00 | 0.10 | 0.00 | - | 2.46 | 3.35 | 2.26 | 2.88 | 4.04 | 2.99 | 2.59 | 2.28 |
| TUL | 0.17 | 0.07 | 0.07 | 0.04 | 0.01 | - | 1.73 | 1.77 | 1.09 | 1.62 | 1.62 | 1.86 | 1.58 |
| CHL | 0.01 | 0.09 | 0.38 | 0.00 | 0.00 | 0.08 | - | 1.61 | 1.93 | 2.11 | 1.94 | 1.81 | 1.72 |
| PIL | 0.15 | 0.41 | 0.42 | 0.01 | 0.02 | 0.08 | 0.11 | - | 1.19 | 1.56 | 0.44 | 1.41 | 0.99 |
| OCO | 0.19 | 0.14 | 0.05 | 0.01 | 0.00 | 0.28 | 0.05 | 0.23 | - | 1.16 | 0.86 | 1.50 | 1.01 |
| MAB | 0.11 | 0.05 | 0.08 | 0.00 | 0.00 | 0.11 | 0.03 | 0.12 | 0.25 | - | 1.49 | 1.60 | 1.55 |
| CHI | 0.10 | 0.32 | 0.23 | 0.01 | 0.00 | 0.11 | 0.05 | 0.66 | 0.39 | 0.14 | - | 1.94 | 0.85 |
| COH | 0.11 | 0.05 | 0.15 | 0.02 | 0.01 | 0.06 | 0.07 | 0.16 | 0.13 | 0.11 | 0.05 | - | 1.97 |
| SLP | 0.08 | 0.39 | 0.27 | 0.05 | 0.02 | 0.11 | 0.09 | 0.32 | 0.31 | 0.12 | 0.40 | 0.05 | - |

^a A significant *p*-value indicates that the D^2 value (Table 6) is statistically different from 0.

^b Abbreviations based on those in Table 4.

Inter-site estimated genetic distances are visualized in Fig 12. While the ordination plot gives a good overall summary of genetic variation among analyzed populations, it compresses genetic variation within central Mexican populations and thus does not exactly reflect estimated genetic distances (Table 6). For example, while Non-Grid 4 B should plot most closely to Pilarillos ($D^2=0.08$, $p=0.023$) populations from the Malpaso Valley-La Quemada settlement system region, followed by Tlajinga 33 ($D^2=0.09$, $p<0.001$) in central Mexico, and Monte Albán ($D^2=0.09$, $p<0.001$) in southern Mexico based on estimated genetic distances, the central Mexican sites of Cholula ($D^2=0.13$, $p<0.001$) and Non-Grid 4 A ($D^2=0.16$, $p<0.001$) appear to be among its nearest neighbors on the ordination plot. For this reason, all subsequent discussion of inter-site genetic distances will focus on the values reported in Table 6 rather than on the ordination plot.

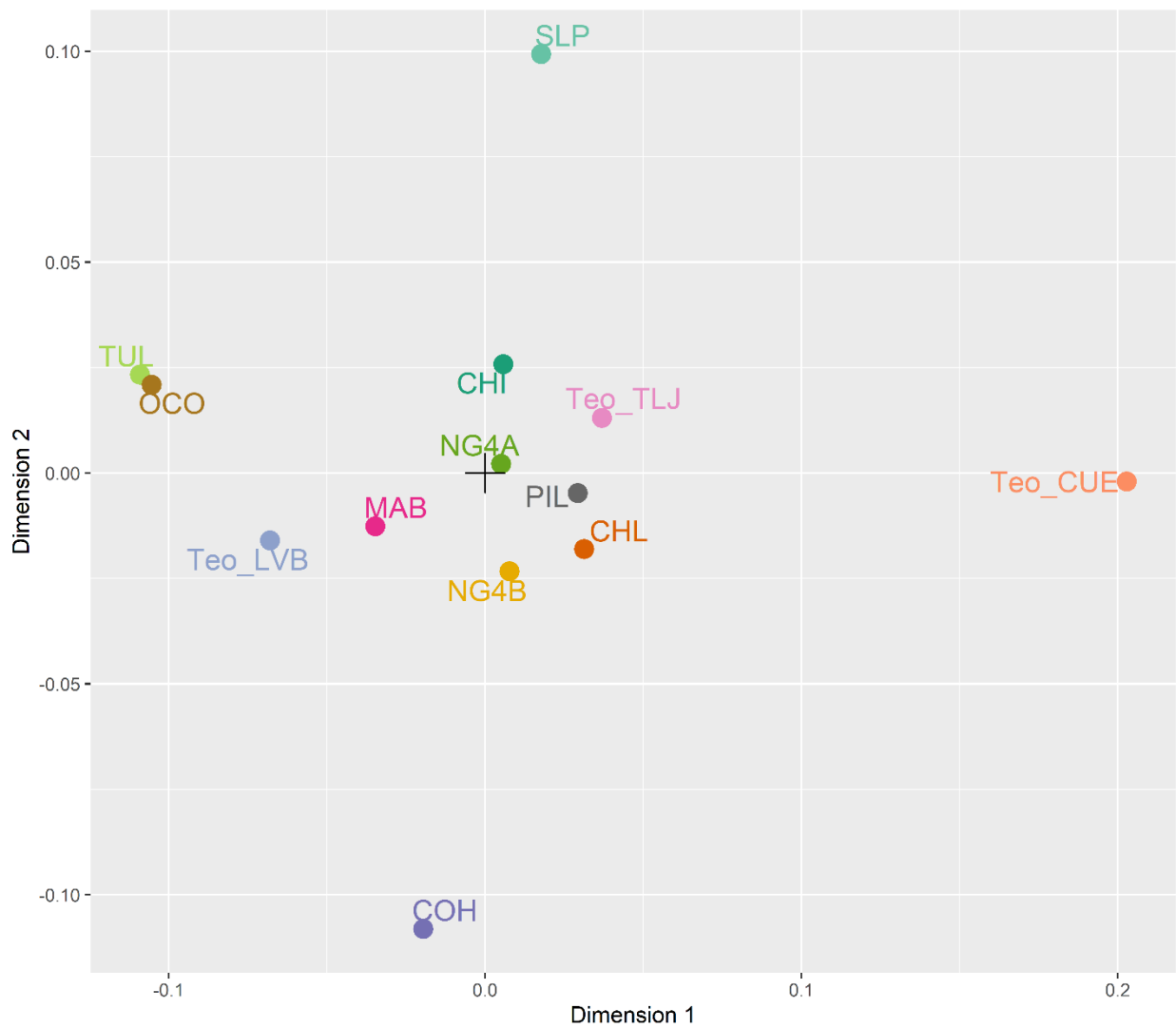


Figure 12. Multidimensional scaling of inter-site estimated genetic distances. Site abbreviations in Table 4. Note that the ordination compresses variation among central Mexican sites, making them appear more similar than estimated genetic distances indicate.

Discussion

The pooled average genetic differentiation among populations of $F_{ST}=0.05$ indicates that analyzed Classic and Epiclassic Mesoamerican populations are significantly different. The low magnitude of the statistic, however, suggests that there

are high levels of gene flow between populations. When considered alongside biodistance and mtDNA studies finding that cultural differences—more so than geography or time period—best explain patterns of biological differentiation in ancient Mesoamerica (Morales-Arce et al., 2019; Willermet et al., 2013), this result provides some support to archaeological and bioarchaeological models of Epiclassic migrations within Mesoamerica. Greater mate exchange between populations implies greater movement of peoples.

Nevertheless, the F_{ST} is not necessarily meaningful in isolation. Instead, it establishes a baseline level of diversity among sampled populations. An examination of changes in overall levels of genetic variation over time (see Stojanowski, 2009) could better characterize changes in levels of inter-populational mate exchange from the Classic to the Epiclassic periods, bolstering support for Epiclassic migration hypotheses. Unfortunately, such an analysis is limited here by a lack of sufficient Classic-period samples.

Consideration of Non-Grid 4 variation at the intra-site level indicates that at least two biologically distinct populations were interred at the site. Finite mixture analysis results dividing Non-Grid 4 individuals into two distinct groups further clarify García Velasco's (2019) finding that phenotypic diversity at Non-Grid 4 was likely indicative of multiple biological populations. This is further supported by the relatively large estimated genetic distance between the two Non-Grid 4 groups ($D^2=0.16$, $p<.001$), which was greater than 80% of all estimated genetic distances in the analysis. By dividing Non-Grid

4 individuals into two groups, this analysis reveals important biological variation within the Non-Grid 4 sample.

Continuity or replacement in Epiclassic Central Mexico?

Archaeological discussions of migration are often polarized between binary models of cultural and biological continuity or replacement. Inter-site estimated genetic distances between the five Epiclassic central Mexican groups (Non-Grid 4 A and B, the Teotihuacan Cuevas burials, Tula, and Cholula) and previous Classic period central Mexican populations (the Teotihuacan La Ventilla B and Tlajinga 33 burials), however, demonstrate that central Mexican population structure during the Classic-Epiclassic transition was far more complex. The three Cuevas individuals are the most genetically isolated group of any considered in the R-matrix analysis (Fig 12). When considering Classic-Epiclassic biological relationships, the Cuevas-La Ventilla B genetic distance is among the largest in the analysis ($D^2=0.26$, $p=.047$). Although the Cuevas-Tlajinga 33 distance is among the smallest ($D^2=0.05$, $p=.592$), it was not statistically significant different from 0, likely due to small sample size. Price and colleagues' (2000) biogeochemical identification of at least three migrants among the Cuevas individuals, however, further supports the finding that the Cuevas individuals were indeed biologically distinct from at least some Classic period populations at Teotihuacan.

Other Epiclassic central Mexican populations range from being relatively genetically distinct from Classic central Mexican populations (Non-Grid 4 B-La Ventilla B, $D^2=0.12$, $p<.001$; Tula-Tlajinga 33, $D^2=0.12$, $p=.068$) to sharing a greater genetic

affinity with them (Cholula-Tlajinga 33, $D^2=0.05$, $p=.087$; Non-Grid 4 A-Tlajinga 33, $D^2=0.07$, $p=.002$; Tula-La Ventilla B, $D^2=0.07$, $p=.171$). This again suggests some level of biological continuity within Epiclassic central Mexico, while also providing evidence of external gene flow—and an accompanying influx of migrants—into the region during the Classic-Epiclassic transition. However, I again note the need for greater representation of Classic period central Mexican samples in future analyses, as I have here only a limited representation of regional Classic period biological diversity.

These results contradict previous archaeological models dismissing the influence of migrants on central Mexican sociopolitical development (e.g., Sanders, 1986, 2006; Sugiura Yamamoto, 2006). Similarly, they also do not support prior models of complete population replacement (e.g., Rattray, 1966, 1996, 2006). The results suggest that rather than assessing Epiclassic central Mexico population structure through polarized models of biological continuity or replacement, a perspective of biological admixture may be more appropriate.

Epiclassic migrant origins at Non-Grid 4 and beyond

Interestingly, results show that Epiclassic central Mexican groups are generally more biologically distinct from each other than they are from Classic central Mexican populations at Teotihuacan. The Cuevas group is particularly divergent from other Epiclassic central Mexican groups (Cuevas-Tula, $D^2=0.31$, $p=.070$; Cuevas-Non-Grid 4 B, $D^2=0.20$, $p=.103$; Cuevas-Non-Grid 4 A, $D^2=0.18$, $p=.135$). Similarly, Non-Grid 4 B is genetically distinct from Tula ($D^2=0.17$, $p=.014$), and as previously discussed, the two

Non-Grid 4 groups are biologically distinct from each other. This suggests that although increased biological diversity into central Mexico is likely consistent with an influx of migrants, these migrants do not represent a biological monolith. Biogeochemical analyses of both Cuevas and Non-Grid 4 individuals support this finding, as identified migrants from both sites exhibit clear differences in isotopic values reflective of several geologically distinct geographic origin locales (Pacheco-Forés, Morehart, et al., in preparation; Price et al., 2000). Instead, migrants appear to be arriving in central Mexico from multiple distinct biological source populations.

Inter-site estimated genetic distances also provide valuable insight when considering possible source populations for Epiclassic migrants into central Mexico. Cuevas individuals share the greatest biological affinity with individuals from Los Pilarillos in the Malpaso Valley-La Quemada settlement system region of north-central Mexico ($D^2=0.09$, $p=.053$). Non-Grid 4 A individuals are most biologically similar to groups from Monte Albán in southern Mexico ($D^2=0.07$, $p<.001$) and Cusihuiriachi in northern Mexico ($D^2=0.08$, $p=.009$). Non-Grid 4 B individuals are most genetically similar to those from Los Pilarillos ($D^2=0.08$, $p=.024$) and Monte Albán ($D^2=0.09$, $p<.001$). Although I focus on sites where biogeochemical analyses have confirmed the presence of migrants (Cuevas and Non-Grid 4), it is interesting to note that Epiclassic individuals at Tula are most similar to those from El Ocote in the Bajío region of north-central Mexico ($D^2=0.08$, $p=.276$), and individuals at Cholula share the greatest biological affinity with Monte Albán ($D^2=0.06$, $p=.035$).

The biological distances discerned here reflect possible source populations for Epiclassic migrants in central Mexico. While these results further support archaeological and bioarchaeological models suggesting Epiclassic migrants came from northern Mexico, including the Bajío and Malpaso Valley-La Quemada settlement system regions, they also suggest potential Epiclassic migration sources not previously considered by archaeologists.

For example, although migration from southern Mexico is rarely included in discussions of central Mexican Epiclassic migrations, individuals from Non-Grid 4 (both A and B) and Cholula all shared close biological affinities with individuals from Epiclassic Monte Albán. Material evidence at Non-Grid 4 further supports this finding. Excavations of the Non-Grid 4 shrine recovered numerous matte finish gray ladle ceramic censers (Morehart et al., 2012). Such censers are uncommon within Classic period central Mexican ceramic assemblages, but strongly resemble Epiclassic period Xoo phase ceramics found at Monte Albán and throughout the Oaxaca Valley in southern Mexico (Martínez López et al., 2000).

The presence of Oaxacan migrants is well documented at the Classic period Teotihuacan residential compound of Tlailotlacan (Price et al., 2000; Spence et al., 2005; White, Spence, et al., 2004), as well as at Teotihuacan-established colonies to the north in the Tula region (Healan & Cobean, 2019; Spence, 2005). Additionally, Xoo phase ceramics recovered from other Epiclassic sites within central Mexico were interpreted as evidence of a southern Mexican merchant enclave (Smith & Lind, 2005). The biological

and cultural evidence thus not only suggest that Oaxaca may indeed have been a source of migrants into central Mexico during the Epiclassic period, but also that Oaxacan diaspora networks persisted in some form throughout the Classic/Epiclassic transition.

These analyses therefore indicate that Epiclassic migrants in central Mexico may have originated in the Bajío or Malpaso Valley-La Quemada settlement system regions of northern Mexico or in the Oaxaca Valley in southern Mexico. However, it is also possible that close genetic distances between these groups reflect limitations in the analyzed sample, which include small sample sizes from northern Mexico and a lack of samples from western Mexico—an important hypothesized source of Epiclassic migrants (e.g., Beekman & Christensen, 2003; Gómez-Valdés et al., 2008). Thus, while the analyses have clarified certain aspects of the role of migration during the Classic-Epiclassic transition in central Mexico, more work remains to be done.

Conclusion

The biodistance analyses presented here demonstrate the complexities of central Mexican population structure during the Classic-Epiclassic transition. Analyses indicate some degree of biological continuity within central Mexico during this transition, while also showing support for an influx of migrants. Thus, rather than supporting archaeological models of biological continuity or replacement within central Mexico, I find support for population admixture during the Epiclassic period.

Furthermore, I find that Epiclassic populations within central Mexico were themselves biologically diverse, indicating that purported Epiclassic migrants originated

from various source populations throughout Mesoamerica. While findings support archaeological and bioarchaeological models positing a northern Mexican origin for these migrants, including the Bajío and Malpaso Valley-La Quemada settlement system regions, they also find evidence of migration from the Oaxaca Valley in southern Mexico. Future work would benefit greatly from further examination of central Mexican connections with southern Mexico during the Epiclassic, as well as the inclusion of western Mexican Epiclassic skeletal samples and more samples from northern Mexico.

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CHAPTER 5

CONCLUSION

The preceding chapters build upon a persistent theme in interpreting the prehistory of ancient Mesoamerica: migration has long been a subject of archaeological research within central Mexico. The vast majority of research has focused on discerning the role of migration in the dramatic social, political, and cultural changes that swept central Mexico during the Epiclassic period. Migration has been implicated in the declining regional political and economic dominance of the Teotihuacan state to varying degrees (Cowgill, 2013, 2015a; Manzanilla, 2015), as well as in the emergence of new polities within the region (Beekman, 2015; Beekman & Christensen, 2011; Healan & Cobean, 2019).

Scholars have sought to reconstruct evidence for Epiclassic central migrations using material correlates in the archaeological record. Central among these is the appearance of the Coyotlatelco ceramic industry in central Mexico during the Epiclassic period (Crider, 2011; Solar Valverde, 2006; Tozzer, 1921). Most authors agree that these ceramics bear little resemblance to antecedent ceramic industries in central Mexico (Cowgill, 2013; Crider, 2013; Rattray, 1966, 1996, 2006), although countervailing opinions do exist (Fournier & Bolaños, 2007; Sanders, 1986, 2006; Sugiura Yamamoto, 2006). Instead, Coyotlatelco ceramics strongly resemble older ceramic industries from outside of central Mexico. Scholars have posited possible origins for Coyotlatelco in the Bajío (Beekman & Christensen, 2003, 2011; Healan, 2012; Hernández, 2016; Hernández

& Healan, 2019) and Malpaso Valley-La Quemada settlement system (Mastache et al., 2002; B. A. Nelson & Crider, 2005) regions of northwestern Mexico.

Archaeologists have also noted other changes in material culture coincident with the appearance of Coyotlatelco ceramics. The Epiclassic also saw the advent of new forms of lithic production throughout central Mexico (Carballo, 2011; Z. Nelson, 2009; Rattray, 1987). These techniques produced projectile points that were distinct from previous Classic period artifacts (Cowgill, 2015a, p. 111). Additionally, central Mexican Epiclassic sites with large numbers of Coyotlatelco ceramics also appeared to be exploiting obsidian sources distinct from those preferred during the Classic period (Healan, 1997). Interestingly, Ucareo-Zinapécuaro, the obsidian source dominating these assemblages, is located in the Bajío region, coinciding with one of the proposed origins of Coyotlatelco ceramics (Hernández & Healan, 2019). Scholars have thus hypothesized that these combined shifts in material culture patterns make up an intrusive material culture complex suggestive of migration into Epiclassic central Mexico (Cowgill, 2013, 2015a; Hernández & Healan, 2019).

Cognizant of critiques of cultural-historical archaeology and its oversimplified approach to migration, however (see Cabana, 2011), many researchers have been hesitant to equate changes in ceramic forms with the incursion of new peoples into central Mexico (Cowgill, 2013; Crider, 2013). Such changes could also be explained by the local emulation of foreign material forms, as well as by an intentional social, political, or moral distancing from previous material forms that may have fallen out of vogue with

Teotihuacan's declining regional political and economic influence during the Epiclassic (Cowgill, 2013). Distinguishing the subtle differences between the introduction of a foreign style by migrants and local emulation of a foreign style, however, is difficult to discern archaeologically and little work has been done on the topic to date.

As such, archaeologists have sought to incorporate additional lines of evidence to further support Epiclassic migration hypotheses. Extensive regional survey data reveal stark discontinuities between Classic and Epiclassic settlement patterns in central Mexico. While the majority of Classic period settlements were located in clusters on the alluvial plain, most Epiclassic settlements are found on dispersed hilltops (Healan, 2012; Healan & Cobean, 2019; Mastache et al., 2002; Parsons et al., 2008; Sanders et al., 1979). Many Classic period settlements were abandoned during the Epiclassic, including sectors of Teotihuacan itself (Cabrera Castro & Gómez Chavez, 2008; Healan & Cobean, 2019; Rattray, 2006). Of the Classic settlements that persisted, few show evidence of interaction with new Epiclassic sites, as inferred through the lack of appreciable amounts of Coyotlatelco ceramics (Healan & Cobean, 2019). The association of new settlement patterns with the introduction of novel forms of material culture thus also suggested the arrival of a migrant population.

Art historical evidence further supports Epiclassic migration hypotheses. Public monumental art during the Epiclassic period is characterized primarily by its hybridity and eclecticism. In particular, murals and public monuments at Epiclassic sites in central Mexico have been described as stylistically foreign. This eclecticism has been attributed

to emulation of foreign art styles (Nagao, 1989, 2014), and it has also been held up as an additional line of evidence supporting Epiclassic migration hypotheses within central Mexico (Turner, 2019).

While these data strongly suggest of substantial migration into central Mexico during the Epiclassic period, George Cowgill noted that “the most decisive evidence about migrations will probably eventually come through bioarchaeological methods” (2013, p. 143). To date, one biogeochemical study has directly identified three Epiclassic migrant individuals interred at Teotihuacan (Price et al., 2000). Interestingly, although ancient DNA (Morales-Arce et al., 2019) and biodistance analyses (Aubry, 2009; Beekman & Christensen, 2003; Christensen, 1997; Gómez-Valdés et al., 2008; González-José et al., 2007; Ragsdale & Edgar, 2018) have produced results that also indicate biological discontinuity between Classic and Postclassic period central Mexican peoples and likely population replacement due to migration, few have examined Epiclassic individuals directly (but see Manzanilla, 2005a). This is due to the rarity of Epiclassic skeletal remains recovered from central Mexican sites.

The discovery and excavation of a minimum of 180 individuals who were sacrificed and interred at the Epiclassic Non-Grid 4 shrine site in the northern Basin of Mexico (Morehart et al., 2012) thus provide a wealth of data for central Mexican migration studies. Biodistance analyses suggests that the individuals represented multiple biological kinship groups (García Velasco, 2019) who were also biologically distinct from previous central Mexican populations (Meza-Peñaloza et al., 2019).

In addition to providing much needed bioarchaeological data on migration in populations living in Epiclassic central Mexico, the Non-Grid 4 individuals also provide a unique opportunity to examine the consequences of such population movements on the migrants themselves. Thus, the goals of this dissertation were threefold:

1. To develop a theoretical framework for investigating identity-based violence in the past using bioarchaeological methods (Chapter 3). Bioarchaeological studies are highly successful in accessing multivalent past social identities, but applications of social identity theory to contexts of violence to understand who is targeted for violence are rare (but see Kurin, 2014, 2016).
2. To identify first-generation immigrants at Non-Grid 4 by reconstructing individuals' residential histories using biogeochemical methods. This involves creating a radiogenic strontium isotope baseline for central Mexico (Chapter 2) in addition to analyzing hard tissue samples from Non-Grid 4 using radiogenic strontium ($^{87}\text{Sr}/^{86}\text{Sr}$) and stable oxygen ($\delta^{18}\text{O}$) isotopic systems (Chapter 3).
3. To characterize patterns of biological affinity at Non-Grid 4 using multi-scalar biodistance analyses to evaluate archaeological models of migration and population dynamics in Classic-Epiclassic central Mexico and Mesoamerica (Chapter 4).

Summary of Results

Chapter 2 creates a baseline of bioavailable radiogenic strontium values for central Mexico. A total of 63 plant and water samples collected from 13 sites across

central Mexico are analyzed for $^{87}\text{Sr}/^{86}\text{Sr}$ on a Thermo-Finnigan Neptune multi-collector inductively couple plasma mass spectrometer (MC-ICP-MS). These data are analyzed alongside 16 previously published soil and faunal bone $^{87}\text{Sr}/^{86}\text{Sr}$ values from two additional central Mexican sites. A five-cluster k-means model is generated to determine which regions of central Mexico can and cannot be distinguished isotopically using $^{87}\text{Sr}/^{86}\text{Sr}$ values.

Results demonstrate that although the two clusters falling within the Basin of Mexico within central Mexico overlap in their expected local $^{87}\text{Sr}/^{86}\text{Sr}$ ranges, many locations within the Basin are distinguishable using site-specific $^{87}\text{Sr}/^{86}\text{Sr}$ local ranges. In the context of central Mexican migration studies, the use of cluster or site-specific $^{87}\text{Sr}/^{86}\text{Sr}$ local ranges should thus be determined by the research question and the scale of mobility to be examined. This study contributes to central Mexican paleomobility studies by expanding knowledge of strontium isotope variability within the region, ultimately allowing researchers to detect intra-regional residential mobility and gain a greater understanding of the sociopolitical interactions within and between central Mexico and other regions of Mesoamerica.

Chapter 3 introduces a theoretical framework for examining identity-based violence in the past. It explores how during periods of scarcity and heightened intergroup competition, universally observed social dynamics like in-group favoritism and out-group bias can coalesce around specific salient social identities and escalate into identity-based

violence. This framework is then applied to understand identity-based violence at Non-Grid 4.

Ethnohistoric and archaeological data indicate that geographic origin was a culturally significant indicator of social difference in pre-Hispanic Mesoamerica. Biogeochemical analyses of $^{87}\text{Sr}/^{86}\text{Sr}$ and $\delta^{18}\text{O}$ isotopes, along with the biogeochemical $^{87}\text{Sr}/^{86}\text{Sr}$ developed in Chapter 2 are therefore used to reconstruct a sample ($n=73$) of sacrificial victims' residential histories to examine whether residential mobility predisposed individuals to suffer violence. Results find that 80% of sacrificed individuals at Non-Grid 4 were born outside of the Basin of Mexico. Furthermore 39 of these foreign-born individuals (53% of the total sample) are identified as first-generation immigrants. In contrast, only 6% of sampled individuals are identified as having lived locally within the Basin of Mexico their whole lives. Although there are no normative central Mexican Epiclassic burial contexts against which to compare these percentages, the overwhelming majority of foreign-born individuals strongly suggests individuals' divergent geographic origins and residential histories predisposed them to be targeted for ritual violence.

This work expands knowledge of migration processes in Epiclassic central Mexico. It uses biogeochemical methods to directly identify a large group of Epiclassic immigrants into central Mexico, providing decisive support for Epiclassic migration models. Furthermore, through the development of a theoretical framework within which to examine identity-based violence bioarchaeologically, it refocuses the central Mexican

Epiclassic migration debate on the migrants themselves, exploring the human impact of population movement and individual residential mobility, which included an increased vulnerability to suffer ritual violence.

Chapter 4 reconstructs patterns of biological affinity among Non-Grid 4 sacrificial victims ($n=115$) and a comparative sample of Classic ($n=79$) and Epiclassic ($n=139$) Mesoamerican populations. It uses multi-scalar biodistance analyses to characterize central Mexican population structure and dynamics during the Epiclassic period. Finite mixture analysis of Non-Grid 4 individuals reveals latent biological population structure within the sample indicating that two biologically distinct kinship groups are represented among the sacrificial victims. When considered alongside biogeochemical results presented in Chapter 3 finding a wide range of non-local $^{87}\text{Sr}/^{86}\text{Sr}$ values in identified migrant individuals, this suggests that migrants at Non-Grid 4 came from multiple origin communities.

Relationship (R) matrix analysis of Non-Grid 4 subpopulations and comparative Classic and Epiclassic Mesoamerican populations reveal a complex picture of population structure in Epiclassic central Mexico. Archaeological models of Epiclassic migration into the region are divided between those that emphasize biological continuity and dismiss the influence of migrants and those that posit complete population replacement of Classic central Mexican populations with Epiclassic migrants. Estimated genetic distances between Classic and Epiclassic central Mexican populations, however, demonstrate that the biological reality of Epiclassic central Mexico was far more

complex, with some support for migration as well as for biological continuity. Thus, rather than assessing Epiclassic central Mexico through the polarized lens of biological continuity or replacement, it is best understood from the perspective of biological admixture.

Additionally, R matrix analysis results point to potential origin communities for central Mexican Epiclassic migrants. Estimated genetic distances support archaeological models suggesting Epiclassic populations in central Mexico originated from the Bajío and Malpaso Valley-La Quemada settlement system regions in northwestern Mexico. However, they also indicate migration from the Oaxaca Valley in southern Mexico may have taken place during the Epiclassic—a somewhat unexpected result. This indicates that Oaxacan migration and diaspora networks may have persisted beyond archaeologically well documented ethnic enclaves dating to the Classic period (e.g., Spence, 2005) into the Epiclassic period and beyond.

Final Thoughts and Future Directions

This dissertation reconstructs patterns of Epiclassic central Mexican migration through paleomobility and biodistance analyses of individuals ritually sacrificed and interred at the Non-Grid 4 shrine site in central Mexico. Additionally, it examines the interaction between perceptions of social difference and ritual violence to determine whether individuals with distinct categorical social identities were at greater risk of suffering violence during the Epiclassic period, a period of dramatic socio-political upheaval and reorganization.

This study reaffirms the efficacy and importance of employing multiple lines of evidence to investigate past migrations. The biogeochemical and biodistance results presented here provide direct support for Epiclassic central Mexican migration models. Not only are migrants directly identified using biogeochemical techniques, but their potential places of origin are proposed using R matrix analysis. Similarly, the integration of archaeological data in these analyses is essential, further clarifying and contextualizing results. For example, in the absence of archaeological data, unexpected results such as the presence of southern Mexican migrants within Epiclassic central Mexico, may have been dismissed as an artifact of small sample size or a non-representative sample. Instead, these results are now an exciting direction for future research.

Moreover, this study demonstrates the promise of identity-based violence as a subject of bioarchaeological inquiry. Although a growing body of anthropological and broader social science research indicates that the perception of social difference can motivate instances of violence today (e.g., Bowman, 1994; Bush & Keyman, 1997; Schmidt & Schröder, 2001), little work has been done to examine such identity-based violence in the past (but see Kurin, 2014, 2016). This project thus provides a theoretical framework within which to examine identity-based violence in the past, allowing bioarchaeologists to understand who was most vulnerable to violence in past societies.

Additionally, examining identity-based violence in the past provides a deep time perspective to social scientists seeking to understand the topic today. For example, although the research presented here investigates the complex relationship between

political instability, migration, and violence in ancient Mesoamerica, many of these processes are increasingly relevant to today's world. Similarly volatile social processes are increasingly evident in the United States and around the world, manifest in the rising tides of nationalism, xenophobia, and violence in response to increasing globalization, demographic change, and mass migration (Bogel-Burroughs, 2019; Cai et al., 2019; Cai & Landon, 2019; Kulish, 2011). By understanding the causes and consequences of identity-based violence in the past, I can connect the past to the present to understand patterns of marginalization and violence through time. This allows researchers to recognize the persistence of identity-based violence as well as the profound challenges to be met in altering patterned violent behaviors.

More importantly, recognizing past instances of identity-based violence shows that although such violence did exist in the past, it is rare in the archaeological record. This dissertation highlights numerous examples of multi-ethnic societies within ancient Mesoamerica that benefitted greatly from sustained migration (e.g., Healan & Cobean, 2019; Manzanilla, 2015; Nichols, 2016; Winter, 1998). When examined within this broader context of migration, it becomes clear that identity-based violence at Non-Grid 4 was a response to a particular set of circumstances that included migration as well as political fragmentation and increased competition for access to resources and land. Identity-based violence is therefore neither ubiquitous nor an inevitable outcome of diverse human interaction in the past, present, or future.

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APPENDIX A
AUTHOR CONTRIBUTIONS

Chapter 2 of this dissertation was adapted from the article “Expanding Radiogenic Strontium Isotope Baseline Data for Central Mexican Paleomobility Studies” previously published in the journal PLoS ONE (2020). This article was coauthored by the author of this dissertation, Gwyneth W. Gordon, and Kelly J. Knudson. Both Gordon and Knudson have granted their permission for a modified version of this publication to be included in this dissertation.

APPENDIX B

CENTRAL MEXICAN BASELINE $^{87}\text{SR}/^{86}\text{SR}$ AND PROVENANCE DATA

| Laboratory Number | Site | Material | $^{87}\text{Sr}/^{86}\text{Sr}$ | UTM-E | UTM-N | Altitude (masl) | Cluster |
|-------------------|---------------------------|-------------------------------|---------------------------------|--------|---------|-----------------|---------|
| ACL-7409-FT | Tequixquiac, Mexico State | spring water | 0.70476 | 484002 | 2199133 | 2239 | 1 |
| ACL-7409-UF | Tequixquiac, Mexico State | spring water | 0.70469 | 484002 | 2199133 | 2239 | 1 |
| ACL-7410-FT | Tequixquiac, Mexico State | spring water | 0.70462 | 480117 | 2200273 | 2533 | 1 |
| ACL-7410-UF | Tequixquiac, Mexico State | spring water | 0.70458 | 480117 | 2200273 | 2533 | 1 |
| TU-1S | Tula, Hidalgo | soil ^a | 0.70500 | 464348 | 2218555 | 2050 | 1 |
| TU-2S | Tula, Hidalgo | soil ^a | 0.70501 | 464348 | 2218555 | 2050 | 1 |
| TU-3S | Tula, Hidalgo | soil ^a | 0.70469 | 464348 | 2218555 | 2050 | 1 |
| ACL-9058 | Texcotzingo, Mexico State | <i>Opuntia ficus</i> | 0.70471 | 519433 | 2155797 | 2513 | 2 |
| ACL-9059 | Texcotzingo, Mexico State | <i>Dahlia pinnata</i> | 0.70459 | 519358 | 2155730 | 2504 | 2 |
| ACL-9060 | Texcotzingo, Mexico State | <i>Agave</i> spp. | 0.70464 | 519020 | 2155859 | 2534 | 2 |
| ACL-7374 | Xaltocan, Mexico State | <i>Kochia scoparia</i> | 0.70480 | 495867 | 2178713 | 2239 | 2 |
| ACL-7375 | Xaltocan, Mexico State | <i>Poa</i> spp. | 0.70479 | 495867 | 2178716 | 2239 | 2 |
| ACL-7376 | Xaltocan, Mexico State | <i>Poa</i> spp. | 0.70480 | 495878 | 2178710 | 2239 | 2 |
| ACL-7377 | Xaltocan, Mexico State | <i>Chenopodium nuttalliae</i> | 0.70482 | 495882 | 2178707 | 2239 | 2 |
| ACL-7378 | Xaltocan, Mexico State | <i>Chenopodium nuttalliae</i> | 0.70480 | 495883 | 2178708 | 2239 | 2 |

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|----------------------------|---------------------------|-------------------------------------|---------|--------|---------|------|---|
| ACL-7379 | Xaltocan, Mexico State | <i>Chenopodium nuttalliae</i> | 0.70479 | 495892 | 2178710 | 2239 | 2 |
| ACL-7380 | Xaltocan, Mexico State | <i>Avena sativa</i> | 0.70477 | 495702 | 2178935 | 2238 | 2 |
| ACL-7381 | Xaltocan, Mexico State | <i>Helianthus</i> spp. | 0.70479 | 494737 | 2178926 | 2238 | 2 |
| ACL-7382 | Xaltocan, Mexico State | <i>Avena sativa</i> | 0.70478 | 494837 | 2178943 | 2239 | 2 |
| ACL-7383 | Xaltocan, Mexico State | <i>Taraxacum officinale</i> | 0.70474 | 497854 | 2178943 | 2239 | 2 |
| ACL-7384 | Xaltocan, Mexico State | <i>Taraxacum officinale</i> | 0.70475 | 495063 | 2178959 | 2239 | 2 |
| ACL-7385 | Xaltocan, Mexico State | <i>Hordeum vulgare</i> | 0.70478 | 495346 | 2178978 | 2239 | 2 |
| ACL-7386 | Xaltocan, Mexico State | <i>Chenopodium nuttalliae</i> | 0.70481 | 495347 | 2178979 | 2239 | 2 |
| ACL-7387 | Xaltocan, Mexico State | <i>Hordeum vulgare</i> | 0.70484 | 495423 | 2178904 | 2239 | 2 |
| ACL-7388 | Xaltocan, Mexico State | <i>Poa</i> spp. | 0.70484 | 495637 | 2178998 | 2239 | 2 |
| ACL-7389 | Xaltocan, Mexico State | <i>Jaltomata procumbens</i> | 0.70497 | 495884 | 2178860 | 2239 | 2 |
| ACL-7390 | Xaltocan, Mexico State | <i>Poa</i> spp. | 0.70488 | 495868 | 2178714 | 2239 | 2 |
| ACL-7391 | Xaltocan, Mexico State | <i>Poa</i> spp. | 0.70490 | 495868 | 2178713 | 2239 | 2 |
| ACL-7394 | Xaltocan, Mexico State | <i>Helianthus</i> spp. | 0.70481 | 495846 | 2178692 | 2238 | 2 |
| ACL-7397 | Xaltocan, Mexico State | <i>Kochia scoparia</i> | 0.70482 | 495826 | 2178687 | 2238 | 2 |
| ACL-7399 | Xaltocan, Mexico State | <i>Agave</i> spp. | 0.70471 | 495251 | 2181200 | 2241 | 2 |
| ACL-7400 | Xaltocan, Mexico State | <i>Opuntia ficus</i> | 0.70490 | 495292 | 2181209 | 2242 | 2 |
| 11203 CV C2 N334 E96 11 | Teotihuacan, Mexico State | <i>Sylvilagus</i> spp. ^b | 0.70459 | 516371 | 2177462 | 2351 | 2 |
| 11145 CV C2 N331 E93 1k | Teotihuacan, Mexico State | <i>Sylvilagus</i> spp. ^b | 0.70458 | 516371 | 2177462 | 2351 | 2 |
| 3110 CV C1 N342 E94 1a | Teotihuacan, Mexico State | <i>Sylvilagus</i> spp. ^b | 0.70468 | 516371 | 2177462 | 2351 | 2 |

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|--------------------------------|------------------------------|-------------------------------------|---------|--------|---------|------|---|
| 8186 CV T N333 E81 2d | Teotihuacan, Mexico State | <i>Sylvilagus spp.</i> ^b | 0.70459 | 516371 | 2177462 | 2351 | 2 |
| 3294 CV C1 N338 E91 1a | Teotihuacan, Mexico State | <i>Sylvilagus spp.</i> ^b | 0.70464 | 516371 | 2177462 | 2351 | 2 |
| 7531 CV NS N334 E91 1a | Teotihuacan, Mexico State | <i>Sylvilagus spp.</i> ^b | 0.70471 | 516371 | 2177462 | 2351 | 2 |
| 22422 CP C5 N348 E116 1f/2a | Teotihuacan, Mexico State | <i>Sylvilagus spp.</i> ^b | 0.70461 | 516371 | 2177462 | 2351 | 2 |
| 790 CB N325 E16 S | Teotihuacan, Mexico State | <i>Sylvilagus spp.</i> ^b | 0.70470 | 516371 | 2177462 | 2351 | 2 |
| 706 CB N332 E31 S | Teotihuacan, Mexico State | <i>Sylvilagus spp.</i> ^b | 0.70465 | 516371 | 2177462 | 2351 | 2 |
| 67145s | Teotihuacan, Mexico State | soil ^a | 0.70435 | 516371 | 2177462 | 2351 | 2 |
| 67145s | Teotihuacan, Mexico State | soil ^a | 0.70432 | 516371 | 2177462 | 2351 | 2 |
| 25166s | Teotihuacan, Mexico State | soil ^a | 0.70438 | 516371 | 2177462 | 2351 | 2 |
| 25166s | Teotihuacan, Mexico State | soil ^a | 0.70441 | 516371 | 2177462 | 2351 | 2 |
| ACL-9046 | Cuicuilco, Mexico City | <i>Agave spp.</i> | 0.70507 | 480790 | 2134234 | 2290 | 3 |
| ACL-9047 | Cuicuilco, Mexico City | <i>Dahlia pinnata</i> | 0.70536 | 480998 | 2134251 | 2288 | 3 |
| ACL-9048 | Cuicuilco, Mexico City | <i>Verbascum giganteum</i> | 0.70502 | 481044 | 2134137 | 2283 | 3 |
| ACL-9049 | Cuicuilco, Mexico City | <i>Opuntia ficus</i> | 0.70591 | 480991 | 2134066 | 2286 | 3 |
| ACL-9050 | Tezozomoc, Mexico City | <i>Schinus molle</i> | 0.70520 | 477880 | 2156155 | 2251 | 3 |

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|----------|-----------------------------------|----------------------------|---------|--------|---------|------|---|
| ACL-9051 | Tezozomoc, Mexico City | <i>Agave</i> spp. | 0.70618 | 478014 | 2156268 | 2251 | 3 |
| ACL-9052 | Naucalli, Mexico State | <i>Yucca filifera</i> | 0.70497 | 474873 | 2155369 | 2264 | 3 |
| ACL-9053 | Naucalli, Mexico State | <i>Opuntia ficus</i> | 0.70503 | 475008 | 2155777 | 2264 | 3 |
| ACL-9054 | Naucalli, Mexico State | <i>Agave</i> spp. | 0.70469 | 474863 | 2155595 | 2264 | 3 |
| ACL-9055 | Cerro Moctezuma, Mexico State | <i>Arctostaphylos</i> spp. | 0.70455 | 473040 | 2154358 | 2385 | 3 |
| ACL-9056 | Cerro Moctezuma, Mexico State | <i>Agave</i> spp. | 0.70489 | 472950 | 2154408 | 2397 | 3 |
| ACL-9057 | Cerro Moctezuma, Mexico State | <i>Dahlia pinnata</i> | 0.70471 | 473033 | 2154438 | 2382 | 3 |
| ACL-9061 | Tlatelolco, Mexico City | <i>Poa</i> spp. | 0.70484 | 485501 | 2150723 | 2231 | 3 |
| ACL-9062 | Tlatelolco, Mexico City | <i>Yucca filifera</i> | 0.70483 | 485523 | 2150718 | 2231 | 3 |
| ACL-9063 | Tlatelolco, Mexico City | <i>Opuntia ficus</i> | 0.70496 | 485597 | 2150719 | 2233 | 3 |
| ACL-9064 | Tlatelolco, Mexico City | <i>Agave</i> spp. | 0.70514 | 485543 | 2150786 | 2233 | 3 |
| ACL-9069 | San Pedro Atocpan, Mexico City | <i>Opuntia ficus</i> | 0.70486 | 494693 | 2122995 | 2239 | 3 |
| ACL-9070 | San Pedro Atocpan, Mexico City | <i>Poa</i> spp. | 0.70481 | 494693 | 2122995 | 2239 | 3 |
| ACL-9071 | San Pedro Atocpan, Mexico City | <i>Amaranthus hybridus</i> | 0.70452 | 494693 | 2122995 | 2239 | 3 |
| ACL-9072 | Santiago Tulyehualco, Mexico City | <i>Pinus</i> spp. | 0.70466 | 498319 | 2128955 | 2252 | 3 |
| ACL-9073 | Santiago Tulyehualco, Mexico City | <i>Agave</i> spp. | 0.70455 | 498319 | 2128955 | 2252 | 3 |
| ACL-9074 | Santiago Tulyehualco, Mexico City | <i>Agave</i> spp. | 0.70462 | 498319 | 2128955 | 2252 | 3 |

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|----------|---------------------|---------------------------------|---------|--------|---------|------|---|
| ACL-9075 | Cholula, Puebla | <i>Dorotheanthus</i> spp. | 0.70543 | 573378 | 2107043 | 2148 | 4 |
| ACL-9076 | Cholula, Puebla | <i>Opuntia ficus</i> | 0.70598 | 573421 | 2107130 | 2150 | 4 |
| ACL-9077 | Cholula, Puebla | <i>Chenopodium nuttalliae</i> | 0.70575 | 573314 | 2107191 | 2154 | 4 |
| ACL-9078 | Cholula, Puebla | <i>Agave</i> spp. | 0.70602 | 573258 | 2107461 | 2157 | 4 |
| ACL-9079 | Cacaxtla, Tlaxcala | <i>Dahlia pinnata</i> | 0.70500 | 569529 | 2127993 | 2298 | 4 |
| ACL-9080 | Cacaxtla, Tlaxcala | <i>Agave</i> spp. | 0.70553 | 569461 | 2128024 | 2302 | 4 |
| ACL-9081 | Cacaxtla, Tlaxcala | <i>Quercus</i> spp. | 0.70525 | 569280 | 2128072 | 2305 | 4 |
| ACL-9082 | Cacaxtla, Tlaxcala | <i>Agave</i> spp. | 0.70541 | 569395 | 2127878 | 2309 | 4 |
| ACL-9065 | Xochicalco, Morelos | <i>Agave</i> spp. | 0.70641 | 468747 | 2079174 | 1349 | 5 |
| ACL-9066 | Xochicalco, Morelos | <i>Enterolobium cyclocarpum</i> | 0.70521 | 468921 | 2079148 | 1329 | 5 |
| ACL-9067 | Xochicalco, Morelos | <i>Agave</i> spp. | 0.70600 | 468629 | 2079292 | 1348 | 5 |
| ACL-9068 | Xochicalco, Morelos | <i>Agave</i> spp. | 0.70539 | 468872 | 2079236 | 1340 | 5 |

^a Data published in Schaaf et al. (2012)

^b Data published in Price et al. (2000)

APPENDIX C

BIOGEOCHEMICAL $^{87}\text{SR}/^{86}\text{SR}$ AND $\delta^{18}\text{O}$ DATA FROM NON-GRID 4

| Laboratory Number | Individual | Sex ^a | Age ^b | Element | Ca/P | U/Ca | Nd/Ca | $\delta^{18}\text{O}_{\text{carbonate}}$ (VPDB) | $^{87}\text{Sr}/^{86}\text{Sr}$ |
|-------------------|------------|------------------|------------------|------------|------|----------|----------|--|---------------------------------|
| ACL-7969 | XAL-C1-1 | M | Adult | LRM1 | 1.99 | 2.39E-07 | 5.05E-07 | -6.6 | 0.70564 |
| ACL-7970 | XAL-C1-1 | M | Adult | URM3 | 1.97 | 1.71E-07 | 5.49E-07 | -6.1 | 0.70558 |
| ACL-7971 | XAL-C1-1 | M | Adult | L sphenoid | 2.23 | 9.04E-06 | 7.54E-07 | -5.3 | 0.70478 |
| ACL-7972 | XAL-C1-2 | M | Adult | LLM1 | 1.96 | 1.21E-07 | 3.67E-07 | -5.9 | 0.70608 |
| ACL-7973 | XAL-C1-2 | M | Adult | LLM3 | 1.98 | 3.68E-07 | 2.15E-07 | -6.3 | 0.70555 |
| ACL-7974 | XAL-C1-2 | M | Adult | occipital | 2.29 | 3.31E-05 | 3.78E-07 | -7.4 | 0.70483 |
| ACL-7975 | XAL-C1-4 | M | Adult | LLM1 | 1.98 | 2.73E-07 | 3.46E-07 | -6.1 | 0.70466 |
| ACL-7976 | XAL-C1-4 | M | Adult | URM3 | 1.96 | 1.46E-07 | 1.49E-07 | -7.2 | 0.70596 |
| ACL-7977 | XAL-C1-4 | M | Adult | L mandible | 2.30 | 3.13E-05 | 2.27E-07 | -8.0 | 0.70471 |
| ACL-7978 | XAL-C2-1 | M | Adult | ULM1 | 1.96 | 2.12E-07 | 2.19E-07 | -6.2 | 0.70555 |
| ACL-7979 | XAL-C2-1 | M | Adult | ULM3 | 1.99 | 3.34E-07 | 1.71E-07 | -7.1 | 0.70541 |
| ACL-7980 | XAL-C2-2 | M | Adult | LRM1 | 1.95 | 1.42E-07 | 1.38E-07 | -6.1 | 0.70451 |
| ACL-7981 | XAL-C2-2 | M | Adult | URM3 | 2.00 | 1.07E-07 | 1.30E-07 | -7.1 | 0.70525 |
| ACL-7982 | XAL-C2-2 | M | Adult | sphenoid | 2.27 | 8.80E-06 | 3.43E-07 | -7.0 | 0.70477 |
| ACL-7983 | XAL-C2-3 | M | Adult | LRM1 | 1.98 | 2.37E-07 | 1.24E-07 | -5.9 | 0.70530 |
| ACL-7984 | XAL-C2-3 | M | Adult | LRM3 | 1.98 | 7.33E-08 | 1.41E-07 | -7.3 | 0.70530 |
| ACL-7985 | XAL-C3-3 | M | Adult | ULM1 | 1.96 | 3.22E-08 | 1.79E-07 | -6.2 | 0.70387 |
| ACL-7986 | XAL-C3-3 | M | Adult | ULM3 | 1.95 | 3.43E-07 | 6.18E-07 | -6.5 | 0.70448 |
| ACL-7987 | XAL-C4-1 | M | Adol. | LLM1 | 1.96 | 1.48E-07 | 2.10E-07 | -7.5 | 0.70602 |
| ACL-7988 | XAL-C4-1 | M | Adol. | LLM3 | 1.97 | 7.53E-08 | 2.11E-07 | -7.0 | 0.70584 |
| ACL-7989 | XAL-C4-1 | M | Adol. | L temporal | 2.29 | 3.41E-05 | 2.17E-06 | -7.4 | 0.70481 |
| ACL-7990 | XAL-C4-2 | M | Adult | LLM1 | 1.98 | 3.40E-07 | 4.48E-07 | -6.1 | 0.70525 |
| ACL-7991 | XAL-C4-2 | M | Adult | LLM3 | 1.99 | 1.97E-07 | 9.25E-08 | -8.6 | 0.70530 |
| ACL-7992 | XAL-C4-2 | M | Adult | L mandible | 2.26 | 2.57E-05 | 6.92E-07 | -8.4 | 0.70479 |
| ACL-7993 | XAL-C4-4 | M | Adult | LRM1 | 1.97 | 1.89E-07 | 5.95E-07 | -5.5 | 0.70719 |

| | | | | | | | | | |
|----------|-----------|------|-------|-------------|------|----------|----------|-------|---------|
| ACL-7994 | XAL-C4-4 | M | Adult | LRM3 | 1.97 | 1.63E-07 | 6.37E-07 | -6.4 | 0.70703 |
| ACL-7995 | XAL-C4-4 | M | Adult | R temporal | 2.29 | 6.97E-05 | 2.81E-07 | -11.3 | 0.70499 |
| ACL-7996 | XAL-C5-2 | M | Adult | URM1 | 1.98 | 6.74E-08 | 1.36E-07 | -5.3 | 0.70598 |
| ACL-7997 | XAL-C5-2 | M | Adult | ULM3 | 1.97 | 1.36E-07 | 1.25E-07 | -5.2 | 0.70540 |
| ACL-7998 | XAL-C5-2 | M | Adult | R mandible | 2.23 | 8.82E-06 | 1.91E-07 | -6.3 | 0.70485 |
| ACL-7999 | XAL-C5-3 | M | Adult | LRM1 | 1.96 | 9.20E-08 | 1.01E-07 | -5.8 | 0.70472 |
| ACL-8000 | XAL-C5-3 | M | Adult | ULM3 | 1.97 | 1.14E-07 | 1.28E-07 | -5.8 | 0.70477 |
| ACL-8001 | XAL-C8-2 | M | Adult | ULM1 | 1.96 | 1.76E-07 | 4.50E-07 | -6.1 | 0.70475 |
| ACL-8002 | XAL-C8-2 | M | Adult | LLM3 | 1.99 | 1.92E-07 | 4.29E-07 | -6.3 | 0.70464 |
| ACL-8003 | XAL-C8-2 | M | Adult | L temporal | 2.26 | 2.69E-05 | 5.86E-07 | -4.5 | 0.70468 |
| ACL-8004 | XAL-C8-3 | M | Adult | ULM1 | 1.97 | 6.61E-08 | 1.11E-07 | -6.7 | 0.70602 |
| ACL-8005 | XAL-C8-3 | M | Adult | URM3 | 1.95 | 4.22E-08 | 1.22E-07 | -6.5 | 0.70603 |
| ACL-8006 | XAL-C9-1 | F | Adult | LLM1 | 1.96 | 7.02E-08 | 1.18E-07 | -6.6 | 0.70416 |
| ACL-8007 | XAL-C9-1 | F | Adult | ULM3 | 1.97 | 6.98E-08 | 1.69E-07 | -7.3 | 0.70408 |
| ACL-8008 | XAL-C9-1 | F | Adult | R parietal | 2.23 | 1.30E-05 | 6.66E-07 | -6.5 | 0.70472 |
| ACL-8009 | XAL-C9-3 | Ind. | Adol. | URM1 | 1.98 | 2.68E-07 | 6.46E-07 | -6.0 | 0.70471 |
| ACL-8010 | XAL-C9-3 | Ind. | Adol. | URM3 | 1.97 | 1.23E-07 | 1.69E-07 | -5.5 | 0.70484 |
| ACL-8011 | XAL-C9-3 | Ind. | Adol. | R maxilla | 2.25 | 2.37E-05 | 3.07E-07 | -4.9 | 0.70477 |
| ACL-8012 | XAL-C9-4 | M | Adult | LRM1 | 1.95 | 1.11E-07 | 3.30E-07 | -5.7 | 0.70469 |
| ACL-8013 | XAL-C9-4 | M | Adult | LRM3 | 1.95 | 7.69E-08 | 3.94E-07 | -8.0 | 0.70464 |
| ACL-8014 | XAL-C9-4 | M | Adult | R zygomatic | 2.25 | 2.00E-05 | 4.94E-07 | -7.5 | 0.70473 |
| ACL-8015 | XAL-C10-4 | M | Adult | URM1 | 1.99 | 5.20E-07 | 7.37E-07 | -5.5 | 0.70451 |
| ACL-8016 | XAL-C10-4 | M | Adult | URM3 | 1.98 | 5.74E-07 | 9.42E-07 | -6.5 | 0.70428 |
| ACL-8017 | XAL-C12-2 | M | Adult | URM1 | 2.04 | 1.73E-07 | 4.22E-07 | -5.0 | 0.70443 |
| ACL-8018 | XAL-C12-2 | M | Adult | URM3 | 2.04 | 1.56E-07 | 4.30E-07 | -5.9 | 0.70423 |
| ACL-8019 | XAL-C12-2 | M | Adult | R temporal | 2.33 | 2.62E-05 | 1.10E-07 | -5.7 | 0.70472 |
| ACL-8020 | XAL-C13-1 | M | Adult | LRM1 | 1.98 | 1.16E-07 | 4.76E-08 | -4.4 | 0.70507 |

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| ACL-8021 | XAL-C13-1 | M | Adult | LRM3 | 1.98 | 7.47E-08 | 4.04E-08 | -5.4 | 0.70493 |
| ACL-8022 | XAL-C13-1 | M | Adult | R sphenoid | 2.20 | 2.92E-05 | 2.75E-07 | -3.9 | 0.70477 |
| ACL-8023 | XAL-C13-3 | M | Adult | LRM1 | 1.96 | 1.44E-07 | 4.23E-08 | -5.8 | 0.70465 |
| ACL-8024 | XAL-C13-3 | M | Adult | URM3 | 1.96 | 1.45E-07 | 4.84E-08 | -7.0 | 0.70476 |
| ACL-8025 | XAL-C13-4 | M | Adult | URM1 | 1.96 | 2.78E-07 | 9.85E-08 | -6.2 | 0.70537 |
| ACL-8026 | XAL-C13-4 | M | Adult | ULM3 | 1.97 | 6.78E-08 | 5.58E-08 | -7.0 | 0.70499 |
| ACL-8027 | XAL-C13-4 | M | Adult | occipital | 2.24 | 1.60E-05 | 4.15E-07 | -7.0 | 0.70475 |
| ACL-8028 | XAL-C14-1 | Ind. | Adult | LRM1 | 1.95 | 4.16E-07 | 7.58E-08 | -5.9 | 0.70451 |
| ACL-8029 | XAL-C14-1 | Ind. | Adult | LRM3 | 1.96 | 2.15E-07 | 9.64E-08 | -6.6 | 0.70561 |
| ACL-8030 | XAL-C14-5 | M | Adult | URM1 | 1.96 | 3.38E-07 | 2.64E-07 | -6.9 | 0.70432 |
| ACL-8031 | XAL-C14-5 | M | Adult | URM3 | 1.96 | 2.73E-07 | 1.45E-07 | -8.4 | 0.70434 |
| ACL-8032 | XAL-C15-3 | M | Adult | LRM1 | 1.97 | 1.65E-07 | 6.33E-08 | -5.3 | 0.70448 |
| ACL-8033 | XAL-C15-3 | M | Adult | LRM3 | 1.94 | 1.58E-07 | 7.43E-08 | -5.4 | 0.70471 |
| ACL-8034 | XAL-C15-3 | M | Adult | occipital | 2.23 | 1.35E-05 | 5.06E-07 | -7.0 | 0.70476 |
| ACL-8035 | XAL-C15-4 | M | Adol. | URM1 | 1.97 | 4.93E-07 | 1.21E-07 | -6.3 | 0.70454 |
| ACL-8036 | XAL-C15-4 | M | Adol. | URM3 | 1.97 | 1.45E-07 | 5.03E-08 | -5.3 | 0.70466 |
| ACL-8037 | XAL-C15-6 | M | Adult | ULM1 | 1.99 | 1.69E-07 | 5.00E-08 | -5.7 | 0.70463 |
| ACL-8038 | XAL-C15-6 | M | Adult | ULM3 | 1.96 | 8.69E-08 | 4.24E-08 | -6.7 | 0.70503 |
| ACL-8039 | XAL-C16-1 | M | Adult | LLM1 | 1.98 | 1.09E-07 | 1.95E-07 | -6.8 | 0.70463 |
| ACL-8040 | XAL-C16-1 | M | Adult | LLM3 | 1.98 | 2.03E-07 | 4.95E-07 | -7.5 | 0.70453 |
| ACL-8041 | XAL-C16-2 | M | Adult | URM1 | 1.97 | 2.33E-07 | 1.60E-07 | -6.0 | 0.70552 |
| ACL-8042 | XAL-C16-2 | M | Adult | URM3 | 1.97 | 8.60E-08 | 5.17E-08 | -6.2 | 0.70591 |
| ACL-8043 | XAL-C16-3 | M | Adult | LRM1 | 1.98 | 2.25E-07 | 1.65E-07 | -6.1 | 0.70473 |
| ACL-8044 | XAL-C16-3 | M | Adult | ULM3 | 1.96 | 1.94E-07 | 2.89E-07 | -6.9 | 0.70473 |
| ACL-8045 | XAL-C16-4 | F | Adult | LRM1 | 1.99 | 1.33E-07 | 4.05E-08 | -6.7 | 0.70550 |
| ACL-8046 | XAL-C16-4 | F | Adult | URM3 | 2.06 | 1.55E-07 | 4.50E-08 | -7.7 | 0.70552 |
| ACL-8047 | XAL-C16-5 | M | Adol. | URM1 | 2.05 | 8.26E-08 | 3.60E-08 | -6.2 | 0.70550 |

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| ACL-8048 | XAL-C16-5 | M | Adol. | URM3 | 2.03 | 7.54E-08 | 4.30E-08 | -7.4 | 0.70559 |
| ACL-8049 | XAL-C16-5 | M | Adol. | R sphenoid | 2.21 | 1.25E-05 | 1.70E-07 | -6.0 | 0.70480 |
| ACL-8050 | XAL-C17-1 | Ind. | Adol. | LLM1 | 1.96 | 6.07E-08 | 4.30E-08 | -6.5 | 0.70591 |
| ACL-8051 | XAL-C17-1 | Ind. | Adol. | LLM3 | 1.98 | 1.17E-07 | 4.30E-08 | -6.6 | 0.70575 |
| ACL-8052 | XAL-C17-1 | Ind. | Adol. | occipital | 2.21 | 3.01E-05 | 3.95E-07 | -0.4 | 0.70483 |
| ACL-8053 | XAL-C17-4 | M | Adult | LRM1 | 1.97 | 1.74E-07 | 2.31E-08 | -5.2 | 0.70472 |
| ACL-8054 | XAL-C17-4 | M | Adult | ULM3 | 1.95 | 1.44E-07 | 2.24E-08 | -5.1 | 0.70459 |
| ACL-8055 | XAL-C17-5 | F | Adol. | ULM1 | 1.96 | 1.95E-07 | 5.13E-08 | -3.2 | 0.70595 |
| ACL-8056 | XAL-C17-5 | F | Adol. | ULM3 | 1.96 | 1.37E-07 | 3.72E-08 | -3.6 | 0.70698 |
| ACL-8057 | XAL-C17-5 | F | Adol. | R zygomatic | 2.26 | 3.42E-05 | 1.14E-06 | -5.0 | 0.70478 |
| ACL-8058 | XAL-C17-6 | M | Adol. | LLM1 | 1.95 | 9.84E-08 | 3.82E-08 | -5.9 | 0.70463 |
| ACL-8059 | XAL-C17-6 | M | Adol. | LLM3 | 1.96 | 7.49E-08 | 3.16E-08 | -6.7 | 0.70452 |
| ACL-8060 | XAL-C17-6 | M | Adol. | L mandible | 2.25 | 4.03E-05 | 1.61E-07 | -5.8 | 0.70468 |
| ACL-8061 | XAL-C18-1 | M | Adol. | URM1 | 1.99 | 2.11E-07 | 1.02E-07 | -6.7 | 0.70402 |
| ACL-8062 | XAL-C18-1 | M | Adol. | URM3 | 1.98 | 9.50E-08 | 6.22E-08 | -7.8 | 0.70419 |
| ACL-8063 | XAL-C18-1 | M | Adol. | L frontal | 2.15 | 2.07E-05 | 1.78E-07 | -6.8 | 0.70465 |
| ACL-8064 | XAL-C18-3 | M | Adult | URM1 | 1.96 | 1.61E-07 | 7.38E-08 | -5.9 | 0.70579 |
| ACL-8065 | XAL-C18-3 | M | Adult | URM3 | 1.95 | 5.91E-08 | 3.67E-08 | -5.8 | 0.70525 |
| ACL-8066 | XAL-C18-3 | M | Adult | R maxilla | 2.19 | 1.54E-05 | 7.13E-07 | -4.0 | 0.70476 |
| ACL-8067 | XAL-C19-1 | M | Adult | URM1 | 1.94 | 1.24E-07 | 6.79E-08 | -5.0 | 0.70478 |
| ACL-8068 | XAL-C19-1 | M | Adult | ULM3 | 1.96 | 2.09E-07 | 5.85E-07 | -7.0 | 0.70457 |
| ACL-8069 | XAL-C19-1 | M | Adult | R occipital | 2.17 | 4.32E-05 | 2.99E-07 | -5.8 | 0.70476 |
| ACL-8070 | XAL-C19-6 | M | Adult | URM1 | 1.97 | 1.13E-07 | 1.27E-07 | -5.8 | 0.70539 |
| ACL-8071 | XAL-C19-6 | M | Adult | ULM3 | 1.96 | 1.42E-07 | 4.06E-07 | -7.1 | 0.70564 |
| ACL-8072 | XAL-C21-3 | M | Adult | URM1 | 1.97 | 3.26E-07 | 2.03E-07 | -4.9 | 0.70448 |
| ACL-8073 | XAL-C21-3 | M | Adult | ULM3 | 1.95 | 9.36E-08 | 1.56E-07 | -6.9 | 0.70475 |
| ACL-8074 | XAL-C22-1 | M | Adult | URM1 | 1.96 | 2.97E-07 | 4.96E-07 | -6.4 | 0.70605 |

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| ACL-8075 | XAL-C22-1 | M | Adult | URM3 | 1.96 | 9.90E-08 | 9.52E-08 | -6.8 | 0.70636 |
| ACL-8076 | XAL-C22-3 | M | Adult | ULM1 | 1.97 | 1.06E-06 | 3.41E-07 | -5.5 | 0.70459 |
| ACL-8077 | XAL-C22-3 | M | Adult | LRM3 | 1.96 | 8.25E-08 | 1.38E-07 | -5.3 | 0.70445 |
| ACL-8078 | XAL-C22-4 | M | Adult | LRM1 | 1.97 | 4.92E-07 | 1.01E-07 | -6.6 | 0.70519 |
| ACL-8079 | XAL-C22-4 | M | Adult | LRM3 | 1.97 | 1.18E-07 | 1.07E-07 | -7.9 | 0.70529 |
| ACL-8080 | XAL-C22-4 | M | Adult | R zygomatic | 2.27 | 4.01E-05 | 1.53E-07 | -5.5 | 0.70477 |
| ACL-8081 | XAL-C23-2 | M | Adult | LRM1 | 1.96 | 2.03E-07 | 1.83E-07 | -5.6 | 0.70539 |
| ACL-8082 | XAL-C23-2 | M | Adult | LRM3 | 1.97 | 3.49E-07 | 1.01E-07 | -6.9 | 0.70525 |
| ACL-8083 | XAL-C23-2 | M | Adult | R mandible | 2.26 | 3.39E-05 | 6.59E-08 | -6.3 | 0.70476 |
| ACL-8084 | XAL-C23-5 | M | Adult | ULM1 | 1.96 | 1.70E-07 | 5.15E-07 | -5.6 | 0.70450 |
| ACL-8085 | XAL-C23-5 | M | Adult | URM3 | 1.96 | 4.99E-08 | 4.85E-08 | -6.9 | 0.70448 |
| ACL-8086 | XAL-C24-2 | M | Adult | URM1 | 1.97 | 1.21E-06 | 2.28E-07 | -5.0 | 0.70546 |
| ACL-8087 | XAL-C24-2 | M | Adult | URM3 | 1.96 | 7.97E-08 | 2.09E-07 | -6.7 | 0.70573 |
| ACL-8088 | XAL-C24-4 | F | Adol. | LLM1 | 1.97 | 6.35E-07 | 2.91E-07 | -7.2 | 0.70546 |
| ACL-8089 | XAL-C24-4 | F | Adol. | ULM3 | 2.05 | 6.88E-06 | 3.02E-07 | -7.0 | 0.70486 |
| ACL-8090 | XAL-C24-4 | F | Adol. | R zygomatic | 2.18 | 4.04E-05 | 1.59E-07 | -5.1 | 0.70476 |
| ACL-8091 | XAL-C24-5 | M | Adol. | ULM1 | 1.95 | 5.32E-08 | 5.19E-08 | -6.0 | 0.70531 |
| ACL-8092 | XAL-C24-5 | M | Adol. | ULM3 | 1.97 | 6.20E-08 | 5.48E-08 | -6.1 | 0.70523 |
| ACL-8093 | XAL-C24-5 | M | Adol. | L sphenoid | 2.25 | 2.24E-05 | 7.56E-07 | -3.7 | 0.70481 |
| ACL-8094 | XAL-C24-6 | M | Adult | ULM1 | 1.97 | 6.01E-07 | 5.46E-07 | -7.2 | 0.70632 |
| ACL-8095 | XAL-C24-6 | M | Adult | ULM3 | 1.96 | 1.55E-07 | 1.05E-07 | -6.9 | 0.70738 |
| ACL-8096 | XAL-C25-2 | M | Adult | LRM1 | 1.96 | 1.98E-07 | 1.29E-07 | -4.5 | 0.70464 |
| ACL-8097 | XAL-C25-2 | M | Adult | URM3 | 1.96 | 1.83E-07 | 1.93E-07 | -5.7 | 0.70472 |
| ACL-8098 | XAL-C25-3 | Ind. | Adol. | ULM1 | 1.96 | 3.29E-07 | 2.33E-07 | -6.2 | 0.70482 |
| ACL-8099 | XAL-C25-3 | Ind. | Adol. | ULM3 | 1.95 | 9.33E-08 | 4.01E-08 | -6.3 | 0.70493 |
| ACL-8100 | XAL-C25-3 | Ind. | Adol. | L sphenoid | 2.21 | 2.02E-05 | 4.22E-07 | -6.0 | 0.70477 |
| ACL-8101 | XAL-C25-4 | M | Adult | URM1 | 1.96 | 2.71E-07 | 2.80E-07 | -5.9 | 0.70473 |

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| ACL-8102 | XAL-C25-4 | M | Adult | ULM3 | 1.96 | 2.42E-07 | 1.15E-07 | -6.5 | 0.70462 |
| ACL-8103 | XAL-C25-4 | M | Adult | R & L parietal | 2.07 | 2.57E-05 | 3.03E-07 | -5.9 | 0.70473 |
| ACL-8104 | XAL-C25-5 | M | Adult | URM1 | 1.96 | 1.46E-06 | 4.26E-07 | -4.0 | 0.70483 |
| ACL-8105 | XAL-C25-5 | M | Adult | ULM3 | 1.97 | 6.99E-07 | 1.28E-07 | -4.2 | 0.70491 |
| ACL-8106 | XAL-C25-5 | M | Adult | R maxilla | 2.24 | 4.00E-05 | 1.58E-07 | -5.3 | 0.70477 |
| ACL-8107 | XAL-C26-1 | M | Adult | URM1 | 1.96 | 2.92E-07 | 1.48E-07 | -6.9 | 0.70438 |
| ACL-8108 | XAL-C26-1 | M | Adult | ULM3 | 1.97 | 1.32E-07 | 2.08E-07 | -5.8 | 0.70446 |
| ACL-8109 | XAL-C26-3 | M | Adult | URM1 | 1.97 | 2.84E-07 | 7.63E-07 | -6.1 | 0.70438 |
| ACL-8110 | XAL-C26-3 | M | Adult | LRM3 | 1.97 | 6.51E-08 | 7.92E-08 | -5.7 | 0.70432 |
| ACL-8111 | XAL-C26-3 | M | Adult | R zygomatic | 1.97 | 9.78E-06 | NA | -6.4 | 0.70468 |
| ACL-8112 | XAL-C26-4 | M | Adult | ULM1 | 1.96 | 2.00E-07 | 1.99E-07 | -7.1 | 0.70447 |
| ACL-8113 | XAL-C26-4 | M | Adult | ULM3 | 1.96 | 2.41E-07 | 2.65E-07 | -6.8 | 0.70439 |
| ACL-8114 | XAL-C26-4 | M | Adult | L zygomatic | 2.29 | 6.20E-05 | 6.31E-07 | -5.6 | 0.70472 |
| ACL-8115 | XAL-C27-2 | M | Adol. | LLM1 | 1.97 | 1.04E-07 | 8.30E-08 | -6.0 | 0.70494 |
| ACL-8116 | XAL-C27-2 | M | Adol. | URM3 | 1.97 | 1.44E-07 | 5.80E-08 | -6.8 | 0.70460 |
| ACL-8117 | XAL-C27-2 | M | Adol. | R parietal | 2.19 | 3.62E-05 | 2.09E-07 | -6.6 | 0.70472 |
| ACL-8118 | XAL-C27-4 | M | Adult | URM1 | 1.96 | 1.32E-07 | 4.82E-07 | -6.3 | 0.70443 |
| ACL-8119 | XAL-C27-4 | M | Adult | LRM3 | 1.99 | 4.79E-07 | 4.64E-07 | -6.8 | 0.70451 |
| ACL-8120 | XAL-C27-4 | M | Adult | L mandible | 2.25 | 2.37E-05 | 1.17E-07 | -5.7 | 0.70472 |
| ACL-8121 | XAL-C27-5 | M | Adult | ULM1 | 1.95 | 1.91E-07 | 4.10E-07 | -4.9 | 0.70481 |
| ACL-8122 | XAL-C27-5 | M | Adult | URM3 | 1.97 | 1.74E-07 | 5.49E-07 | -7.8 | 0.70464 |
| ACL-8123 | XAL-C27-5 | M | Adult | L parietal | 2.19 | 3.30E-05 | 7.92E-08 | -5.8 | 0.70474 |
| ACL-8124 | XAL-C28-2 | M | Adol. | URM1 | 1.98 | 1.71E-07 | 1.16E-07 | -4.9 | 0.70595 |
| ACL-8125 | XAL-C28-2 | M | Adol. | ULM3 | 2.07 | 7.00E-06 | 1.76E-06 | -4.8 | 0.70510 |
| ACL-8126 | XAL-C28-2 | M | Adol. | R parietal | 2.20 | 3.35E-05 | 2.09E-07 | -5.2 | 0.70489 |
| ACL-8127 | XAL-C28-3 | M | Adult | LRM1 | 1.97 | 1.10E-07 | 1.21E-07 | -4.1 | 0.70599 |
| ACL-8128 | XAL-C28-3 | M | Adult | LLM3 | 1.99 | 4.95E-07 | 1.59E-06 | -6.3 | 0.70627 |

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| ACL-8129 | XAL-C28-3 | M | Adult | L zygomatic | 2.22 | 3.68E-05 | 2.80E-07 | -4.8 | 0.70482 |
| ACL-8130 | XAL-C28-4 | M | Adult | URM1 | 1.97 | 2.24E-07 | 5.84E-07 | -5.7 | 0.70459 |
| ACL-8131 | XAL-C28-4 | M | Adult | URM3 | 1.96 | 4.76E-07 | 5.97E-07 | -6.4 | 0.70459 |
| ACL-8132 | XAL-C28-4 | M | Adult | L zygomatic | 2.21 | 2.37E-05 | 8.39E-07 | -5.3 | 0.70475 |
| ACL-8133 | XAL-C29-1 | F | Adol. | LRM1 | 1.95 | 6.53E-08 | 3.87E-08 | -5.2 | 0.70632 |
| ACL-8134 | XAL-C29-1 | F | Adol. | LRM3 | 1.97 | 2.05E-07 | 8.24E-08 | -6.6 | 0.70522 |
| ACL-8135 | XAL-C29-1 | F | Adol. | R mandible | 2.22 | 1.25E-05 | 2.51E-07 | -5.7 | 0.70485 |
| ACL-8136 | XAL-C29-3 | M | Adult | LRM1 | 1.96 | 5.85E-07 | 2.74E-07 | -4.9 | 0.70462 |
| ACL-8137 | XAL-C29-3 | M | Adult | LRM3 | 1.97 | 1.42E-06 | 1.65E-06 | -7.2 | 0.70464 |
| ACL-8138 | XAL-C29-3 | M | Adult | L parietal | 2.21 | 1.85E-05 | 1.02E-06 | -6.6 | 0.70474 |
| ACL-8139 | XAL-C30-2 | M | Adult | ULM1 | 1.96 | 6.58E-08 | 7.19E-08 | -6.2 | 0.70620 |
| ACL-8140 | XAL-C30-2 | M | Adult | ULM3 | 1.96 | 3.28E-07 | 6.66E-07 | -7.1 | 0.70454 |
| ACL-8141 | XAL-C30-2 | M | Adult | R occipital | 2.29 | 5.29E-06 | 2.22E-07 | -3.8 | 0.70482 |
| ACL-8142 | XAL-C30-5 | M | Adult | LLM1 | 1.96 | 3.65E-07 | 7.16E-07 | -4.9 | 0.70494 |
| ACL-8143 | XAL-C30-5 | M | Adult | LLM3 | 1.98 | 7.19E-07 | 2.49E-07 | -5.0 | 0.70484 |
| ACL-8144 | XAL-C30-5 | M | Adult | R parietal | 2.26 | 1.24E-05 | 1.49E-07 | -5.5 | 0.70476 |
| ACL-8145 | XAL-C31-1 | F | Adol. | URM1 | 1.97 | 2.46E-07 | 5.50E-07 | -4.8 | 0.70526 |
| ACL-8146 | XAL-C31-1 | F | Adol. | URM3 | 1.95 | 1.82E-07 | 7.08E-07 | -5.4 | 0.70556 |
| ACL-8147 | XAL-C31-1 | F | Adol. | R frontal | 2.23 | 2.11E-05 | 1.29E-07 | -4.5 | 0.70479 |
| ACL-8148 | XAL-C32-4 | M | Adult | LLM1 | 1.97 | 1.49E-07 | 5.31E-08 | -6.1 | 0.70452 |
| ACL-8149 | XAL-C32-4 | M | Adult | ULM3 | 1.94 | 5.54E-08 | 6.04E-08 | -6.1 | 0.70432 |
| ACL-8150 | XAL-C32-4 | M | Adult | R occipital | 2.23 | 3.17E-05 | 6.86E-07 | -5.8 | 0.70471 |
| ACL-8151 | XAL-C32-5 | M | Adult | URM1 | 1.94 | 1.93E-07 | 5.63E-07 | -4.9 | 0.70622 |
| ACL-8152 | XAL-C32-5 | M | Adult | LRM3 | 1.95 | 3.56E-07 | 4.34E-07 | -6.7 | 0.70617 |
| ACL-8153 | XAL-C32-5 | M | Adult | R parietal | 2.16 | 2.73E-05 | 1.93E-07 | -7.2 | 0.70485 |
| ACL-8154 | XAL-C34-1 | M | Adult | URM1 | 1.98 | 1.19E-06 | 5.25E-07 | -5.0 | 0.70465 |
| ACL-8155 | XAL-C34-1 | M | Adult | URM3 | 1.98 | 3.57E-07 | 2.58E-07 | -6.2 | 0.70453 |

| | | | | | | | | | |
|----------|-----------|---|-------|------------|------|----------|----------|------|---------|
| ACL-8156 | XAL-C34-1 | M | Adult | L frontal | 2.20 | 2.01E-05 | 2.62E-07 | -6.4 | 0.70472 |
| ACL-8157 | XAL-C34-2 | M | Adult | LLM1 | 1.94 | 2.09E-07 | 8.24E-08 | -6.4 | 0.70455 |
| ACL-8158 | XAL-C34-2 | M | Adult | LRM3 | 2.04 | 6.73E-06 | 1.38E-07 | -4.2 | 0.70474 |
| ACL-8159 | XAL-C34-2 | M | Adult | R parietal | 2.19 | 2.46E-05 | 9.23E-08 | -5.7 | 0.70476 |
| ACL-8160 | XAL-C34-3 | M | Adult | LRM1 | 1.96 | 5.26E-07 | 3.26E-07 | -5.6 | 0.70450 |
| ACL-8161 | XAL-C34-3 | M | Adult | LLM3 | 1.95 | 3.14E-07 | 1.22E-07 | -6.6 | 0.70448 |
| ACL-8162 | XAL-C34-3 | M | Adult | R parietal | 2.23 | 2.79E-05 | 1.53E-07 | -6.2 | 0.70472 |

^a Sex estimates based on cranial morphology following Buikstra and Ubelaker (1994). M = male; F = female; Ind = Indeterminate

^b Age estimates based on dental eruption following Buikstra and Ubelaker (1994). Adol = Adolescent